

The impact of a Neogene basalt intrusion on the optical properties and internal structure of the dispersed organic matter in Carboniferous strata (SW-part USCB)

ZDZISŁAW ADAMCZYK¹, MAGDALENA KOKOWSKA-PAWŁOWSKA¹, JOANNA KOMOREK¹,
AGNIESZKA KLUPA², MAŁGORZATA LEWANDOWSKA¹ and JACEK NOWAK¹

¹ Silesian University of Technology, Faculty of Mining and Geology, Department of Applied Geology,
Akademicka 2 Street, 44-100 Gliwice, Poland.

E-mails: zdzislaw.adamczyk@polsl.pl; magdalena.kokowska-pawlowska@polsl.pl; joanna.komorek@polsl.pl;
malgorzata.lewandowska@polsl.pl; jacek.nowak@polsl.pl

² Central Mining Institute, Plac Gwarków 140-166 Katowice, Poland.
E-mail: aklupa@gig.eu

ABSTRACT

Adamczyk, Z., Kokowska-Pawłowska, M., Komorek, J., Klupa, A., Lewandowska, M. and Nowak, J. 2018. The impact of a Neogene basalt intrusion on the optical properties and internal structure of the dispersed organic matter in Carboniferous strata (SW-part USCB). *Acta Geologica Polonica*, **68** (2), 249–262. Warszawa.

The S-7 borehole log from the Sumina area (USCB Poland) revealed the presence of three basaltic veins originating from a basalt dyke. Coal interlayers in the rocks surrounding the basaltic veins have been coked to form natural coke. Photometric measurements revealed that the optical properties of the studied natural coke samples are characteristic of semi-graphite ($R_{\max} > 9\%$). The natural coke matrix of all of the analyzed samples has a biaxial negative optical character. Vitrinite in the examined natural coke samples is characterized by a lower optical anisotropy than that of the natural matrix and it has a biaxial positive optical character. Vitrinite in almost all samples taken at locations more distant from the intrusion has a biaxial positive optical character. A reversal of the changes of the true maximum vitrinite reflectance and bireflectance with changing distance from the second basaltic vein has been observed. The temperature regime that acted upon the dispersed organic matter located in the immediate vicinity of the intrusion, estimated on the basis of the selected experimental data, is suggested to be higher than 750 °C.

Key words: Optical properties; Natural coke; Neogenic intrusion; Carbonisation.

INTRODUCTION

With increasing rank (degree of coalification) of coal, its chemical, processing, and physical properties resulting from the degree of internal structural order of the coal matter change with some regularity.

Three factors that affect the coalification process are distinguished: rise of temperature, pressure, and geologic time. As opposed to other coal macerals, changes in vitrinite reflectance during the coalifi-

cation process proceed in a uniform manner, and for this reason among others the optical properties of vitrinite have been adopted in petrography as a maturity parameter (Ergun and McCartney 1960; Stach *et al.* 1982; Van Krevelen 1993; Taylor *et al.* 1998). Reflectance measurements in incident polarized white light can show, that vitrinite in coals is optically anisotropic (Vries *et al.* 1968; Hevia and Virgos 1977; Davis 1978; Hower and Davis 1981a, b; Grieve 1991; Komorek 1996; Littke *et al.* 2012).

Changes in vitrinite reflectance with changing orientation of the polished section are described by means of a triaxial ellipsoid (indicatrix), the axis of which, in any direction, is proportional to reflectance. The three main principal reflectance axes of the indicatrix correspond to the maximum (R_{\max}), intermediate (R_{int}), and minimum (R_{\min}) reflectance values (Hevia and Virgos 1977). The values thereof are associated with the presence of a directional stress field during the coalification process of the peat-precursors and/or during structural deformation (Vries *et al.* 1968; Cook *et al.* 1972; Hower and Davis 1981a, b; Levine and Davis 1984; Houseknecht and Weesner 1997; Bruns and Littke 2015). On the basis of the shape of the indicatrix (optical character of vitrinite) and the orientation of its main axes, conclusions can be drawn related to the tectonic influence of the coal-bearing basin and coalification process (Stone and Cook 1979; Hower and Davis 1981a, b; Levine and Davis 1984; Kilby 1986, 1988; Levine and Davis 1989a, 1989b; Levine and Davis 1990; Grieve 1991; Kilby 1991; Langerberg and Kalkreuth 1991, 1991a; Middleton 1991; Reinhardt 1991; Tsai 1991; Houseknecht and Weesner 1997; Littke *et al.* 2012; Bruns and Littke 2015). The relationships between the axes define the reflectance character of vitrinite:

- isotropic $R_{\max} = R_{\text{int}} = R_{\min}$,
- uniaxial negative $R_{\max} = R_{\text{int}} > R_{\min}$,
- uniaxial positive $R_{\max} > R_{\text{int}} = R_{\min}$,
- biaxial negative $R_{\max} > R_{\text{int}} \gg R_{\min}$,
- biaxial positive $R_{\max} \gg R_{\text{int}} > R_{\min}$ (Stone and Cook 1979).

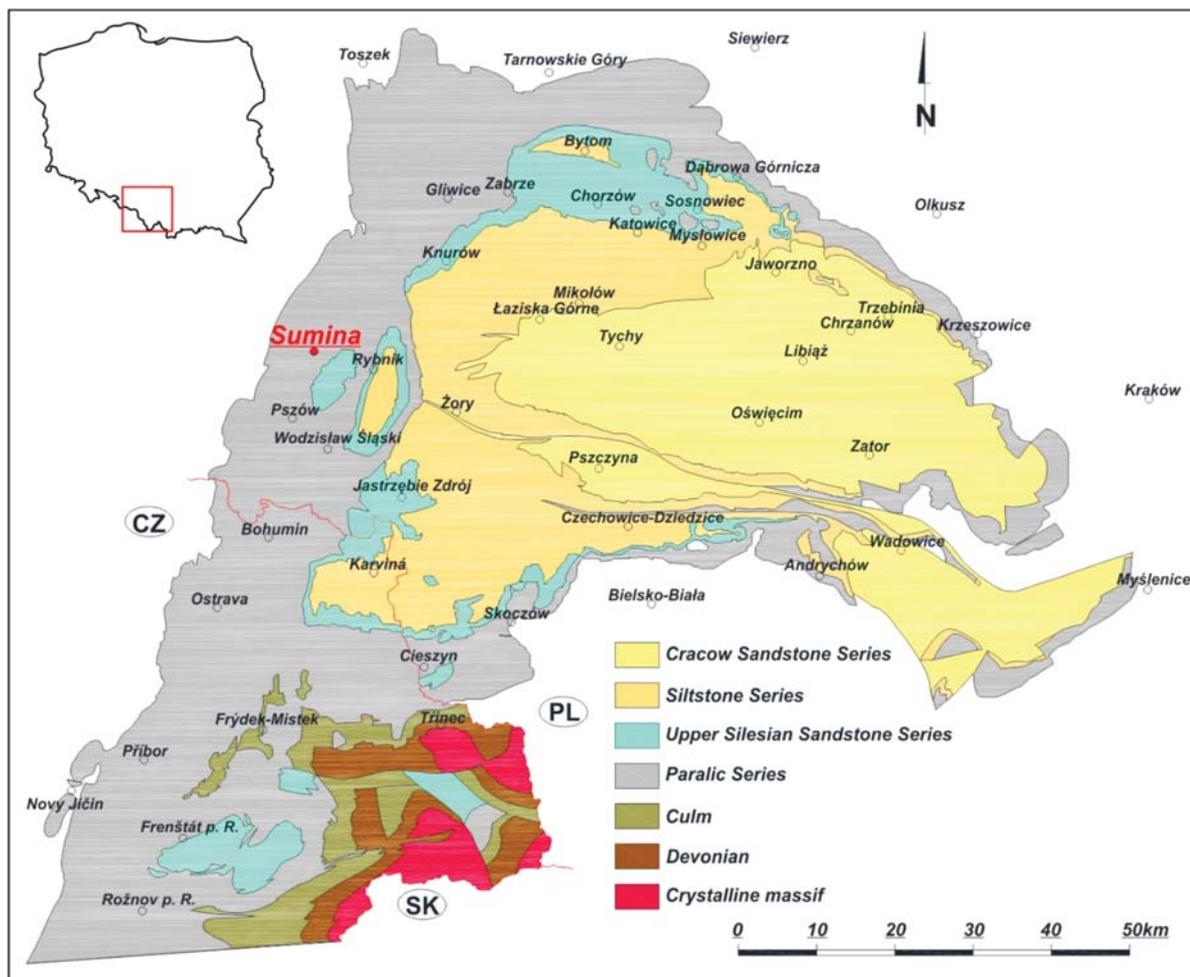
Vitrinite of uniaxial negative optical character is characteristic of tectonically undeformed coal deposits. The biaxial negative or positive optical character of vitrinite emerges, when in the process of coalification the coal is subjected, in addition to overburden pressure and rise of temperature, to a tectonic stress field at a non-perpendicular direction in relation to the bedding (Ting 1981; Levine and Davis 1984; Kilby 1986, 1988; Levine and Davis 1989a, b, 1990; Littke *et al.* 2012; Bruns and Littke 2015).

There is therefore a relationship established between the optical properties of coal (vitrinite) and the tectonic stress field of the deposit (Stone and Cook 1979). Results of the researches on the effects of tectonic processes on the optical characteristics of vitrinite in the bituminous coal and anthracites of the Upper Silesian Coal Basin (USCB) were presented, among others, in the works of Komorek, Morga and Pozzi (Komorek *et al.* 1995; Komorek 1996; Komorek and Pozzi 1996; Pozzi 1996; Morga 2000).

The temperature prevailing in the rock mass is reflected in the degree of coalification of the coaly matter or in the thermal maturity of the organic matter, and is usually expressed by the value of the vitrinite reflectance. It is accepted that the degree of coalification in USCB which occurred subsequent to the deposition of the Carboniferous sediments displays both synorogenic and postorogenic characters, while further coalification was not generated by burial- or tectonic related regional metamorphism, being described as a locally occurring phenomenon. It is usually associated with an additional source of heat resulting from intrusive bodies (Gabzdyl and Probiez 1987; Probiez 1989; Klika and Kraussova 1993). The type of coal alteration depends on the temperature of the igneous intrusion, the duration of the magmatic heating, the distance from the igneous rock, and the original coal rank prior to thermal alteration. The influence of thermal metamorphism on coal is diverse. The width of the contact zone is usually limited from a few centimeters up to several meters (Sarana and Kar 2011). The optical changes of the coal induced by the thermal metamorphism depend on the initial rank of the coal at the time of intrusion (Kwiecińska and Petersen 2004; Hartkopf-Fröder *et al.* 2015).

A common effect of thermal metamorphism of coal is its transformation in the intrusion area up to a meta-anthracite and/or natural coke. The role of pressure and temperature in the transformation processes is widely discussed in the literature (Kwiecińska *et al.* 1992, 1995; Yule *et al.* 2000; Gurba and Weber 2001; Stewart *et al.* 2005; Amijaya and Littke 2006; Cooper *et al.* 2007; Mastalerz *et al.* 2009; Borrego and Martin 2010; Morga 2010; Wang *et al.* 2010; Littke *et al.* 2012; Suarez-Ruiz *et al.* 2012; Valentim *et al.* 2013; Rahman and Rimmer 2014). It is generally believed that pressure counteracts temperature rise, and thus prevents an increase in coalification, which is illustrated by a non-linear increase in the degree of coalification of organic matter with increasing depth. The regional field of burial metamorphism is affected by anomalies associated with thermal metamorphism caused by numerous intrusions found in the USCB (Chodyncka and Sankiewicz 1978; Gabzdyl and Probiez 1987; Probiez 1989; Probiez and Lewandowska 2004; Matuszewska *et al.* 2015).

Scientific research on coal aimed at determining the rate of change of the optical properties of coal with increasing temperature has shown that after short heating periods at 350–400 °C, the reflectance and bireflectance of vitrinite has increased (Goodarzi and Murchison 1972; Murchison 1991; Jimenez *et al.* 1999).



Text-fig. 1. Location of the study area on the background of the geological structure of the Upper Silesian Coal Basin (Poland) (according to Kotas 1994, modified)

Changes in the optical properties of coals studied under laboratory conditions are associated with thermal changes that occur in both the internal structure of vitrinite and the products of its transformation (mesophases and matrix). The nature and intensity of these changes depend on the original degree of coalification and the coking capacity of the precursor material. It has been found that the highest optical anisotropy of vitrinite and matrix is observed in coke formed at a temperature of 1200 °C, after heat treatment of concentrates from typically graphitable substances (coking coal) (Komorek *et al.* 2000; Komorek and Morga 2001; Morga and Komorek 2002; Komorek and Morga 2003; Morga and Komorek 2004).

The mesophase and matrix show a stronger optical anisotropy than the vitrinite. In coke obtained from coking coal the reflectance and bireflectance of the

matrix attain values that are typical of graphite (Taylor *et al.* 1998; Komorek and Morga 2001, 2003; Morga and Komorek 2004; Komorek and Morga 2007).

During heating at 1200 °C vitrinite retains its optical character determined in the raw concentrate, despite the changes occurring at lower temperatures, particularly in the plasticization phase (Komorek *et al.* 2000, 2001; Komorek and Morga 2001; Morga and Komorek 2002; Komorek and Morga 2003; Komorek and Morga 2007; Morga and Komorek 2004).

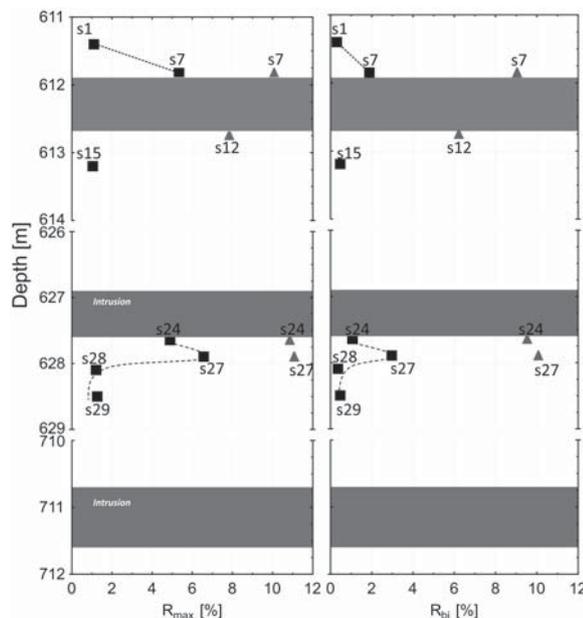
Research aimed at determining the effect of temperature on the optical properties of vitrinite indicates that the changes in the internal structure of vitrinite that take place with increasing heating temperature can be divided into two phases. The first phase occurs in the temperature range of 400 °C to 600 °C, when rapid changes are observed in the op-

tical parameters that characterize the internal structure of the heated vitrinite. Mesophase appears in vitrinite from coking coal at 500 °C or 600 °C. At higher temperature mesophase is transformed into matrix. These changes are associated with the process of degassing and plasticization. The second phase of changes is observed in the temperature range between 800 °C and 1200 °C. Within this range there is a further increase in mean reflectance R_r and bireflectance R_{bi} of vitrinite. The content as well as the value of mean reflectance R_r and bireflectance R_{bi} of the matrix also increase. This means that with an increasing heating temperature the degree of internal structure arrangement of the products of vitrinite carbonization also increases (Komorek and Morga 2001, 2003, 2007; Komorek 2013).

Under geological conditions, magmatic events provide an additional stimulus, leading to thermal transformations of coal. Numerous occurrences of volcanic rocks are present in the southern part of the USC B. Tuff and breccias of basic volcanic rocks as well as basalts have been found in the same part of the USC B (Kuhl 1954; Gabzdyl 1964; Chodyniecka and Sankiewicz 1972, 1978). Veins of basalt originating from a basalt dyke have been encountered during drilling in the Sumina area (well S-7). This well is situated on the northwest slope of the Jejkowice basin (Text-fig. 1).

The results of investigations of the basalt and sedimentary rocks identified in well S-7 were presented in the paper by L. Chodyniecki and J. Sankiewicz (1978). The S-7 well, which was drilled from the surface, penetrates Quaternary formations (claystones and sandstones), Neogene formations (sandy claystones with marl and gypsum interlayers), and Carboniferous formations represented by Gruszów (marginal) layers (Upper Mississippian). The Gruszów layers contain mostly claystones with minor interbedding of mudstones and sandstones. Basalt veins were found in three sections of the sediment core occurring within Carboniferous formations at depths of 611.95–612.8, 626.9–627.70, and 710.70–711.60 m. Their thicknesses were as follows: 0.85, 0.80 and 0.90 m, respectively (Text-fig. 2).

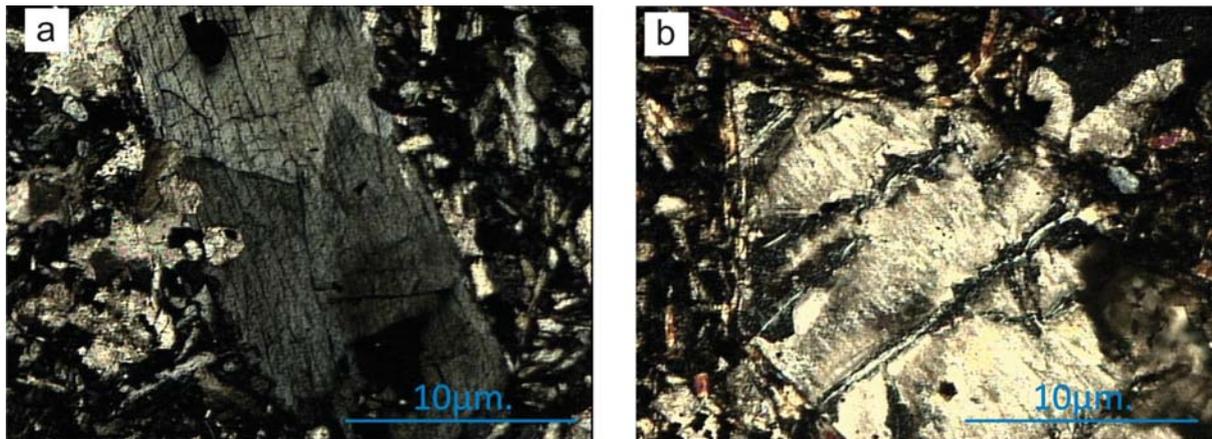
The basalt from the Sumina region belongs to the Central European Volcanic Province (CEVP). It has a holocrystalline porphyric texture and disordered structure. The petrographic composition of the basalts from the individual sections is similar. Phenocrysts include idiomorphic pyroxene (augite, Text-fig. 3a), and olivine pseudomorphs (Text-fig. 3b). Products of pyroxene transformations in the form of chlorites and iron oxides are observed sporadically



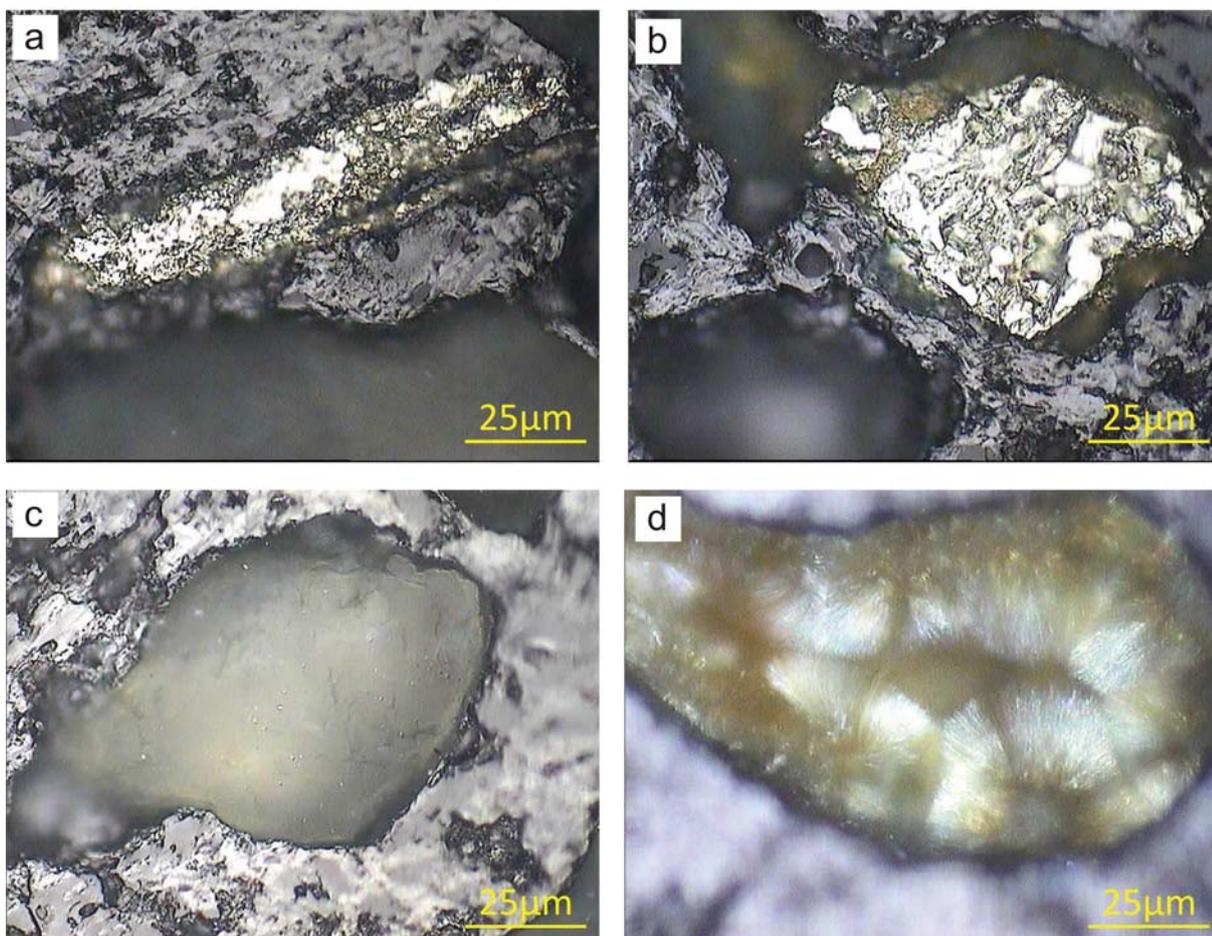
Text-fig. 2. Variation of the optical properties of coke matrix (▲) and vitrinite (■) in the profile of Sumina borehole S7

on the edges of pyroxene grains. Olivine phenocrysts are also highly transformed. Carbonates (calcite, rarely dolomite) and serpentines (chrysotile, rarely antigorite) are the major secondary minerals that fill olivine. The groundmass is composed of fine-grained augite, magnetite, nepheline, zeolites (anal-cime), and carbonates. Magnetite and nepheline form fine crystals evenly distributed in the groundmass. Zeolites usually fill vesicles and form a mixture of several minerals.

L. Chodyniecka and J. Sankiewicz observed an interface between basalt and sandstone at a depth of 626.90 m. The S-7 well log also revealed the presence of claystones. Basalt/claystone interfaces were found at the following depths: 611.95, 612.80, 627.9, 710.70 and 711.60 m. A brownish-yellow glassy coat has formed on the claystone/basalt endocontact. Coal interlayers in the claystones have been transformed to form natural coke. The presence of some constituents (e.g. natural coke) as well as structural and thermal transformations observed in the claystones (brownish-yellow rock coat) display a clear evidence of thermal metamorphism. This is the effect of the contact metamorphism of basalt on claystone. The presence of numerous vesicles filled with sulphides (Text-fig. 4a, b), carbonates (Text-fig. 4c), and zeolites (Text-fig. 4d) may also be an indication of postmagmatic hydrothermal activity on the solidified



Text-fig. 3. Microphotographs of basalt samples: a – phenocryst including idiomorphic pyroxene (augite); b – calcite-silica pseudomorphs after olivine; (magnification 100 \times)



Text-fig. 4. Vesicles filled with sulphides (a,b), carbonates (c), and zeolites (d); (magnification 500 \times)

magma and surrounding rocks (Gabzdyl and Probiez 1987; Probiez 1989; Klika and Kraussova 1993).

This paper is a continuation of a previous study (Chodyniewska and Sankiewicz 1978) and concerns

the optical properties of dispersed organic matter in claystones and sandstones occurring in the vicinity of basaltic intrusions (basaltic dykes). Samples from well S-7 containing dispersed organic matter occurring in the vicinity of basalt intrusions are unique in the USCB.

SAMPLING AND METHODS

Studies on the optical properties were carried out on dispersed organic matter examined in samples of claystones and sandstones collected near the basaltic intrusions. The cut samples were used to prepare polished grain mounts for an examination in reflected white light. The samples were prepared and the examination performed in according with PN-ISO 7404-2: 2005. These studies included reflectance measurements (according to PN-ISO 7404-5: 2002) of randomly oriented vitrinite grains and the products of thermal transformations of organic matter (i.e.: matrix of natural coke).

In incident polarized white light every anisotropic constituent shows, when performing a rotation of the microscopic stage through 360°, an apparent maximum R'_{\max} and minimum R'_{\min} reflectance. Random values of reflectance R'_{\max} and R'_{\min} were registered for each measurement point. About 100 reflectance measurements were made for each analyzed constituent (vitrinite, and matrix of natural coke) in the individual samples. Based on these results, using Kilby's method and a computer program (Kilby 1986, 1988, 1991), the principle axes of the reflectance indicating surface, RIS: a true maximum R_{\max} , intermediate R_{int} and minimum R_{\min} reflectance values were calculated. The value of optical anisotropy was expressed by means of bireflectance R_{bi} and by parameters introduced by W.E. Kilby (Kilby 1986, 1988, 1991): R_{am} (am – anisotropy magnitude) and R_{st} (st – style of reflectance indicating surface).

In isotropic bodies $R_{\text{am}} = 0$, whereas when $R_{\text{am}} > 0$ the value of that factor describes the deviation from the isotropic state. The higher the value the higher is the anisotropy, and at $R_{\text{am}} = 0.1$ anisotropy is described as very high (Kilby 1988, 1991).

The optical character of the studied constituents is defined by means of R_{st} (st – style). R_{st} may take on values between -30 and +30. When $R_{\text{st}} = -30$, the optical character is described as uniaxial negative,

whereas when $R_{\text{st}} = +30$, the optical character is uniaxial positive. The values between -30 and +30 allow us to classify coal constituents as biaxial bodies, and the minus and plus signs indicate a negative or positive optical character (Kilby 1988, 1991).

Microscopic examinations were carried out using a Zeiss optical light microscope equipped with a microphotometer. An immersion oil was used with a refractive index $n_o = 1.5176$ at 23 °C and light wavelength $\lambda = 546$ nm.

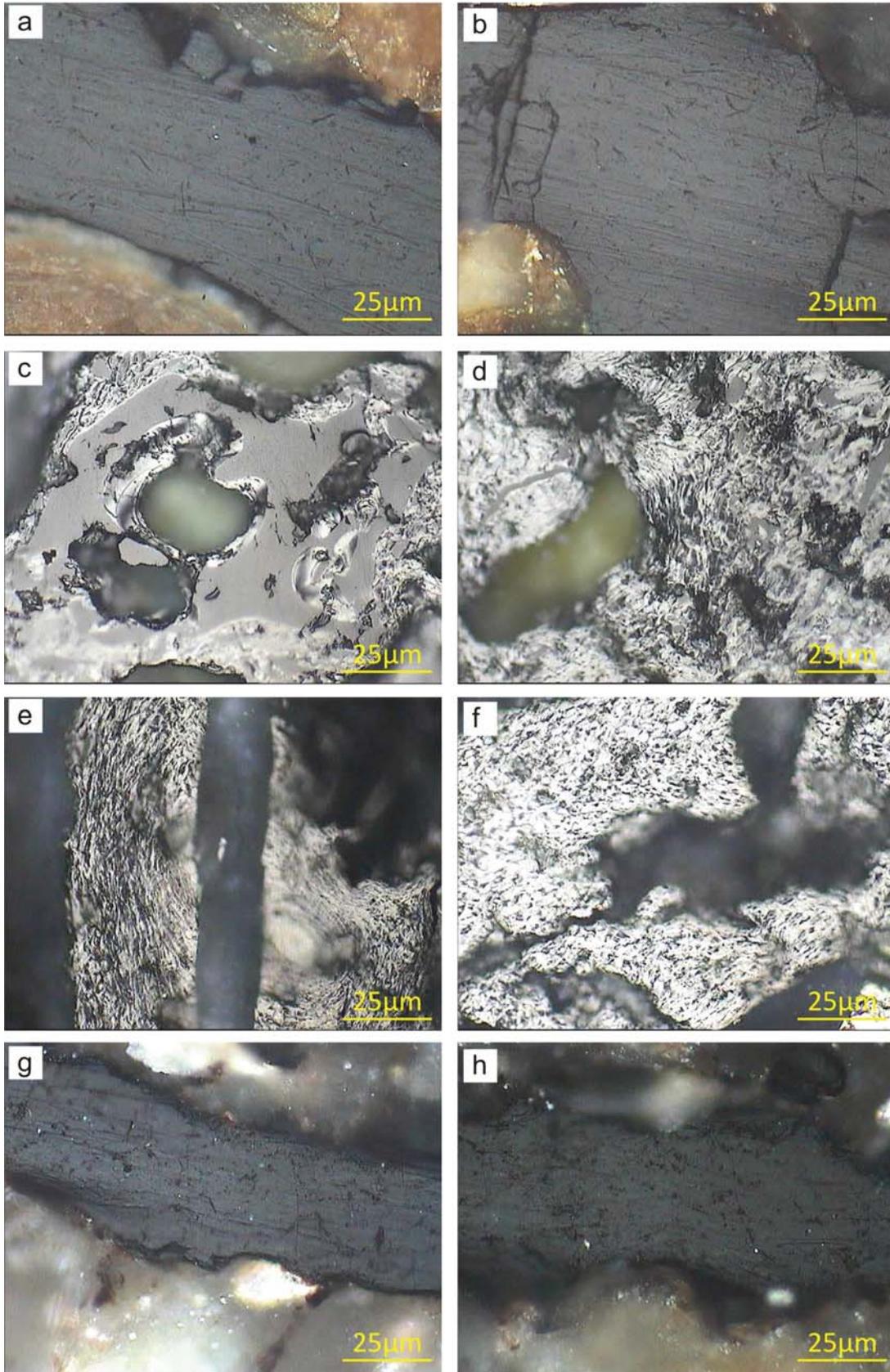
RESULTS AND DISCUSSIONS

At a distance of 25 cm from the upper margin of the uppermost basalt vein and the lower margin of the second vein the presence of porous natural coke was observed (Text-fig. 2). The natural coke is a product of the thermal transformation of organic matter dispersed within the examined sedimentary rocks that surround the basaltic intrusion. The matrix of the natural coke displays mainly a fine- and coarse-grained mosaic texture accompanied by a minor content of vitrinite grains (samples s7, s24, s27). One exception is sample s12 collected from the lower margin of the second vein and lacking vitrinite particles. In addition also some fusinite, semifusinite, and funginite macerals were observed in the natural coke matrix. The coke matrix showed optical anisotropy. Microscopic observations revealed the presence of areas optically uniformly oriented – anisotropic domains (Table 1).

Kilby's diagrams for the natural coke matrix in the examined samples are very clear. The true maximum reflectance values R_{\max} calculated from Kilby's diagrams varies from 7.88% to 11.13%, whereas that of intermediate reflectance R_{int} – from 5.00% to 6.97%, and that of minimum reflectance R_{\min} – from 1.01% to 1.63% (Table 2).

The bireflectance R_{bi} of the matrix varies from 6.25% to 10.12%. R_{am} varies between 0.22 and 0.27, which means that the internal structure of the coke matrix is highly ordered. The values of R_{st} fall within the interval of -2.58 to -7.99, meaning that the coke matrix in all the analyzed samples has a biaxial negative optical character (Table 2). The optical properties of the studied natural coke matrix are characteristic of semigraphite ($R_{\max} > 9\%$) (Taylor *et al.* 1998; Kwiecińska and Petersen 2004).

Text-fig. 5. Microphotographs of vitrinite and natural coke matrix from examined samples: sample S1 – vitrinite isolated particle and lamina (a, b); sample S7 – natural coke grains that constitute vitrinite residues (c), natural coke matrix – visible anisotropic domains (d); sample S12 – natural coke matrix – visible elongated anisotropic domains (e, f); sample S15 – vitrinite laminae (g, h), (magnification 500×) →



Sample	Distance from basalt intrusion [m]	Rock	Forms of organic matter
s1	0.65–0.70	fine-grained sandstone	Vitrinite – laminas, clasts, in addition to vitrinite (Text-fig. 5a, b), scarce grains of fusinite and semifusinite
s7	0.00–0.05	thermally altered claystone	Natural coke – laminas, lenses, clasts; matrix of mosaic texture, visible anisotropic domains (Text-fig. 5d), grains that constitute vitrinite residues (Text-fig. 5c), scarce clasts of fusinite and semifusinite
s12	0.00–0.05	claystone in direct contact with basalt	Natural coke – clasts and lenses; coke matrix of mosaic texture with visible elongated anisotropic domains, no vitrinite residues (Text-fig. 5e, f)
s15	0.45–0.75	claystone	Vitrinite – laminas and sharp-edged clasts (Text-fig. 5g, h), visible clasts of fusinite and semifusinite
s24	0.00–0.05	thermally altered claystone	Natural coke – laminas, porous coke matrix of mosaic texture with visible anisotropic domains; visible grains that constitute vitrinite residues (Text-fig. 6a, b) and clasts of fusinite and semifusinite
s27	0.25–0.45	claystone	Natural coke – laminas, lenses, clasts; matrix with elongated anisotropic domains visible on pore walls, grains that constitute vitrinite residues (Text-fig. 6c, d) and scarce clasts of fusinite, semifusinite and funginite
s28	0.45–0.85	claystone	Vitrinite – laminas, clasts (Text-fig. 6e, f)
s29	0.85–1.15	claystone	Vitrinite – fine clasts and lenses (Text-fig. 6g, h)

Table 1. Characteristics of the examined samples from the Sumina S7 borehole

Sample	R' _{max} [%]	s _{max} [%]	R' _{min} [%]	s _{min} [%]	R _{max} [%]	R _{int} [%]	R _{min} [%]	R _{bi} [%]	R _{am}	R _{st}
Vitrinite										
s1	0.97	0.06	0.87	0.05	1.09	0.89	0.75	0.34	0.06	5.89
s7	4.56	0.39	4.21	0.36	5.33	4.28	3.44	1.89	0.07	3.73
s15	0.86	0.11	0.77	0.11	1.07	0.82	0.59	0.48	0.10	1.42
s24	4.54	0.22	4.14	0.24	4.93	4.37	3.82	1.11	0.04	0.39
s27	5.24	0.68	4.51	0.47	6.24	4.68	3.80	2.44	0.08	9.20
s28	1.05	0.08	0.96	0.07	1.22	0.98	0.82	0.40	0.07	6.65
s29	1.10	0.09	1.04	0.13	1.26	1.05	0.78	0.48	0.08	-4.08
Matrix										
s7	8.70	0.85	3.59	1.38	10.12	6.70	1.08	9.04	0.25	-7.99
s12	6.03	0.80	3.37	0.87	7.88	5.00	1.63	6.25	0.22	-2.58
s24	8.86	1.02	4.05	1.54	10.88	6.97	1.31	9.57	0.25	-6.02
s27	8.72	1.23	4.05	1.70	11.13	6.83	1.01	10.12	0.27	-4.95

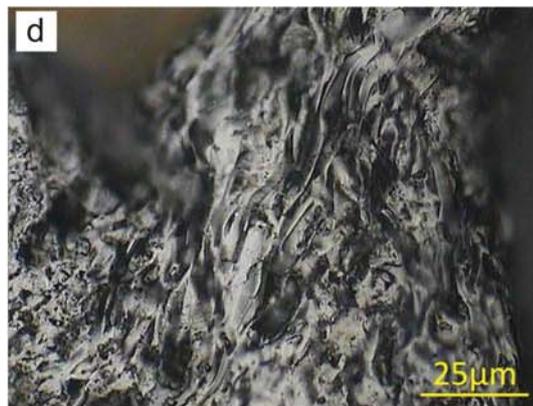
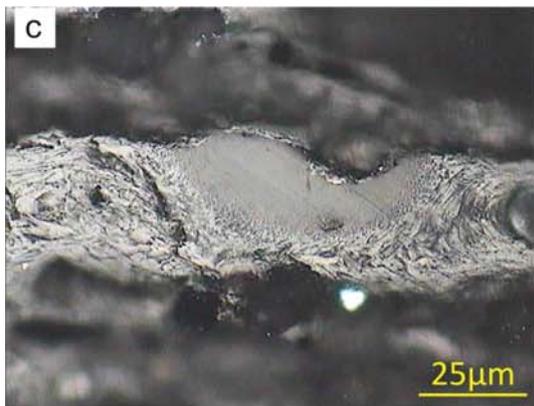
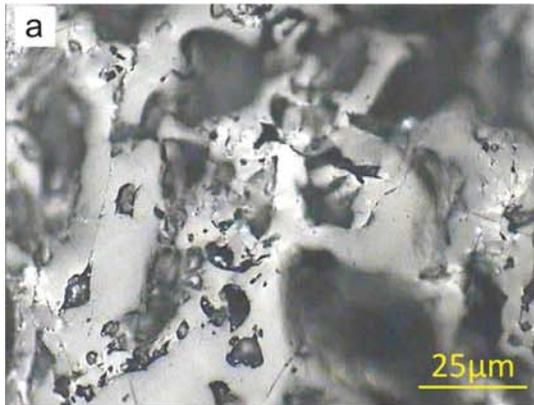
Table 2. Optical properties of vitrinite and natural coke matrix: R' _{max} – mean apparent maximum reflectance, s _{max} – standard deviation of mean apparent maximum reflectance value, R' _{min} – mean apparent minimum reflectance value, s _{min} – standard deviation of mean apparent minimum reflectance value, R _{max} – true maximum reflectance, R _{int} – true intermediate reflectance, R _{min} – true minimum reflectance, R _{bi} – bireflectance, R _{am} – anisotropy magnitude, R _{st} – reflectance indicating surface – style)

In the vicinity of the analyzed area it has been found that the vitrinite from coal seams (not subjected to thermal alteration) has a reflectance of from 0.67% to 1.08% (Adamczyk and Komorek 1999). Thermally altered vitrinite, which is present in the natural coke matrix in minor amounts, has the form of small particles with rounded edges. The true maximum reflectance

value of R _{max} calculated from Kilby's diagrams varies from 4.93% to 6.24%, whereas that of intermediate reflectance R _{int} – from 4.28% to 4.68%, and that of minimum reflectance R _{min} – from 3.44% to 3.82% (Table 2).

The bireflectance R _{bi} of thermally altered vitrinite varies from 1.11% to 2.44%. R _{am} assumes values

Text-fig. 6. Microphotographs of vitrinite and natural coke matrix from examined samples: sample S24 – natural coke grains that constitute vitrinite residues (a), natural coke matrix – visible anisotropic domains (b); sample S27 – natural coke grains that constitute vitrinite residues (c), natural coke matrix – visible anisotropic domains (d); sample S28 – vitrinite clast and lamina (e, f); sample S29 – vitrinite lamina and isolated particle (g, h), (magnification 500×)



in the range of 0.04 to 0.08. The results show that thermally affected vitrinite grains in coke samples are characterized by lower values of true reflectance and by lower optical anisotropy than that of the natural coke matrix. The values of R_{st} fall within the interval of 0.39 to 9.20, meaning that the examined vitrinite in all analyzed samples has a biaxial positive optical character (Table 2).

Thermally altered dispersed organic matter was also observed in samples taken at a distance greater than 45 cm from the respective basalt veins (samples s1, s15, s28, s29) and it is characterised by a lower thermal degree of alteration. Organic matter in these samples was represented mainly by vitrinite occurring in the form of laminae, lenses, and as single grains. Apart from vitrinite, scarce macerals of semifusinite and fusinite were encountered (Table 1).

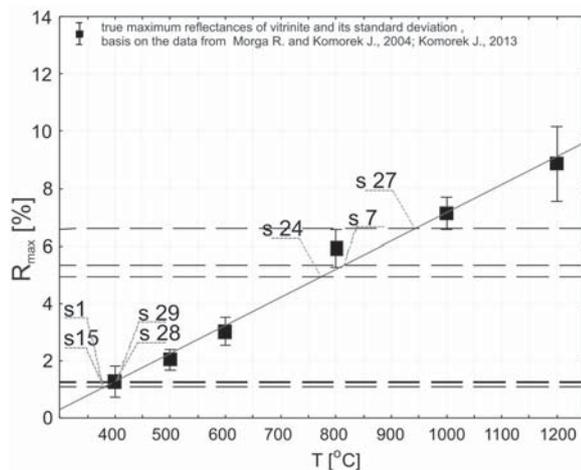
It should be noted that Kilby's diagrams of the thermally altered vitrinite in these samples are less ordered and more challenging to interpret than those of the severely thermally altered vitrinite in a natural coke matrix discussed above. This may be due to disordering of the vitrinite structure caused by an increased temperature (Komorek and Morga 2001, 2003; Morga and Komorek 2004; Komorek and Morga 2007; Komorek 2013). Therefore, the values of reflectance, bireflectance, and R_{am} and R_{st} calculated for vitrinite in these samples should be treated as approximate. The true maximum reflectance R_{max} of vitrinite varies from 1.07% to 1.26%, whereas the true intermediate reflectance R_{int} – from 0.82% to 1.05%, and the true minimum reflectance R_{min} – from 0.59% to 0.82%. Bireflectance R_{bi} attains values from 0.34% to 0.48%. R_{am} varies between 0.06 to 0.10. The values of R_{st} in nearly all of the analysed samples are positive, meaning that vitrinite displays a biaxial positive optical character (Table 2). Only vitrinite in sample s29 has a biaxial negative optical character ($R_{st} = -4.08$). Note that most of the vitrinites in coals from USCB are characterized by negative, biaxial or uniaxial optical characters (Komorek 1996; Pozzi 1996; Morga 2000; Adamczyk *et al.* 2014). Differentiation in the optical character of the analyzed samples of these thermally altered vitrinites influenced by a lower degree of thermal alteration may be evidence of changes in their internal structure caused by heat sourced from the respective basaltic intrusions. A confirmation of the effect of temperature on the microstructure of vitrinite in samples taken at a larger distance from the basalt veins may be represented by the pores observed in some vitrinite grains, these pores being probably formed by degassing (s28, s29). The differing optical character of these vitrinites may

also result from the fact that the dispersed organic matter observed in the sedimentary rocks and occurring in the form of clasts may be resedimented.

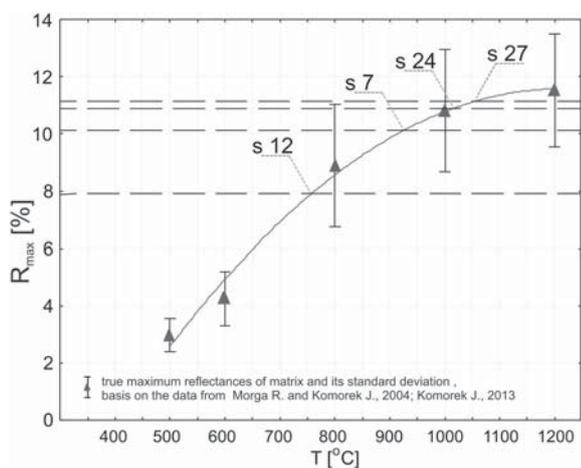
The analysis of the variation of the true vitrinite maximum reflectance value with changing distance from the basaltic intrusion indicates in general that higher values are observed in samples located closer to the basaltic veins. Noteworthy are samples s24, s27, s28 and s29 taken at the lower margin of the second vein (at depths of more than 627m). Here we observe a reversal of reflectance changes with distance from the intrusion similar to that previously described in the literature (Murchison 2006). This can be explained by changes in the arrangement of the internal structure of vitrinite resulting from the influence of heat (Khorasani *et al.* 1990; Murchison 2006). The bireflectance R_{bi} shows a similar reversal of the trend in changes depending on the distance from the second basalt vein (Text-fig. 2).

The calculated values of true maximum reflectance of vitrinite and matrix of natural coke were used to estimate the maximum temperature that acted on the organic matter dispersed in the rocks surrounding the intrusion.

The values of the true maximum vitrinite reflectance were compared with the data obtained for vitrinite subjected to thermal treatment in the laboratory experiments within the range of 400 °C to 1200 °C. The true maximum reflectance values of the natural coke matrix were compared with the data for cokes obtained in the laboratory from coals of different ranks (Morga and Komorek 2004; Komorek 2013). The arithmetic mean and standard deviation of the true maximum reflectance of vitrinite and matrix were calculated for every temperature of heat treatment (400, 500, 600, 800, 1000 and 1200 °C). The relationship was then determined between temperature and the mean value of the true maximum vitrinite and matrix reflectance $R_{max} = f(T)$. In the case of vitrinite that relationship was linear $R_{max} = 0.0099T - 2.7147$ ($r = 0.99$, $p < 0.05$) (Text-fig. 7). In the case of the matrix the relationship $R_{max} = f(T)$ was described with a second-degree polynomial $R_{max} = -1.7748 \cdot 10^{-5}T^2 + 0.0430T - 14.5013$, $p < 0.05$ (Text-fig. 8). These functions were used to obtain general estimates of the paleotemperature that acted on the rocks surrounding the basaltic intrusion. The obtained temperature data are only approximations used to correlate the determined reflectance data with the temperatures derived from laboratory conditions (Text-figs 7, 8). These functions were used to determine the paleotemperature that acted on the rocks surrounding the intrusion.



Text-fig. 7. Comparison of maximum vitrinite reflectance with the data obtained for cokes produced in laboratory conditions (Morgia and Komorek 2004; Komorek 2013)



Text-fig. 8. Comparison of maximum coke matrix reflectance with the data obtained for cokes produced in laboratory conditions (Morgia and Komorek 2004; Komorek 2013)

It can be suggested that in the immediate vicinity of the basalt veins (samples: s7, s12, s24, s27) the estimated temperature was higher than 750 °C (Text-fig. 7, 8). Samples taken from locations more distant from the intrusion (samples: s1, s15, s28 and s29) were found to be very likely subjected to temperatures of around 400 °C. It should be noted that no matrix was found in the samples situated more distant from the intrusion, which may indicate that the temperature acting on these samples was possibly lower than 500 °C. Under laboratory conditions, matrix appears only at temperatures exceeding 500 °C (Komorek and Morgia 2001, 2003; Morgia and Komorek 2004; Komorek and Morgia 2007; Komorek 2013).

CONCLUSIONS

In the Sumina region three Neogene basaltic veins were encountered in the S-7 well, intruding into Carboniferous formations at the depths from 611.95 m to 711.60 m. At a distance of 0.25 m from two of the veins the presence of porous natural coke was observed. This coke consists mainly of fine- and coarse-grained matrix of mosaic texture and is associated with a small content of vitrinite grains as well as fusinite, semifusinite, and funginite macerals.

Dispersed organic matter, represented mainly by vitrinite in the form of laminas, lenses, and isolated particles, was also found in samples taken at a distance greater than 0.45 m from the basalt veins.

The optical properties of the studied natural coke samples are characteristic of semi-graphite ($R_{\max} > 9\%$). The examined matrix of the natural coke for all of the analysed samples has a biaxial negative optical character. Vitrinite in the natural coke samples is characterized by lower optical anisotropy than that of the matrix and it has a biaxial positive optical character.

Vitrinite in almost all samples taken at locations more distant from the intrusion has a biaxial positive optical character, with the exception of vitrinite in sample s29, having biaxial negative optical character.

A reversal of the changes of true maximum vitrinite reflectance and bireflectance with changing distance from the second basalt vein, previously described in the literature, has been observed. This is suggested to be related to changes in the arrangement of the internal structure of vitrinite, resulting from the action of heat.

The temperature that had acted on the organic matter in the immediate vicinity of the intrusion, estimated on the basis of experimental data, was higher than 750 °C, while the samples taken from locations more distant from the intrusion (samples: s1, s15, s28 and s29) were found to have been subjected to temperatures of around 400 °C.

Acknowledgments

The authors would like to thank Professor Lidia Chodyniecka for providing rock samples from borehole Sumina S-7 for researches. The authors are grateful to reviewers: Prof. B. Kwiecińska and Dr Jolanta Kus for their valuable comments on the manuscript. Special thanks are to the journal editor Piotr Luczyński for his editorial work.

REFERENCES

- Adamczyk, Z. and Komorek, J. 1999. Reflectance of vitrinite from barren rock inserts of coal seams of the upper marginal beds in Jejkowice trough (USCB). *Zeszyty Naukowe Politechniki Śląskiej, Górnictwo*, **241**, 21–35. [In Polish]
- Adamczyk, Z., Komorek, J. and Lewandowska, M. 2014. Specific types of coal macerals from Orzesze and Ruda beds from “Pniówek” coal mine (USCB – Poland) as a manifestation of thermal metamorphism. *Archives of Mining Sciences*, **59**, 1, 77–91.
- Amijaya, H. and Littke, R. 2006. Properties of thermally metamorphosed coal from Tanjung Enim area South Sumatra Basin, Indonesia with special reference to the coalification path of maceral. *International Journal of Coal Geology*, **66**, 271–295.
- Borrego, A.G. and Martín, A.J. 2010. Variation in the structure of anthracite at a fast heating rate as determined by its optical properties: An example of oxy-combustion conditions in a drop tube reactor. *International Journal of Coal Geology*, **81**, 301–308.
- Bruns, B. and Littke, R. 2015. Lithological dependency and anisotropy of vitrinite reflectance in high rank sedimentary rocks of the Ibbenbüren area, NW-Germany: Implications for the tectonic and thermal evolution of the Lower Saxony Basin. *International Journal of Coal Geology*, **137**, 124–135.
- Chodynicka, L. and Sankiewicz, J. 1972. Magmatic intrusion in Lower Namurian in the Marklowice region, Upper Silesia Coal Basin Poland. *Rocznik Polskiego Towarzystwa Geologicznego*, **42**, 309–326. [In Polish]
- Chodynicka, L. and Sankiewicz, J. 1978. Basalt from the Sumina area (Rybnik Coal Basin). *Geological Quarterly*, **22**, 119–133. [In Polish]
- Cook, A.C., Murchison, D.G. and Scott E. 1972. Optically biaxial anthracitic vitrinite. *Fuel*, **51**, 180–184.
- Cooper, J.R., Crelling, J.C., Rimmer, S.M. and Whittington, A.G. 2007. Coal metamorphism by igneous intrusion in the Raton Basin, CO and NM: implication for generation of volatile. *International Journal of Coal Geology*, **71**, 15–27.
- Davis, A. 1978. The Reflectance of Coal. In: Karr, C. (Ed.), *Analytical Methods for Coal and Coal Products 1*, pp. 27–81. Academic Press; New York.
- Ergun, S. and McCartney, J.T. 1960. Reflectance of coals, graphite and diamond. *Fuel*, **39**, 449–454.
- Fink, R., Virgo, S., Arndt, M., Visser, W., Littke, R. and Urai, J.L. 2016. Solid bitumen in calcite veins from the Natih Formation in the Oman Mountains: Multiple phases of petroleum migration in a changing stress field. *International Journal of Coal Geology*, **157**, 39–51
- Gabzdyl, W. 1964. Przejawy metamorfizmu kontaktowego na kopalni Jastrzębie. *Zeszyty Naukowe Politechniki Śląskiej*, **12**, 107–121, Gliwice. [In Polish]
- Gabzdyl, W. and Proberz, K. 1987. The occurrence of anthracites in an area characterized by lower rank coals in the Upper Silesian Coal Basin of Poland. *International Journal of Coal Geology*, **7**, 209–225.
- Goodarzi, F. and Murchison, D. G. 1972. Optical properties of carbonized vitrinite. *Fuel*, **51**, 322–328.
- Grieve, D.A. 1991. Biaxial vitrinite reflectance in coals of the Elk Valley coalfield, Southeastern British Columbia, Canada. *International Journal of Coal Geology*, **19**, 185–200.
- Gurba, L.W. and Weber, C.R. 2001. Effects of igneous intrusions on coal bed methane potential, Gunnedah Basin, Australia. *International Journal of Coal Geology*, **46**, 113–133.
- Hartkopf-Fröder, C., Königshof, P., Littke, R. and Schwarzbauer, J. 2015. Optical thermal maturity parameters and organic geochemical alteration at low grade diagenesis to anchimetamorphism: A review. *International Journal of Coal Geology*, **150–151**, 74–119.
- Hevia, V. and Virgos, J.M. 1977. The rank and anisotropy of anthracites: the indicating surface of reflectivity in uniaxial and biaxial substance. *Journal of Microscopy*, **109**, 23–28.
- Houseknecht, D.W. and Weesner, C.M.B. 1997. Rotational reflectance of dispersed vitrinite from the Arkoma basin. *Organic Geochemistry*, **26**, 191–206.
- Hower, J.C. and Davis, A. 1981a. Application of vitrinite reflectance anisotropy in the evaluation of coal metamorphism. *Geological Society America Bulletin*, **92**, part I, 350–366.
- Hower, J.C. and Davis, A. 1981b. Vitrinite reflectance anisotropy as a tectonic fabric element. *International Journal of Coal Geology*, **9**, 165–168.
- Jimenez, A., Iglesias, J.M., Laggoun-Defarge, F. and Suarez-Ruiz, I. 1999. Effect of the increase in temperature on the evolution of the physical and chemical structure of vitrinite. *Journal of Analytical and Applied Pyrolysis*, **50**, 117–148.
- Khorasani, K. G., Murchison, D.G. and Raymond, A.G. 1990. Molecular disordering in natural cokes approaching dyke and sill contact. *Fuel*, **69**, 1037–1046.
- Kilby, W.E. 1986. Biaxial reflecting coals in the Peace River coalfield. British Columbia Ministry of Energy, Mines and Petroleum, Geological Fieldwork, paper 1986-1, 127–137.
- Kilby, W.E. 1988. Recognition of vitrinite with non-uniaxial negative reflectance characteristic. *International Journal of Coal Geology*, **9**, 267–285.
- Kilby, W.E. 1991. Vitrinite reflectance measurement – some technique enhancements and relationship. *International Journal of Coal Geology*, **9**, 201–218.
- Klika, Z. and Kraussova, J. 1993. Properties of altered coals associated with Carboniferous red beds in the Upper Silesian Coal Basin and their tentative classification. *International Journal of Coal Geology*, **22**, 217–235.
- Komorek, J. 1996. Optical properties of coals of types 31–42 from the seams of the Upper Silesian Coal Basin. *Prace Geologiczne*, **140**, Monograph. [In Polish with English abstract]

- Komorek, J. 2013. Changes of internal structure of vitrinite and liptinite subjected thermal treatment within the range 400–1200 °C. Monograph. Wydawnictwo Politechniki Śląskiej; Gliwice. [In Polish with English abstract]
- Komorek, J. and Morga, R. 2001. Characteristics of optical properties of vitrinite subjected to thermal treating exemplified by coal from seam 833/1 the “Gliwice” mine. *Zeszyty Naukowe Politechniki Śląskiej, Górnictwo*, **249**, 59–73. [In Polish]
- Komorek, J. and Morga, R. 2003. Vitrinite reflectance property change during heating under inert condition. *International Journal of Coal Geology*, **54**, 125–136.
- Komorek, J. and Morga, R. 2007. Evolution of optical properties of vitrinite, sporinite and semifusinite in response to heating under inert condition. *International Journal of Coal Geology*, **71**, 389–404.
- Komorek, J., Morga, R. and Krzeszowska, E. 2001. Relationship between optical properties of sporinite and vitrinite subjected to the thermal treatment in laboratory condition. *Acta Universitatis Carolinae – Geologica*, **45**, 65–74.
- Komorek, J., Morga, R. and Lewandowski, M. 2000. Optical properties of heat treated vitrinite-coal from Chwałowice Colliery as an example. *Przegląd Górniczy*, **11**, 36–40. [In Polish with English abstract]
- Komorek, J., Morga, R. and Pozzi, M. 1995. Optical anisotropy of vitrinite in coal seams from the fold area in the Upper Silesian Coal Basin (Poland). Abstracts of the XIII International Congress on Carboniferous–Permian. Kraków, 78.
- Komorek, J. and Pozzi, M. 1996. Optical anisotropy of coal from the Jastrzębie Fold (Upper Silesian Coal Basin). *Geological Quarterly*, **40**, 393–405.
- Kotas, A. (Ed.) 1994. Coal-bed methane potential of Upper Silesian coal Basin, Poland. *Prace PIG, CXLII*, Warszawa.
- Kuhl, J. 1954. Tufogenic rocks in the Carboniferous of Upper Silesia. *Rocznik Polskiego Towarzystwa Geologicznego*, **22**, 181–208, Kraków. [In Polish]
- Kwiecińska, B. and Petersen, H.I. 2004. Graphite, semi-graphite, natural coke, and natural char classification – ICCP system. *International Journal of Coal Geology*, **57**, 99–116.
- Kwiecińska, B., Hamburg, G. and Vleeskens, J.M. 1992. Formation temperatures of natural coke in the Lower Silesian Coal Basin, Poland. Evidence from pyrite and clays by SEM-EDX. *International Journal of Coal Geology*, **21**, 217–235.
- Kwiecińska, B., Muszyński, M., Vleeskens, J. and Hamburg, G. 1995. Natural coke from the La Rasa mine, Tineo, Spain. *Mineralogia Polonica*, **26**, 2, 3–14.
- Langerberg, W. and Kalkreuth, W. 1991. Reflectance anisotropy and syn-deformational coalification of the Jewel seam in the Cadomin area, Alberta, Canada. *International Journal of Coal Geology*, **19**, 303–317.
- Langerberg, W. and Kalkreuth, W. 1991a. Tectonic control on regional coalification and vitrinite reflectance anisotropy of Lower Cretaceous coals in the Alberta Foothills, Canada. *Bulletin de la Société géologique de France*, **162**, 2, 375–383.
- Levine, J.R. and Davis, A. 1984. Optical anisotropy of coals as an indicator of tectonic deformation, Broad Top Coal Field, Pennsylvania. *Geological Society America Bulletin*, **95**, 100–108.
- Levine, J.R. and Davis, A. 1989a. Reflectance anisotropy of Upper Carboniferous coals in the Appalachian foreland basin, Pennsylvania, USA. *International Journal of Coal Geology*, **13**, 341–373.
- Levine, J.R. and Davis, A. 1989b. The relationship of coal optical fabrics to Alleghanian tectonic deformation in the central Appalachian fold and thrust belt, Pennsylvania. *Geological Society America Bulletin*, **101**, 1333–1347.
- Levine, J.R. and Davis, A. 1990. Reflectance anisotropy of Carboniferous coals in the Appalachian Foreland Basin, Pennsylvania, USA. *International Journal of Coal Geology*, **16**, 201–204.
- Littke, R., Urai, J.L., Uffmann, A.K. and Risvanis, F. 2012. Reflectance of dispersed vitrinite in Palaeozoic rocks with and without cleavage: Implications for burial and thermal history modeling in the Devonian of Rursee area, northern Rhenish Massif, Germany. *International Journal of Coal Geology*, **89**, 41–50.
- Mastalerz, M., Drobnik, A. and Schimmelmann, A. 2009. Changes in optical properties, chemistry, and micropore and mesopore characteristics of bituminous coal at the contact with dikes in the Illinois Basin. *International Journal of Coal Geology*, **77**, 310–319.
- Matuszewska, A., Pusz, S. and Duber, S. 2015. Evaluation of the structure of bituminous coal from Sośnica mine in the Upper Silesian Coal Basin (Poland) using reflectance indicating surface (RIS) parameters. *International Journal of Coal Geology*, **152**, Part B, 177–188.
- Middleton, M.F. 1991. Tectonic influence on vitrinite reflectance. *International Journal of Coal Geology*, **16**, 235–237.
- Morga, R. 2000. Optical anisotropy of coal in tectonically deformed seams in the Upper Silesian Coal Basin. *Prace Geologiczne*, **148**, Monograph. [In Polish with English abstract]
- Morga, R. 2010. Chemical structure of semifusinite and fusinite of steam and coking coal from the Upper Silesian Coal Basin (Poland) and its changes during heating as inferred from micro-FTIR analysis. *International Journal of Coal Geology*, **84**, 1–15.
- Morga, R. and Komorek, J. 2002. Influence of thermal treatment on optical properties and internal structure of anthracite. *Zeszyty Naukowe Politechniki Śląskiej, Górnictwo*, **254**, 111–128. [In Polish]
- Morga, R. and Komorek, J. 2004. Changes of optical properties and internal structure of vitrinite subjected to thermal treatment within the range of 400–1200 °C. *Prace Geologiczne*, **152**, Monograph. [In Polish with English abstract]

- Murchison, D.G. 1991. Petrographic aspect of coal structure: reactivity of macerals in laboratory and natural environment. *Fuel*, **70**, 296–315.
- Murchison, D.G. 2006. The influence of heating rates on organic matter in laboratory and natural environment. *International Journal of Coal Geology*, **67**, 145–157.
- PN-ISO 7404-2: 2005. Węgiel kamienny – Analiza petrograficzna – Przygotowanie próbek węgla do badań mikroskopowych.
- PN-ISO 7404-5: 2002. Metody analizy petrograficznej węgla kamiennego (bitumicznego) i antracytu – Część 5: Metoda mikroskopowa oznaczania refleksyjności wityrynytu.
- Pozzi, M. 1996. Optical anisotropy of coal from the seams of Jastrzębie area as a manifestation of tectonic stress. *Zeszyty Naukowe Politechniki Śląskiej, Górnictwo*, **229**, 98 p. [In Polish with English and French abstracts]
- Proberz, K. 1989. Effect of thermal metamorphism on coalification degree (rank) and petrographic composition of the coal seams in Jastrzębie region (Upper Silesia Coal Basin of Poland). *Zeszyty Naukowe Politechniki Śląskiej, Górnictwo*, **176**, Gliwice. [In Polish with English abstract]
- Proberz, K. and Lewandowska M. 2004. Paleotemperatures of Upper Carboniferous sedimentary rocks in the NW part of the Upper Silesian Coal Basin, Poland. *Geologica Belgica*, **7**, 3–4, 313–318.
- Rahman, M.W. and Rimmer, S.M. 2014. Effects of rapid thermal alteration on coal: Geochemical and petrographic signatures in the Springfield (No. 5) Coal, Illinois Basin. *International Journal of Coal Geology*, **131**, 214–226.
- Reinhardt, M. 1991. Vitrinite reflectance, illite crystallinity and tectonics: results from the Northern Apennines (Italy). *Organic Geochemistry*, **17**, 175–184.
- Sarana, S. and Kar, R. 2011. Effect of igneous intrusive on coal microconstituents: Study from an Indian Gondwana coalfield. *International Journal of Coal Geology*, **85**, 161–167.
- Stach, E., Mackowsky, M.-Th., Teichmüller, M., Taylor, G.H., Chandra, D. and Teichmüller, R. 1982. Stach's Textbook of Coal Petrology, 535 p. Gebruder Borntraeger; Berlin-Stuttgart.
- Stewart, A.K., Massey, M., Padgett, P.L., Rimmer, S.M. and Hower, J.C. 2005. Influence of a basic intrusion on the vitrinite reflectance and chemistry of the Springfield (No. 5) coal, Harrisburg, Illinois. *International Journal of Coal Geology*, **63**, 58–67.
- Stone, I.J. and Cook, A.C. 1979. The influence of some tectonic structures upon vitrinite reflectance. *International Journal of Coal Geology*, **87**, 497–508.
- Suárez-Ruiz, I., Flores, D., Mendonça Filho, J.G. and Hackley, P.C. 2012. Review and update of the applications of organic petrology: Part 1, geological application. *International Journal of Coal Geology*, **99**, 54–112.
- Taylor, G.H., Teichmüller, M., Davis, A., Diessel, C.K.F., Litke, R. and Robert, T. 1998. Organic Petrology, xvi + 704 p. Gebruder Borntraeger; Berlin-Stuttgart.
- Ting, F.T.C. 1981. Uniaxial and biaxial vitrinite reflectance models and their relationship to paleotectonics. In: Brook, J. (Ed.), Organic Maturation Studies and Fossil Fuel Exploration, pp. 379–392. Academic Press; London.
- Tsai, L.L.Y. 1991. A study of the reflectance indicatrix of vitrinite. International Conference on Coal Science Proceedings, pp. 127–130. Oxford.
- Valentim, B., Rodrigues, S., Ribeiro, S., Pereira, G., Guedes, A. and Suárez-Ruiz, I. 2013. Relationships between the optical properties of coal macerals and the chars resulting from fluidized bed pyrolysis. *International Journal of Coal Geology*, **111**, 80–89.
- Van Krevelen, D.W. 1993. Coal. 3rd edition. Elsevier; Amsterdam.
- Vries, H. A., Habets, P. J. and Bokhoven, C. 1968. Das Reflexionsvermögen von Steinkohle II. *Die Reflexionsanisotropie. Brennstoff-Chemie*, **49**, 47–52.
- Wang, J., Du, J., Chang, L. and Xie, K. 2010. Study on the structure and pyrolysis characteristics of Chinese western coal. *Fuel Processing Technology*, **91**, 430–433.
- Yule, B.L., Roberts, S. and Marshall, J.E. A. 2000. The thermal evolution of sporopollenin. *Organic Geochemistry*, **31**, 859–870.

Manuscript submitted: 6th November 2017

Revised version accepted: 20th February 2018