

Zbigniew KLIMOWICZ and Stanisław UZIAK

Department of Soil Science  
Institute of Earth Sciences  
Maria Curie-Skłodowska University  
Akademicka 19  
20-033 Lublin, POLAND

## Soil-forming processes and soil properties in Calypsostranda, Spitsbergen

**ABSTRACT:** Results of field investigations in 1986 in Calypsostranda, southern coast of Bellsund, are presented. Soils of dry, wet and very wet tundra were studied. Strong skeleton, presence of carbonates, neutral or alkaline reaction, low content of available phosphorus and potassium, high content of organic carbon (in mineral soils) are the most characteristic properties of investigated soils. Regularities in vertical distribution of some components were distinguished.

**Key words:** Arctic, Spitsbergen, soils.

### Introduction

Little has been written on soils of the arctic areas (Jahn 1946, Rieger 1974, Stäblein 1977, Andreiev 1980, Låg 1981, Linell and Tedrow 1981), Spitsbergen included (Szerszeń 1965, 1968, 1974; Boratyński *et al.* 1968; Baranowski and Szerszeń 1968; Plichta 1977, Szerszeń and Chodak 1983; Låg 1980, 1983; Klimowicz and Uziak 1987). Detailed and almost complete description of some soil types from Hornsund region was presented by Szerszeń (1965, 1968, 1974). Soils from the southern coast of Bellsund have not been fully described yet. In this paper results of investigations carried out in 1986 during the Scientific Expedition of the Institute of Earth Sciences, Maria Curie-Skłodowska University of Lublin, are presented.

### Area and methods

The studied area forms a part of coastal plain of southern Bellsund shore. It is constituted of several raised marine beaches that slope northeastwards.

Geologic structure of this area is described by Stäblein (1969), Hauser (*unpubl.*), Harasimuk (1987), Repelewska-Pękalowa (1987) and Pekala (1987). Climatic conditions are presented by Rodzik and Stepko (1985) as well as by Rodzik and Ryzyk (1987).

Soil profiles were studied in places characteristic for dry tundra (which predominates in Calypsostranda: Fig. 1; Pl. 1, Fig. 1), and also for wet and very wet tundra. Samples were collected from locations with dense vegetation and for a comparison also with tussock vegetation. In dry tundra moss and vascular plants as *Salix polaris*, *Polygonum viviparum*, *Cerastium arcticum*, *Silene acaulis*, *Saxifraga oppositifolia* and *Saxifraga caespitosa* prevail.

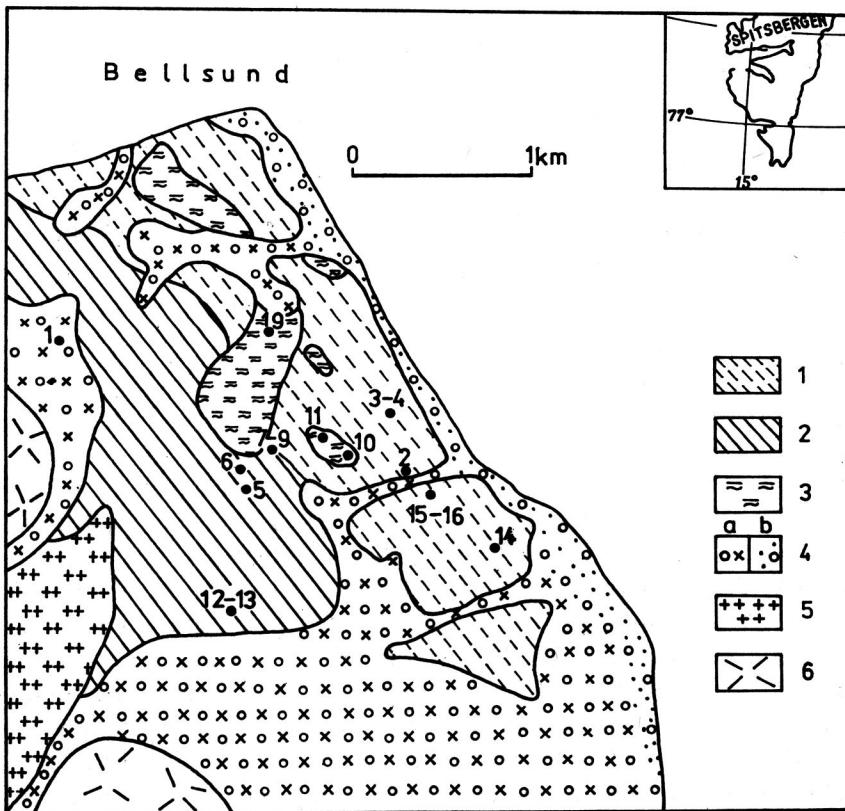


Fig. 1. Scheme of soil distribution in Calypsostranda (with locations and numbers of analyzed soil profiles): 1 — complex of poorly developed and brown soils formed from slightly loamy and loamy sands, 2 — gley and poorly gley soils formed from light and medium loam, and locally lithosols, 3 — gley soils formed from medium, heavy loam and clay, locally peat and shallow peaty soils, 4 — areas devoid of soils: a — stones and gravels, b — gravels and sands, and locally regosols; 5 — mountains, 6 — glaciers

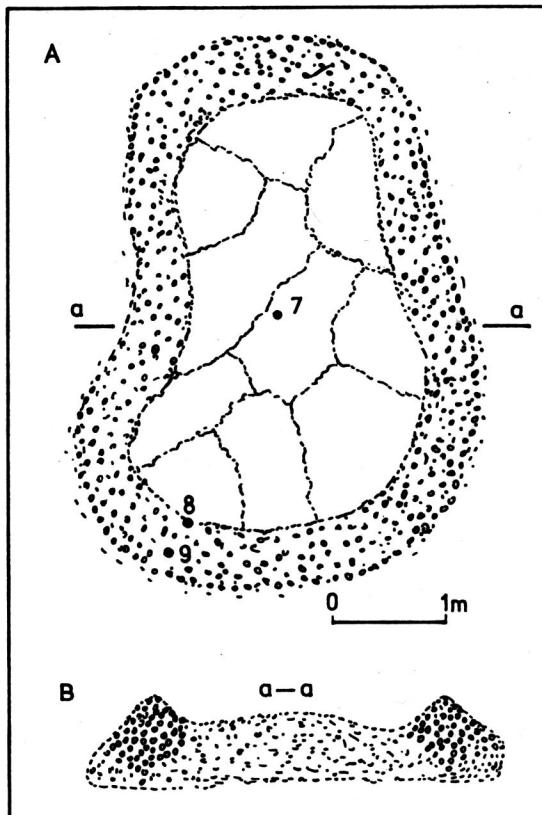


Fig. 2. Stone polygon with locations and numbers of profiles A — general view, B — cross-section (a—a)

In wet and very wet areas there are grass-moss communities with grass species as *Phippsia algida*, *Deshampsia alpina* and *Arctophila fulvia* (Rzętkowska 1987). During field works an additional section was prepared across a stony polygon surrounded by a stone circle (Fig. 2; Pl. 1, Fig. 2). Samples of soils accompanied by loamy outflows were also collected, as well as from centers and along fissures of fissure polygonal soils (Pl. 2).

The other profiles represent organic-mineral soil accompanied by characteristic turf hillocks, fossil soil, soil formed on roche moutonnée, mountain soil (outside the described area) and also soils formed on a pingo (in Chamberlindalen, several kilometers to the south from Calypsostranda). Several additional samples with undisturbed structure were collected to define physical properties of soils.

The first investigations included: grain size composition of soils by Bouyoucos' aerometric method, bulk density with a use of steel barrels of  $100 \text{ cm}^3$ , total porosity as well as capillary and noncapillary porosity with Leobell's apparatus, calcium carbonate content with Scheibler's apparatus,

pH by electrometric method, contents of organic mater by ignition, humus by Tiurin's method, total nitrogen after Kjeldahl and easily available phosphorus and potassium by Egner's method modified by Riehm (for organic soils the method of the Institute of Reclamation and Grasslands, Warsaw, was applied).

## Distribution of soils and their morphology

Field and laboratory studies of soils from Calypsostranda enabled to distinguish the following soil types:

1 — initial rocky soils (lithosols after FAO), 2 — initial loose soils (regosols after FAO), 3 — slightly developed soils, 4 — slightly developed brownish soils, 5 — slightly developed brown soils, 6 — brown soils (gelic cambisols after FAO), 7 — slightly gley soils, 8 — gley soils (gelic gleysols after FAO), 9 — shallow peaty soils, 10 — peat soils (histosols after FAO). These soils occur either on nondifferentiated or on polygonal surfaces. Genetic taxonomy of discussed soils is very similar to proposals of Plichta (1977). Some typical soils in the studied area are presented by the following profiles.

Profile 3 — slightly developed soil formed from silty loamy sands. Marine beach 29 m a.s.l., dry moss tundra, scarce vegetation (moss and cup-moss tufts, *Salix polaris*, *Saxifraga oppositifolia*, *Silene acaulis*, *Oxyria digyna*).

A<sub>1</sub> 0—6 cm — humus horizon, dark-grey almost black brownish (7.5 YR 4/2) heavy loamy sand with gravels and stones, bottom boundary unclear.

A<sub>1/C</sub> 6—20 cm — humus horizon with feature of parent rock, grey to dark-grey (5 YR 4/2) heavy loamy sand, more compact than A<sub>1</sub>, few plant roots, bottom boundary unclear.

C 20 cm — parent rock horizon, grey-dark yellowish (10 YR 5/2), heavy loamy sand with numerous stones, below 20 cm reaction with HCl.

Profile 6 — gley soil formed from silty medium loam. Marine beach 42 m a.s.l., humid tundra, rare vegetation (mainly moss, small grass tufts and *Saxifraga caespitosa*, *Saxifraga cernua*, *Cochlearia officinalis*).

A<sub>1</sub> 0—3 cm humus horizon, greenish-grey (2.5 GY 6/1) heavy silty loam with small amount of skeleton, bottom boundary unclear.

G 3—19 cm — gley horizon, greenish (5 GY 6/1) medium silty loam with small amount of skeleton, gradual transition into CG, beneath 12 cm reacts with HCl.

CG 19 cm parent rock horizon with gley features, greenish (5 GY 6/1) light loam with considerable admixture of gravels and stones.

Profile 16 — brown polygonal soil formed from low loamy sand, center of fissure polygon at 15 m a.s.l., covered in half by vegetation (moss, cup-moss, *Salix polaris*, *Silene acaulis*, *Polygonum viviparum*, *Saxifraga hirculus*).

Table 1

## Soil texture (nd — not determined)

profile no. and soil name	location and altitude (a.s.l.)	horizon	depth (cm)	skele- ton	grain size content in %							
					1—0.1 0.05	0.1— 0.05	0.05— 0.02	0.02— 0.005	0.005— 0.002	Σ <0.002	Σ 0.002	Σ <0.02
1	2	3	4	5	6	7	8	9	10	11	12	
3. poorly developed soil, heavy silty loamy sand, medium — deep	marine terrace tundra, rare vegetation, 29 m	A <sub>1</sub>	0—6	45	45	18	19	11	3	4	18	
		A <sub>1/C</sub>	10—15	35	48	19	14	14	4	1	19	
		C	20—25	55	55	17	11	11	3	3	17	
4. poorly developed brownish soil, light silty loamy sand, medium-deep	marine terrace tundra, dense vegetation, 28 m	A <sub>1</sub>	0—10	50	nd	nd	nd	nd	nd	nd	—	
		(B)C	10—20	65	58	16	13	9	3	1	13	
		D	20—30	75	30	21	25	16	4	4	24	
6. gley soil, medium silty loam	marine terrace tundra, rare vegetation, 42 m	A <sub>1</sub>	0—3	5	28	13	17	18	13	11	42	
		G	7—15	10	30	14	16	20	10	10	40	
		CG	20—30	50	48	10	12	14	10	6	30	
5. gley soil, medium silty loam medium-deep	marine terrace tundra, dense vegetation, 43 m	A <sub>1</sub>	0—4	45	25	15	18	17	16	9	42	
		C <sub>1</sub> G	8—15	20	28	15	13	15	16	13	44	
		C <sub>2</sub> G	20—30	70	40	14	17	13	8	8	29	
19. gley soil, clay medium-deep	ancient bay, bottom, very wet tundra, 17 m	A <sub>1</sub> G	0—5	40	28	39	15	11	4	3	18	
		G <sub>1</sub>	10—20	15	10	7	17	19	13	24	56	
		G <sub>2</sub>	27—32	30	10	18	16	19	14	23	56	
7. gley polygonal soil light loam, center of stone polygon, 34 m deep		A <sub>1</sub>	0—4	55	45	12	10	15	9	9	33	
		G <sub>1</sub>	7—14	40	48	11	10	14	10	7	31	
		CG <sub>2</sub>	18—25	25	40	17	11	18	9	5	32	
		CG <sub>2</sub>	45—55	35	43	14	11	19	9	4	32	
8. gley polygonal soil, heavy silty loamy sand, deep	periphery of stone polygon, 34 m	A <sub>1</sub>	0—5	75	58	14	11	12	4	1	17	
		G	10—20	25	38	19	13	18	8	4	30	
		CG	35—45	30	38	17	13	18	9	5	32	
9. gley polygonal soil heavy loamy sand, medium-deep	beneath stone circle, 34 m	A <sub>1</sub>	0—5	85	65	9	10	12	3	1	16	
		G	7—13	35	53	14	11	15	4	3	22	
		CG										
12. poorly gley polygonal soil, light silty loam, medium-deep	marine terrace, center of loamy outflow, 68 m	A <sub>1</sub>	0—10	20	43	14	17	19	6	1	26	
		G	14—20	40	43	16	13	18	6	4	28	
		CG	25—30	50	53	4	8	16	11	8	35	
13. poorly gley polygonal soil, silty clay, medium-deep	marine terrace, beneath vegetation circle, 68 m	A <sub>1</sub>	0—8	20	nd	nd	nd	nd	nd	nd	—	
		G	13—19	15	23	19	21	23	19	5	37	
		DG	22—28	45	50	12	11	15	7	5	27	
16. brown soil, slightly loamy silt loamy sand, medium deep	center of fissure polygon, 15 m	A <sub>1</sub>	0—8	20	53	31	8	8	0	0	8	
		(B)	15—25	10	63	18	7	9	2	1	12	
		C	30—40	55	70	12	7	9	1	1	11	
15. poorly developed soil, slightly loamy sand, deep	along frost fissure 15 m	A <sub>1</sub>	0—10	5	nd	nd	nd	nd	nd	nd	—	
		A <sub>1/C</sub>	15—25	10	nd	nd	nd	nd	nd	nd	—	
		C	45—50	45	83	5	4	6	1	1	8	
14. peat soil, medium deep	along frost fissure 18 m	AdT	0—12	0	nd	nd	nd	nd	nd	nd	—	
		T <sub>1</sub>	20—30	10	nd	nd	nd	nd	nd	nd	—	
		D	36—42	25	nd	nd	nd	nd	nd	nd	—	
		D	44—47	40	nd	nd	nd	nd	nd	nd	—	
2. Brown soil, silty loamy sand, center of fissure polygon, 22 m medium deep		A <sub>1</sub>	0—3	40	30	50	13	5	1	2	8	
		A <sub>1/(B<sub>1</sub>)</sub>	8—14	45	43	22	15	13	5	2	20	
		(B <sub>2</sub> )	16—19	35	65	13	9	8	2	3	13	
		D	25—35	25	90	0	7	1	1	1	3	
10. shallow peaty soil on loamy sand	turf hillock in solifluction kettle, 33 m	AdT	0—10	0	nd	nd	nd	nd	nd	nd	—	
		D	15—25	40	65	14	8	11	2	0	13	
11. peat soil on silty loamy sand	turf hillock in solifluction kettle, 34 m	AdT	0—10	0	nd	nd	nd	nd	nd	nd	—	
		T <sub>1</sub>	15—25	0	nd	nd	nd	nd	nd	nd	—	
		D	25—35	55	48	22	14	10	4	2	16	
1. fossil brown soil on light silty loam, shallow	Beneath terminal moraine of Scott Glacier, 66 m	A <sub>1</sub>	0—5	55	40	24	1	16	0	3	22	
		A <sub>1/(B)</sub>	15—20	55	38	19	16	16	6	5	27	
		D <sub>1</sub>	30—40	60	80	3	7	6	2	2	10	
		D <sub>2</sub>	60—70	60	90	5	1	2	0	2	4	
17. brown silty soil, very shallow	roche moutonnée between terminal and nival moraine of Scott Glacier, 85 m	A <sub>1</sub>	0—3	20	40	25	17	15	2	1	18	
		(B)	4—8	5	28	26	19	18	7	2	27	
		C	9—12	40	23	26	21	20	7	3	30	
18. initial loamy soil	bench at foot of Wijkanderberget, 500 m	(A <sub>1</sub> )	0—1	20	28	21	16	19	8	8	35	
		C	8—15	20	31	20	13	17	9	10	36	
20. slightly developed brown silty soil, very shallow	crest of pingo (Chamberlindalen)	(A <sub>1</sub> )	0—2	20	nd	nd	nd	nd	nd	nd	—	
		(B)	2—5	20	35	35	21	7	2	0	9	
		D	7—12	35	48	28	11	9	2	2	13	
21. slightly developed silty soil on clay, very shallow	slope of pingo (Chamberlindalen)	A <sub>1</sub>	0—4	5	nd	nd	nd	nd	nd	nd	—	
		A <sub>1/C</sub>	5—8	15	36	32	18	9	2	3	14	
		D	10—15	10	9	12	21	30	16	12	58	

Table 3

## Some chemical properties of soils

profile no. and soil name	horizon	depth (cm)	CaCO <sub>3</sub> (%)	pH 1 N KCl	organic matter (%)	C (%)	N (%)	C:N	available mg/100 g	
									10	11
1	2	3	4	5	6	7	8	9	P <sub>2</sub> O <sub>5</sub>	K <sub>2</sub> O
3. poorly developed soil, heavy silty loamy sand, A <sub>1</sub> medium deep	A <sub>1/C</sub>	0—6	0	7,3		3,84	0,32	12	2,4	4,9
	C	10—15	0	7,1		2,17	0,21	10	1,7	2,6
4. poorly developed brownish soil, light silty loamy sand medium deep	A <sub>1</sub>	20—25	10,5	7,1		0,99	0,09	11	4,1	4,1
	(B)/C	0—10	0	6,6	22,24		0,95		44,0	21,6
	D	10—20	0	6,7		2,45	0,23	11	2,4	3,9
6. gley soil, medium silty loam, medium deep	A <sub>1</sub>	20—30	0	7,1		2,36	0,21	11	2,2	2,9
	G	0—3	0	6,9		1,26	0,15	8	0,5	3,6
	CG	7—15	0,7	7,0		0,99	0,10	10	1,6	3,9
	CG	20—30	15,2	7,2		0,72	0,07	10	1,7	4,9
5. poorly gley soil, medium silty loam, medium deep	A <sub>1</sub>	0—4	0	6,7		1,48	0,18	8	2,1	3,2
	C <sub>1</sub> G	8—15	0	6,9		0,96	0,10	10	2,1	4,5
	C <sub>2</sub> G	20—30	4,2	7,0		0,66	0,07	9	1,7	2,4
19. gley soil, clay medium deep	A <sub>1</sub> G	0—5	5,5	7,5		2,07	0,22	9	2,1	2,4
	G <sub>1</sub>	10—20	13,5	7,6		0,66	0,07	9	1,7	6,5
	G <sub>2</sub>	27—32	13,5	7,5		0,74	0,08	9	3,5	6,5
7. gley polygonal soil, light loam, deep	A <sub>1</sub>	0—4	37,3	7,5		0,30	0,02	15	0,6	2,4
	G <sub>1</sub>	7—14	28,4	7,4		0,23	0,02	11	3,4	0,8
	CG <sub>2</sub>	18—25	27,9	7,6		0,43	0,04	11	1,0	0,8
	CG <sub>2</sub>	45—55	29,6	7,5		0,40	0,04	10	1,4	0,8
8. gley polygonal soil, heavy silty loamy sand, deep	A <sub>1</sub>	0—5	25,5	7,6		2,38	0,19	13	2,9	4,9
	G	10—20	30,7	7,7		0,34	0,04	8	2,1	0,8
	CG	35—45	29,6	7,6		0,35	0,05	7	1,7	1,6
9. gley polygonal soil, heavy loamy sand, medium deep	A <sub>1</sub>	0—5	23,6	7,5		1,33	0,14	9	1,7	3,2
	G	7—13	15,1	7,5		0,77	0,08	10	1,0	1,6
	CG									
12. slightly gley polygonal loam, medium deep	A <sub>1</sub>	0—10	0	7,1		1,30	0,13	10	1,0	1,0
	G	14—20	27,1	7,3		0,53	0,06	9	1,7	0
	CG	25—30	35,8	7,5		0,24	0,04	6	1,4	0
13. slightly gley polygonal soil, silty clay, medium deep	A <sub>1</sub>	0—8	0	7,3	14,06		0,50		24,0	7,6
	G	13—19	2,9	7,1		1,14	0,11	10	1,0	0,8
	DG	22—28	34,9	7,2		0,31	0,05	6	1,7	0,8
16. brown soil, slightly loamy silty sand, medium deep	A <sub>1</sub>	0—8	3,2	7,2		2,77	0,25	11	0,6	1,6
	(B)	15—25	0	7,2		0,78	0,08	10	0,6	0,8
	C	30—40	17,4	7,1		0,48	0,04	12	1,9	0,8
15. poorly developed soil, slightly loamy sand, deep	A <sub>1</sub>	0—10	0	6,6	15,97		0,45		44,0	27,0
	A <sub>1/C</sub>	15—25	0	6,7	10,73		0,40		36,0	14,3
	C	45—50	24,6	7,2		0,59	0,11	5	2,2	0,8
14. peat soil, medium deep	AdT	0—12	0	5,9	42,80		1,51		60,0	14,2
	T <sub>1</sub>	20—30	0	6,3	42,02		1,67		1700,0	9,6
	D	36—42	0	7,2	12,71		0,59		2000,0	6,5
	D	44—47	0	7,1	12,26		0,56		1600,0	6,5
2. brown soil, silty loamy sand, medium deep	A <sub>1</sub>	0—3	0	7,5		2,17	0,15	14	0,5	2,9
	A <sub>1(B)</sub>	8—14	0	7,4		2,15	0,17	13	0,7	1,3
	(B) <sub>2</sub>	16—19	10,6	7,3		1,08	0,08	13	2,1	0,8
	D	25—35	14,7	7,4		0,34	0,03	11	4,1	2,4
10. shallow peaty soil on loamy sand	AdT	0—10	0	6,4	53,13		2,22		30,0	25,7
	D	15—25	0	7,3		1,31	0,12	11	1,4	1,6
11. peat soil on loamy sand	AdT	0—10	0	5,3	60,02		1,85		53,0	16,3
	T <sub>1</sub>	15—25	0	5,6	61,81		1,68		36,0	28,2
	D	25—35	22,2	7,3		1,44	0,11	5	1,0	1,6
1. fossil brown soil, light silty loamy sand, shallow	A <sub>1</sub>	0—5	44,2	7,6		0,21	0,04	5	2,9	6,5
	A <sub>1/(B)</sub>	15—20	28,4	7,4		0,81	0,07	12	2,1	3,2
	D <sub>1</sub>	30—40	45,5	7,2		0,20	0,03	7	1,7	3,2
	D <sub>2</sub>	60—70	46,3	7,5		0,18	0,01	18	1,4	10,6
17. brown silty soil, very shallow	A <sub>1</sub>	0—3	0	7,3		2,06	0,20	10	1,2	2,6
	(B)	4—8	0	7,1		1,91	0,19	10	1,0	1,6
	C	9—12	0	7,0		1,91	0,18	11	0,3	1,6
18. initial loamy soil	(A <sub>1</sub> )	0—1	14,8	7,2		0,49	0,05	10	1,4	2,4
	C	8—15	11,1	7,5		0,45	0,06	8	4,8	1,6
20. poorly developed brown silty soil, very shallow	(A <sub>1</sub> )	0—2	0	6,6	13,97		0,53		108,0	26,0
	(B)	2—5	0	7,2		3,53	0,30	12	0,5	1,9
	D	7—12	0	7,1		2,00	0,15	13	1,0	1,6
21. poorly developed silty soil on clay, very shallow	A <sub>1</sub>	0—4	0	7,0	9,85		0,39		266,0	23,9
	A <sub>1/C</sub>	5—8	0	6,8		3,78	0,31	12	1,0	3,9
	D	10—15	1,0	6,8		1,06	0,12	9	2,7	4,2

A<sub>1</sub> 0—10 cm humus horizon, grey to dark grey (5 YR 4/2) low loamy silty sand with gravel and stones, bottom boundary clear, reacts with HCl.

(B) 10—26 cm — browning horizon, brown-grey (7.5 YR 4/4), light silty loamy sand, sometimes with flora roots, bottom boundary unclear.

C 26—45 cm — parent rock horizon, grey brownish (10 YR 5/3) light silty loamy sand with large amount of gravels and stones, strongly reacts with HCl.

Development of soils and their properties considerably depend on lithology of bedrock as well as frost segregation, ground swelling and cracking, solifluction, which are all typical processes in arctic areas. Many authors indicated that development of arctic soils depends more on cryogenic processes than on soil-forming processes. Local water conditions, microclimate and vegetation are also very important.

A considerable part of Calypsostranda does not possess any but occasionally initial soil cover (Fig. 1). Its higher marine beaches are composed of rock outcrops or debris with primitive soils and very rare vegetation (locally there are lithosols).

A lower part of the coastal plain and wide river valleys have sandy gravel trains with stone admixture (locally there are regosols).

## Physical and chemical properties of soils

Soils of Calypsostranda are composed of light, medium and heavy formations (Table 1). But light and medium, often silty loams significantly prevail. They are mainly gley and poorly gley soils, usually connected with wet tundra, stone polygons and loamy outflows. Poorly developed brown or brownish soils have usually a lighter grain-size composition (silty sands with gravel). Such soils are most common in dry tundra and coastal plain, from the present beach up to the marine terrace at 25—30 m a.s.l. (Fig. 1). Polygons, sometimes very large, and surrounded by frost fissures prevail there. The heaviest grain-size composition (clays and silty clay) has been found in gley soils of very wet tundra, and in a poorly developed soil at lateral part of a pingo.

Section across a stone polygonal soil shows that the heaviest granulometric composition is connected with the center of the polygon. A lighter composition was observed at its border and the lightest beneath a stone circle.

Grain size distribution in described soils contains relatively more coarse silt than colloidal clay (particles below 0.002 mm). It probably results in a low resistance of soils to crushing. Insignificant predominance of the fraction smaller than 0.02 mm was observed in some horizons (B) only.

Most investigated soils show a strong skeleton (to about 85%). But in a vertical distribution of gravels and stones distinct regularities are only occasionally visible.

Bulk density of studied soils changes from 1.3 g/cm<sup>3</sup> in light sandy soil with much organic matter, to 2.2 g/cm<sup>3</sup> in gley loamy soil in center

Table 2

## Physical properties of soils

profile no.	soil horizon	depth (cm)	bulk density (%)	total porosity (%)	bulk-capillary porosity (%)	noncapillary porosity
4	A <sub>1</sub>	3—10	1,3	60	50	10
	(B)	12—18	1,5	53	44	9
7	A <sub>1</sub>	0—4	1,9	39	29	10
	G <sub>1</sub>	7—14	2,0	34	29	5
	CG <sub>2</sub>	18—25	2,2	26	23	3
	CG <sub>2</sub>	47—53	2,1	27	26	1
12	A <sub>1</sub>	4—10	1,7	45	42	3
	G	14—20	1,9	39	35	4
16	A <sub>1</sub>	4—10	1,4	57	51	6
	(B)	15—21	1,5	54	45	9

of a stone polygon (Table 2). Total porosity of these soils is widely varied. In general the total porosity decreases whereas bulk density increases with depth. Highest and lowest non-capillary porosities were observed in the same profile at different horizons of gley and stony polygonal soils (profile 7). This soil indicated strong swelling properties.

Calcium carbonate occurs in most soils (Table 3) and its content increases with depth, reaching 46% in a brown fossil soil. No carbonates were found in brown and similar soils formed on roche moutonnée and pingo, and also in peat and shallow peaty soils. Carbonate reaction with 10% HCl was often weak and slow. It indicates that in described soils calcium carbonate is accompanied by magnesium carbonate.

Reaction of studied soils does not indicate a great differentiation (Table 3) and is usually neutral or basic. Only upper horizons of organic (peat and shallow peaty) soils are slightly acidic. In some profiles a small increase of pH with depth was noted. Soils under dense vegetation cover in dry and wet tundra have a slightly lower pH than under scarce vegetation. Similarly in fissure polygonal soil lower pH values are noted along fissure where more organic matter occurs. Such regularities were neither found in stone polygonal soils nor in soils of loamy outflows. Alkaline reaction of most soils on Calypsostranda is due to occurrence of carbonates in parent rocks, poor leaching and intensive frost processes which generate permanent vertical movements of soil material.

Highest content of organic matter (max. 61.8%) was found in peat and shallow peaty soils formed in turf hummocks and polygonal soils around fissures (Table 3). Organic horizons are also the thickest in these soils. In poorly developed brownish soils in dry tundra but only under dense floristic cover, a considerable content of organic matter was noted as well.

In dry and wet tundra more organic carbon was found under dense than under thin vegetation cover. A small content of carbon was recorded in soils connected with stone polygons (especially with their centers). They are gley and poorly developed gley soils. A little bit more carbon occurs in brown soils (except brown fossil soil). Organic carbon is present in all soil profiles, but most of it occurs in their upper horizons.

Total nitrogen content is differentiated (from 2.22 to 0.01%) and depends largely on humus content. The lowest nitrogen content is noted in stone polygonal soils (in centers of polygons) and in fossil soils. More nitrogen occurs in mineral-organic and organic soils. Content of total nitrogen usually decreases with depth. C/N ratio varies widely (even in a single section) but usually it is narrow or sometimes very narrow.

Phosphorus and potassium contents are low in mineral soils. Only in organic soils with abundant bird and reindeer excrements, a considerable quantity of phosphorus was found. Potassium content is however very differentiated.

Soils in the described area as well as in the Hornsund region (Szerszeń 1968, 1974) are rich in organic carbon in contrast to available phosphorus and potassium in mineral soils. Noted pH was higher than in Hornsund area. The studied soils indicate regular patterns of carbonates, pH, organic carbon and total nitrogen in their profiles.

## Conclusions

1. Cryogenic processes as well as gleization, peat-formation and browning are very important in soil-forming processes of the arctic periglacial zone. They influenced typology and properties of described soils.
2. Strong skeleton, general occurrence of carbonates, neutral or basic reaction, low content of available phosphorus and potassium, and high content of organic carbon are the characteristic features of the soils in Calypsostranda.
3. Despite of slightly distinct morphology of the studied soils, regularities in vertical distribution of some components have been distinguished.

## References

- Andreiev V. N. 1980. Rastitielnost i pochvy subarkticheskoi tundry. — Izdatielstvo „Nauka”, Novosibirsk, 209 pp.
- Baranowski S. and Szerszeń L. 1968. Some properties of sub-fossil mineral and organic deposits from the region of Werenskioldbreen, Vestspitsbergen. In: K. Birkenmajer (ed.), Polish Spitsbergen Expeditions 1957—1960. — Wyd. Geol., Warszawa, 239—274.
- Boratyński K., Kowaliński S. Szerszeń L. and Wilk K. 1968. Preliminary research on the fractional composition of soil humus from the Hornsund area, Vestspitsbergen. In: K. Birkenmajer (ed.), Polish Spitsbergen Expeditions 1957—1960. — Wyd. Geol., Warszawa, 229—237.
- Harasimiuk M. 1987. Współczesny rozwój wybrzeży południowego Bellsundu i fiordu Recherche (Zachodni Spitsbergen). — XIV Symp. Polarne, Ref. i Komunikaty: 99—102. Lublin.
- Jahn A. 1946. O niektórych formach gleb strukturalnych Grenlandii Zachodniej. — Przegl. Geogr., 20: 73—89.
- Klimowicz Z. and Uziak S. 1987. Gleby na obszarze Calypsostrandy (Spitsbergen Zachodni). — XIV Symp. Polarne, Ref. i Komunikaty: 200—202. Lublin.
- Låg J. 1980. Special peat formation in Svalbard. — Acta Agric. Scand., 30: 205—210.
- Låg J. 1981. Humus accumulation in steep-slopes at the inner part of Söudre Strømfjord, Greenland. — Acta Agric. Scand., 31: 242—244.
- Låg J. 1983. Jordforgiftning fra gruveavfall brukt som fyllmasse i Longyearbyen, Svalbard. — Jord og Myr, 6: 208—211.
- Linell K. A. and Tedrow J. C. F. 1981. Soil and permafrost surveys in the Arctic. — Clarendon Press, Oxford, 279 pp.
- Pękal K. 1987. Rzeźba i utwory czwartorzędowe przedpolu lodowców Scotta i Renarda (Spitsbergen). — XIV Symp. Polarne, Ref. i Komunikaty: 84—87. Lublin.
- Plichta W. 1977. Systematics of soils of the Hornsund region, West Spitsbergen. — Acta Univ. N. Copernici, Geografia, 13: 175—180.
- Repelewska-Pękalowa J. 1987. Rozwój równiny nadmorskiej pod wpływem procesów erozji (na przykładzie Calypsostrandy, rejon Bellsundu, Zachodni Spitsbergen). — XIV Symp. Polarne, Ref. i Komunikaty: 103—105. Lublin.
- Rieger S. 1974. Arctic Soils. — In: J. D. Ives and R. G. Barry (eds.). Arctic and Alpine Environments. — Methuen, London, 749—769.
- Rodzik J. and Ryzyk E. 1987. Zróżnicowanie przestrzenne warunków termiczno-wilgotnościowych południowego obrzeżenia Bellsundu w sierpniu 1986 roku. — XIV Symp. Polarne, Ref. i Komunikaty: 195—199, Lublin.
- Rodzik J. and Stepko W. 1985. Climatic conditions in Hornsund (1978—1983). — Pol. Polar Res., 6: 561—576.
- Rzętkowska A. 1987. Wstępna charakterystyka roślinności rejonu Calypsostranda (Wedel-Jarlsberg Land, Spitsbergen). — XIV Symp. Polarne, Ref. i Komunikaty: 218—220. Lublin.
- Stäblein G. 1969. Bellsund — Expedition 1968. Sonderdruck aus „Gipfel in der Arktis”, Bericht über die Hanseatische Spitzbergenkundfahrt des Deutschen Alpenvereins 1968. Hrsg. von W. Schlüter, J. Ruf. u.a., Bielefeld.
- Stäblein G. 1977. Arktische Böden West-Grönlands: Pedovarianz in Abhängigkeit vom Geoökologischen Milieu. — Polarforschung, 47: 11—25.
- Szerszeń L. 1965. Studia nad glebami strefy klimatu arktycznego na przykładzie południowo-zachodniego Spitsbergenu. — Zesz. Nauk. WSR Wrocław, 60: 39—79.
- Szerszeń L. 1968. Preliminary investigations of soil cover in the region of Hornsund. In: K. Birkenmajer (ed.), Polish Spitsbergen Expeditions 1957—1960. — Wyd. Geol., Warszawa, 217—227.

- Szerszeń L. 1974. Wpływ czynników bioklimatycznych na procesy zachodzące w glebach Sudetów i Spitsbergenu. — Rocznik glebozn., 25: 53—95.
- Szerszeń L. and Chodak T. 1983. Clay minerals in soils of the arctic climatic zone exemplified by South-west Spitsbergen. — 5th Meet. Europ. Clay Groups, Prague, 1983: 445—451.

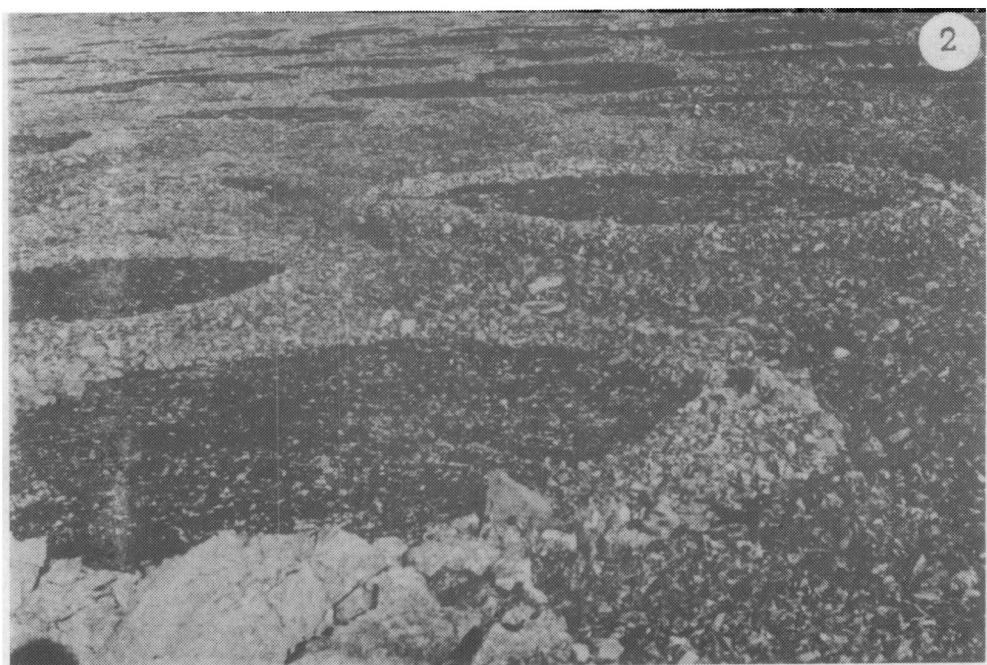
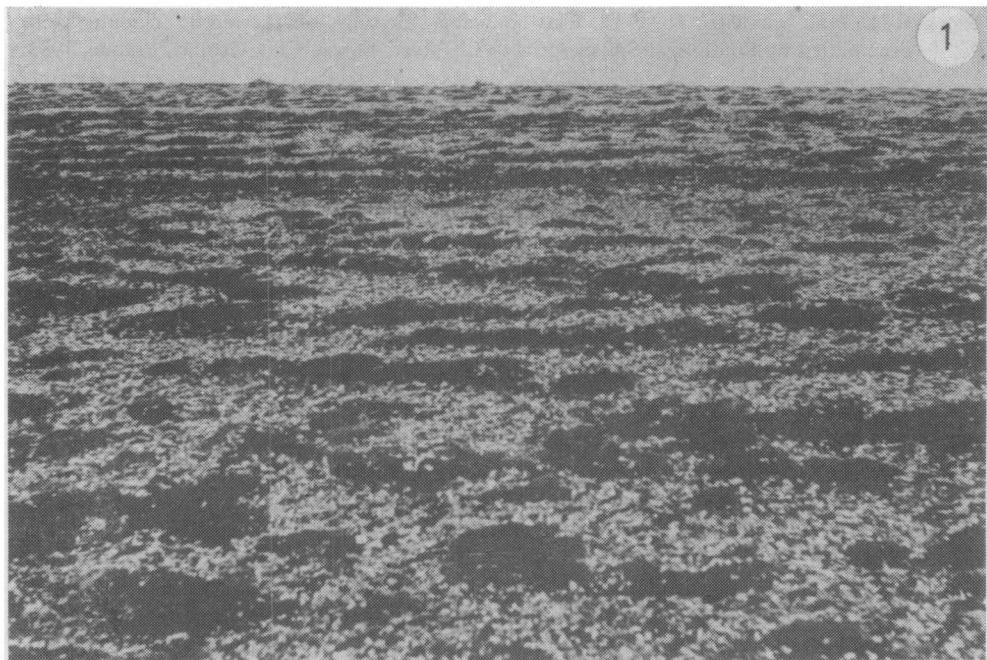
Received June 9, 1987

Revised and accepted October 21, 1987

## Streszczenie

Badania gleb przeprowadzono w 1986 roku w miejscach charakterystycznych dla tundry arktycznej, to jest na tundrze suchej (pl. 1, fig. 1), tundrze wilgotnej oraz bardzo wilgotnej. Ponadto zbadano gleby poligonalne związane z poligonami kamienistymi (fig. 2; pl. 1, fig. 2), poligonami szczelinowymi (pl. 2, fig. 2) oraz wylewami gliniastymi (pl. 2, fig. 1). Wydzielono dziesięć jednostek glebowych (fig. 1).

W tworzeniu się gleb w arktycznej strefie peryglacialnej obok procesów glebotwórczych (glejowych, torfotwórczych, brunatnienia), istotną rolę odgrywają procesy kriogeniczne. Te ostatnie wywierają duży wpływ na typologię a także na właściwości omawianych gleb. Dominującymi cechami gleb Calypsostrandy są: silna szkieletowość, dość powszechnie występowanie węglanów, odczyn obojętny lub zasadowy, mała zawartość (w glebach mineralnych) przyswajalnych form fosforu i potasu oraz stosunkowo znaczna węgiel organiczny (tab. 1—3). Pomimo mało wyraźnej (wizualnie) morfologii badanych gleb, zaznaczają się pewne prawidłowości w pionowym rozmieszczeniu niektórych składników.



1. Dry moss tundra (photo K. Pękala)  
2. Stone polygons (photo Z. Klimowicz)



1. Loamy outflows (photo K. Pękala)  
2. Fissure polygons (photo Z. Klimowicz)