

# Calcareous plankton bio-chronostratigraphy of the Maltese Lower Globigerina Limestone member

NICCOLO' BALDASSINI\*, ROBERTO MAZZEI, LUCA MARIA FORESI, FEDERICA RIFORGIATO AND GIANFRANCO SALVATORINI

Dipartimento di Scienze della Terra, Università di Siena, Via Laterina 8, 53100 Siena, Italy.

\*E-mail: [n.baldassini@hotmail.it](mailto:n.baldassini@hotmail.it), 00393393514706

## ABSTRACT:

Baldassini, N., Mazzei, R., Foresi L.M., Riforgiato, F. and Salvatorini, G. 2012. Calcareous plankton bio-chronostratigraphy of the Maltese Lower Globigerina Limestone member. *Acta Geologica Polonica*, **63** (1), 105–135. Warszawa.

The planktonic foraminifera and calcareous nannofossil biostratigraphy of the Maltese Lower Globigerina Limestone member has been investigated. The member was dated to early planktonic foraminiferal P22 Zone and nannofossil NP25 Zone (upper Chattian). A climate-stratigraphic approach, based on the quantitative analyses of calcareous nannofossils, was used additionally to achieve a more precise chronology. The species *Coccolithus pelagicus* (diameter <11 µm) and the genus *Umbilicosphaera* were selected for the recognition of cold and warm surface waters intervals respectively. The ratio of their percentages enabled the construction of a Climatic Factor (CLF) curve. The CLF values were consistent with a warm climatic phase, which is probably represented by the portion of the oxygen stable isotope curve of Miller *et al.* above the Oi2c event and below the beginning of the cooling trend that culminates in the Mi1 event. Considering these two climatic events and the upper boundary of the NP25 Zone, it can be inferred that the deposition of the Lower Globigerina Limestone member took place between 25.1 and 24.3 Ma.

**Key words:** Calcareous plankton; Biostratigraphy; Chattian; Lower Globigerina Limestone member; Maltese Archipelago.

## INTRODUCTION

The Globigerina Limestone formation (Murray 1890) crops out widely in the Maltese Archipelago, both on Malta and Gozo islands (Text-fig. 1). The type locality and section of the formation was established by Felix (1973) at Il-Qaws (NW of Dingli, on the southwestern coast of Malta Island).

The thickness estimates of the Globigerina Limestone formation vary from c. 20 m to as much as 60 m (Hyde 1955; Felix 1973), 70 m (Bennett 1979; Challis

1979) or even 207 m (House *et al.* 1961; Pedley 1978a, b; Pedley *et al.* 1976). Based on lithological variability and the presence of two distinct phosphatic beds, the formation is commonly subdivided into the Lower, Middle and Upper Globigerina Limestone informal members ("sub-formations" of Rizzo 1932, who introduced the tripartite subdivision of the formation). The phosphatic beds are placed variously at the top of the Lower and Middle Globigerina Limestone members (Pedley *et al.* 1976; Carbone *et al.* 1987; Jacobs *et al.* 1996), at the bases of the Middle and Upper Globigerina Limestone

members (Giannelli and Salvatorini 1972, 1975; Bennett 1979; Challis 1979; Rose *et al.* 1992), or at the base and at the top of the Middle Globigerina Limestone member (Kienel *et al.* 1995; Rehfeld and Janssen 1995).

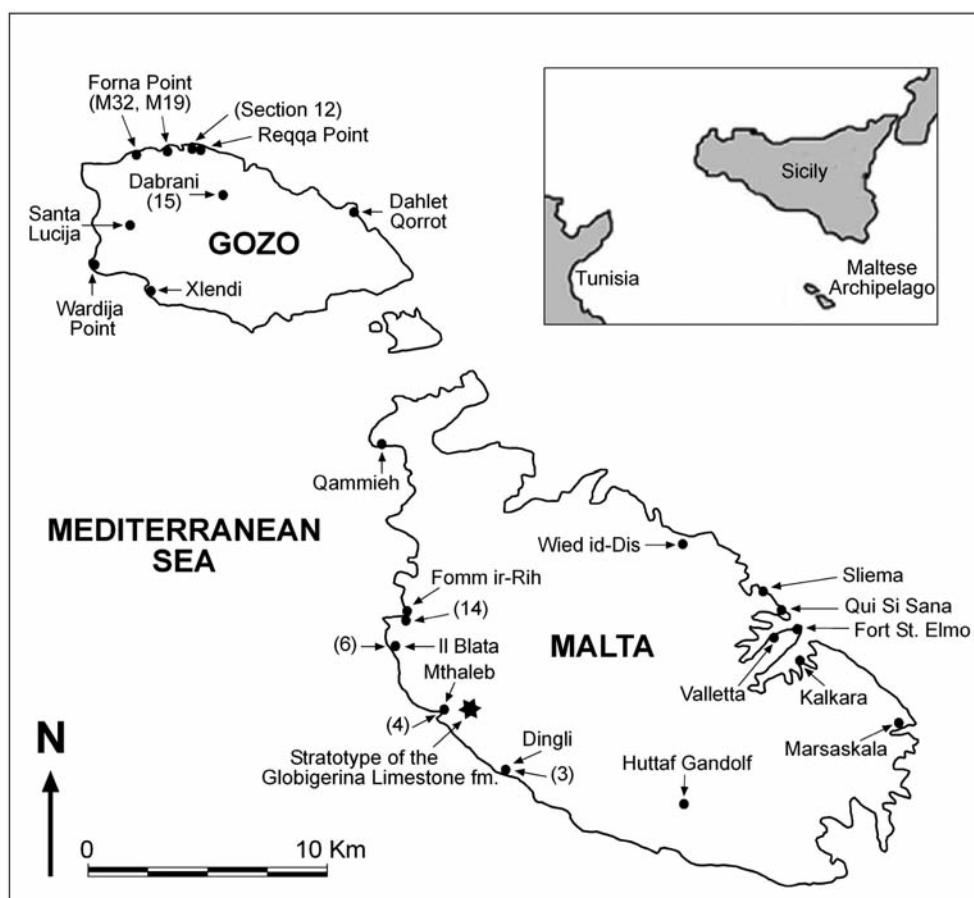
Despite the fact that the first chronostratigraphic data on the Globigerina Limestone formation were published in the late nineteenth century (e.g., Wright 1855); Gregory (1890-91); and Cooke (1896a, b), the first reliable reports date from the 1970s and 1980s century, when Giannelli and Salvatorini (1972) and Mazzei (1980, 1986) recognised the Aquitanian and Burdigalian age of the Middle Globigerina Limestone member and the Langhian age of the Upper Globigerina Limestone member. Moreover, these authors identified sedimentary hiatuses associated with the two phosphatic beds. Their chronostratigraphic conclusions were substantially confirmed by Felix (1973). The researches concerning the Lower Globigerina Limestone member enabled Giannelli and Salvatorini (1972) to suggest a Chattian (Late Oligocene) age for the member, and Felix (1973) to refer it to the Aquitanian (Early Miocene). Subsequently, even though the Chattian age of the member was confirmed by several Sr-isotope analyses (Pratt 1990, unpublished Ph.D. thesis; Rose *et al.* 1992; Jacobs

*et al.* 1996; Follmi *et al.* 2008), the attribution of the member to the Aquitanian was quoted in a number of reports (e.g., Rose 1974; Pedley *et al.* 1976; Challis 1979; Menesini 1979a, b; Theodoridis 1984; Pedley 1993; Kienel *et al.* 1995; Rehfeld and Janssen 1995).

The aim of this study is to clarify the age of the Lower Globigerina Limestone member by quantitative and semiquantitative analyses of the planktonic foraminiferal and calcareous nannofossil assemblages. This research was carried out in several sections (including those sampled by Giannelli and Salvatorini 1972), located both on Malta and Gozo islands (Text-fig. 1). The preliminary results were presented by Foresi *et al.* (2007).

## LITHOSTRATIGRAPHIC OUTLINE

The Lower Globigerina Limestone member is composed of massive to thick-bedded, yellow-brown biotrital limestones (biomicrosparites and biomicrites). In the lower part of the member the limestones are coarse-grained, yellow-brown and strongly bioturbated. In the upper part they are fine-grained, marly and pale yellow



Text-fig. 1 Location of the Maltese sections considered in the present work. In brackets are the sections of Giannelli and Salvatorini (1972) and Mazzei (1980, 1986)

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in colour. Macrofossils are frequent, represented mostly by bryozoans and the echinoid *Scutella* in the lower part, and by pectinid bivalves (mainly *Flabellipecten*) and echinoids (e.g. *Schizaster*) in the upper part. Challis (1979) recognized the *Scutella* biofacies at the base of the unit (the “*Scutella* bed” in the literature) and the *Schizaster/Hemiaster* biofacies in the remaining part.

Pedley (1975, unpublished Ph.D. thesis; and 1989, 1992) reports the maximum thickness of the member as well above 100 m in the Valletta area, with a distinct reduction in thickness near the Comino Straits and on the Rabat Axis (Malta Island), where the member even disappears. In the Malta and Gozo islands, Pedley (1993) reported thicknesses between 0 and 80 m and 5 and 40 m respectively. Bennett (1979), Rose (1985), and Dart *et al.* (1993), reported thicknesses of up to 32 m and up to 20 m respectively.

Previously, the so-called “*Scutella* bed” or “transitional bed” (a layer with a high concentration of large flat echinoids) was considered the boundary level between the underlying Lower Coralline Limestone formation and the Globigerina Limestone formation. Some authors place the base of the Lower Globigerina Limestone member at the top of the “*Scutella* bed”. Below the “*Scutella* bed” Carbone *et al.* (1987) defined the “Basal Globigerina Limestone Phosphatic Bed”, which is particularly well developed in the northern part of Malta Island and in the western part of the Island of Gozo. This phosphatic bed covers the upper surface of an ubiquitous hardground that Bennett (1979, 1980 unpublished Ph.D. thesis) named “Terminal Lower Coralline Limestone Hardground”. Felix (1973), Bennett (1980), Carbone *et al.* (1987) and Rose *et al.* (1992) located the base of the Lower Globigerina Limestone member just above this hardground.

The upper boundary of the Lower Globigerina Limestone member is marked by a prominent hard bed composed of abundant brown phosphatic nodules, glauconite granules, phosphatized (molluscs, echinoids, pteropods, corals, etc.) and non-phosphatized (ostreids, pectinids, *Nautilus*, *Eupatagus*, *Spatangus*, bryozoans, etc.) fossils and shark teeth, in a more or less abundant pale coloured biomicrite matrix. The upper surface of this bed is mainly planar and characterized by a phosphatic polycyclic layer (Pedley and Bennett, 1985) which covers the clasts. The bed was referred to as “C1”, or “Lower main conglomerate” (Pedley 1975; see also Pedley *et al.* 1976), “Qammieh bed” (Bennett 1980), or as “Lower Main Phosphorite Conglomerate Bed” (Pedley and Bennett 1985). Subsequently, it was formally referred to as the “Qammieh Conglomerate Bed” by Rose *et al.* (1992). The up to one metre thick Qammieh Conglomerate Bed overlies a phosphatized

hardground (the “Terminal Lower Globigerina Limestone Hardground” of Rose *et al.* 1992). The top of the Lower Globigerina Limestone member is defined by some authors at the top of the Qammieh Conglomerate Bed (Giannelli and Salvatori 1972, 1975; Bennett 1979; Challis 1979; Rose 1985; Rose *et al.* 1992; Pedley 1993; Dart *et al.* 1993; Rehfeld and Janssen 1995; Kienel *et al.* 1995), and by some others at the top of the underlying hardground (Rizzo 1932; Roman and Roger 1939; House *et al.* 1962; Pedley *et al.* 1976; Carbone *et al.* 1987; Jacobs *et al.* 1996).

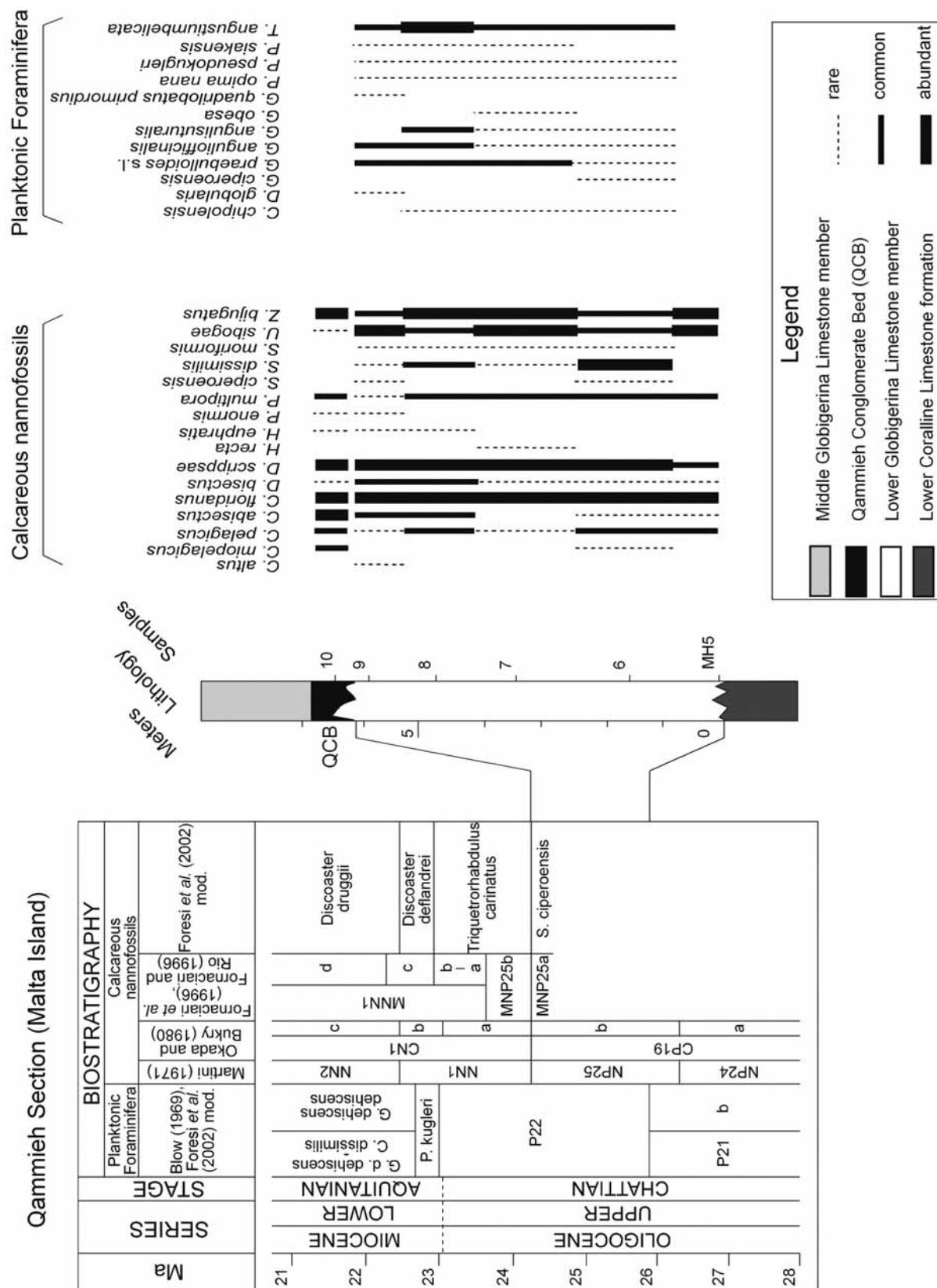
In this study, we consider the upper surfaces of the Terminal Lower Coralline Limestone hardground and Terminal Lower Globigerina Limestone hardground as the base and the top of the Lower Globigerina Limestone member respectively. The arguments for such definitions are the following: (i) the Terminal Lower Coralline Limestone hardground is ubiquitous, readily recognizable, and it marks a discontinuity surface; (ii) all of the several levels rich in *Scutella*, present in the Lower Coralline Limestone formation–Lower Globigerina Limestone member transition lie above the Terminal Lower Coralline Limestone hardground and show lithological characters typical of the Lower Globigerina Limestone member.

## SECTIONS STUDIED

Twenty sections of the Lower Globigerina Limestone member of the Maltese Archipelago (twelve on Malta Island and eight on the Island of Gozo) were studied (Text-fig. 1). Although most of the sections were included in the earlier report by the authors (Foresi *et al.* 2007), all have been subsequently re-sampled. Further details on the sections are summarized in Tables 1 and 2 (in the Appendix).

## BIOSTRATIGRAPHY

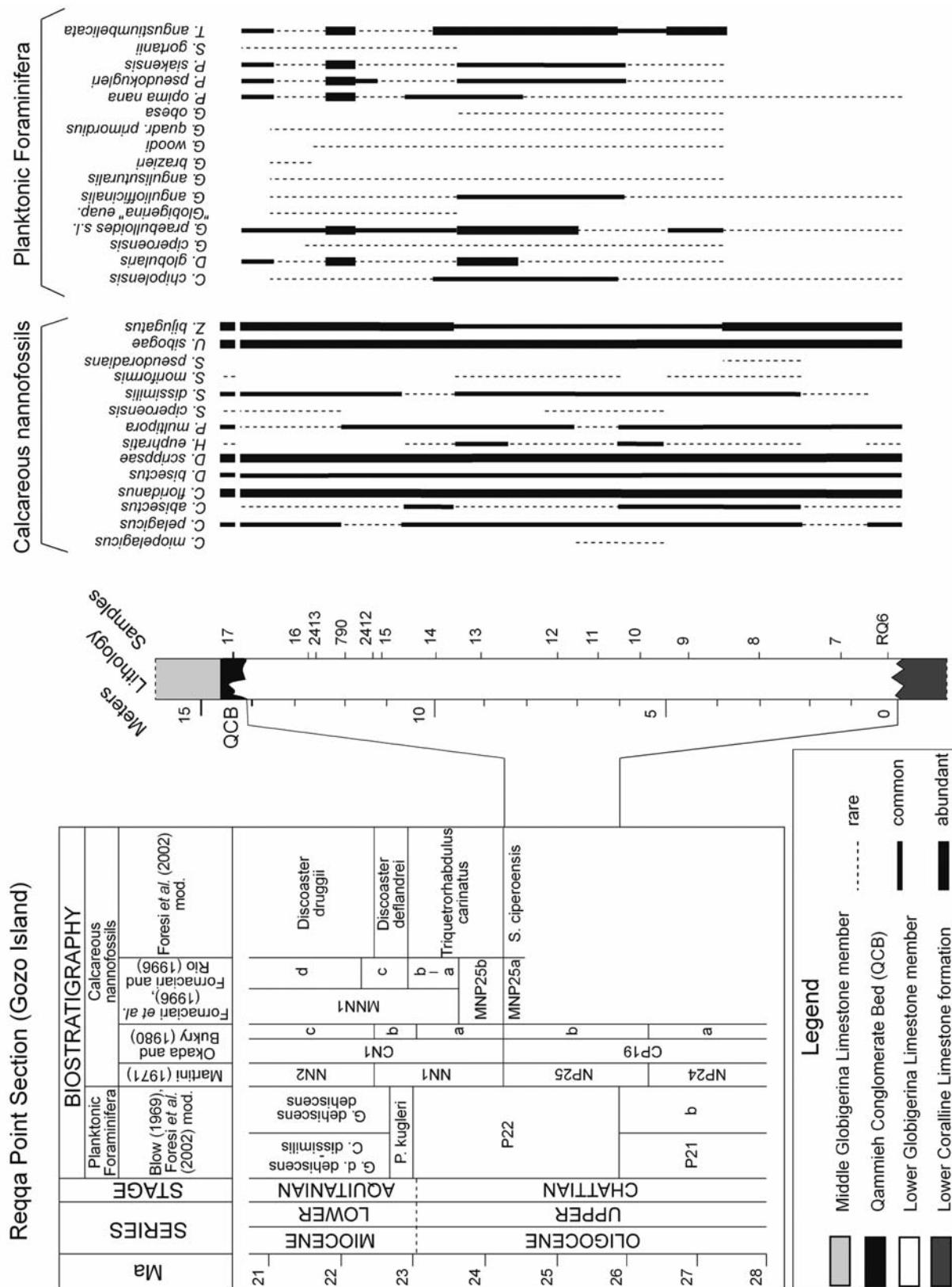
A semi-quantitative analysis was carried out on the planktonic foraminiferal and calcareous nannofossil assemblages (see Tables 3, 4, 5, 6 in the Appendix). Regarding the calcareous nannofossils, a quantitative analysis (a calculation of the percentage of each taxon in relation to a total of 300 specimens) was added for some sections (see Table 7 in the Appendix). The zonal schemes of Blow (1969), Martini (1971), Okada and Bukry (1980), Fornaciari *et al.* (1990), Fornaciari and Rio (1996) and Foresi *et al.* (2002) are compared and used in the biostratigraphic discussion (Text-figs 2, 3, 7).



Text-fig. 2 Semiquantitative calcareous plankton content of the Lower Globigerina Limestone member which crops out in the section of Qammieh (Malta Island).

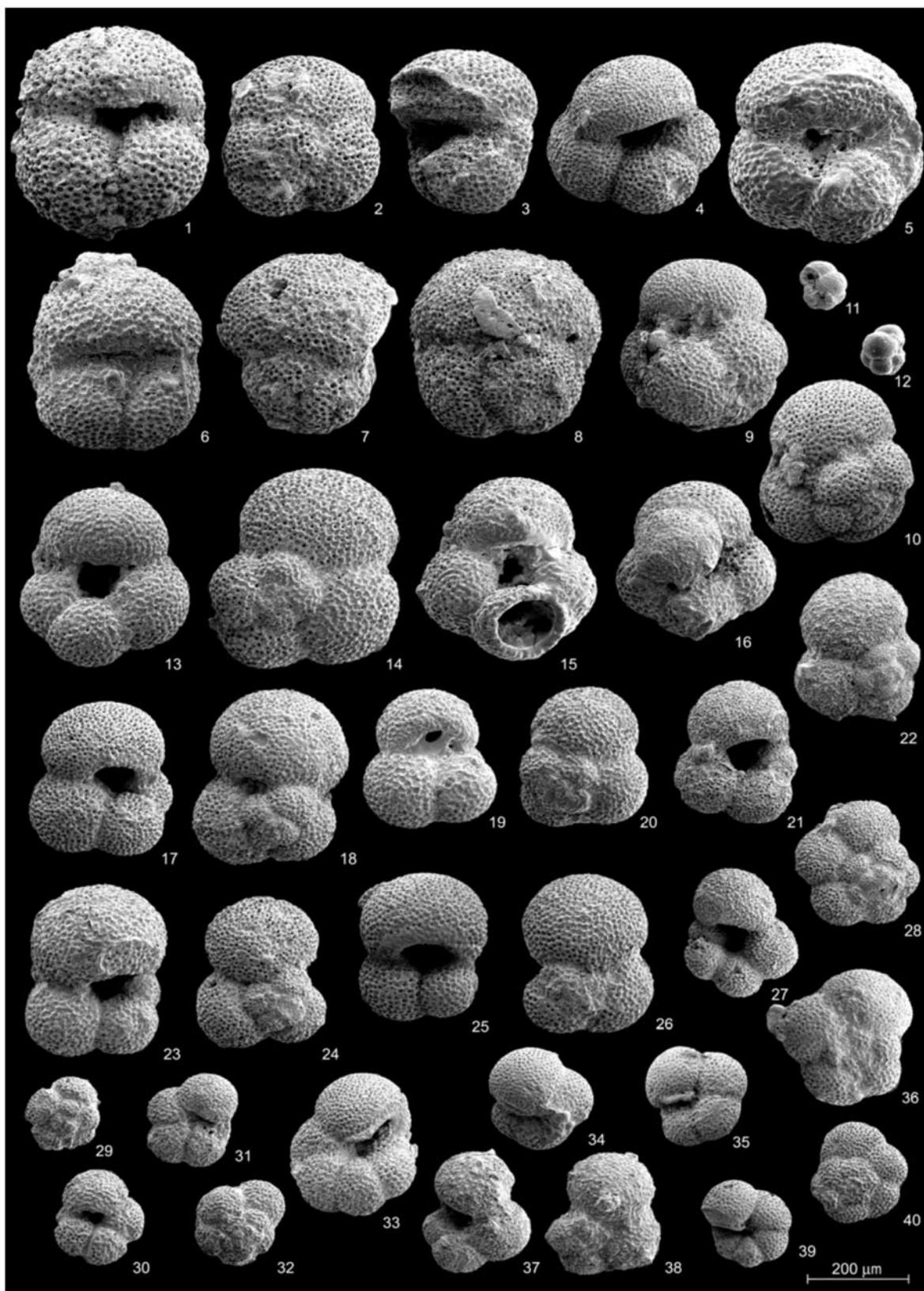
R – rare; C – common; A – abundant

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**Text-fig. 3** Semiquantitative calcareous plankton content of the Lower Globigerina Limestone member which crops out in the section of Rqqa Point (Gozo Island).

R = rare; C = common; A = abundant



## Planktonic foraminifera

The assemblages are generally of low diversity and represent a single biostratigraphic unit. All of the taxa identified are listed (genus-level taxonomy follows Iaccarino *et al.* 2005, with some exceptions) and their known stratigraphic ranges are expressed in terms of Blow's (1969) zonal scheme. The ranges of the taxa recognized through quantitative and semiquantitative analyses in the sections of Malta and Gozo islands are shown in Text-figs 2 and 3, and Tables 3 and 4; the critical forms are illustrated in Text-fig. 4.

*Cassigerinella chipolensis* (Cushman and Ponton, 1932): P17–N14 (N9 in the Mediterranean area, see Foresi *et al.* 2001). Common in many samples, rarely abundant (Text-fig. 4.11, 4.12).

*Dentoglobigerina globularis* (Bermúdez, 1961): P18–N6 (Last Occurrence (LO) occurring in Chron C6Bn.1r, at 22.8 Ma in Berggren *et al.* 1995). Often common or abundant (Text-fig. 4.13, 4.14).

*Globigerina ciperoensis* Bolli, 1954: P18–N4. According to ATNTS 2004 (Astronomical Tuned Neogene Time Scale; Gradstein *et al.* 2004a, b) the recalculated age estimate of the LO of *G. ciperoensis* is 22.90 Ma, which is within the earliest chronozone, N4, according to the timescales of both Berggren *et al.* (1995) and Gradstein *et al.* (2004a, b). Present in many samples, but only common in a few (Text-fig. 4.27, 4.28).

*Globigerina praebulloides* s.l., including the three subspecies *G. praebulloides leroyi* Blow and Banner, 1962 (P15–N8), *G. praebulloides occlusa* Blow and Banner, 1962 (P13–N21) (Text-fig. 4.23, 4.24) and *G. praebul-*

*loides praebulloides* Blow, 1959 (P15–N18) (Text-fig. 4.21, 4.22). Ubiquitous, and abundant in many samples.

“*Globigerina*” *euapertura* Jenkins, 1960: P15–P22 (N3) (up to N8?). Often common, rarely abundant (Text-fig. 4.2, 4.3).

“*Globigerina*” *sellii* Borsetti, 1959. (synonyms: *Globigerina clarae* Bermúdez, 1961; *Globigerina oligocaenica* Blow and Banner, 1962): P19–N6. Present in many samples, but common in only a few (Text-fig. 4.6, 4.7).

“*Globigerina*” *tripartita* Koch, 1926: P14–N6. Rare and only present in a few samples from some sections (Text-fig. 4.1).

*Globigerinella obesa* (Bolli, 1957): P19–Recent. Present, but rare, in many samples (Text-fig. 4.37, 4.38).

*Globigerinoides quadrilobatus primordius* Blow and Banner, 1962: from the base of P22 (N3) (from P21 in Biolzi 1985; Biolzi *et al.* 1981; Boersma and Premoli Silva 1991; Berggren *et al.* 1995 reported its First Occurrence (FO) in Chron C8r at 26.7 Ma) to N4 (up to N7?). It occurs in most samples, but is never abundant (Text-fig. 4.17, 4.18).

*Globoturborotalita anguliofficinalis* (Blow, 1969): P16–P22 (N4?). Ubiquitous and often common (Text-fig. 4.31, 4.32).

*Globoturborotalita angulisuturalis* (Bolli, 1957): base of P21 (N2) (P20 (N1) for Jenkins and Orr 1972)–P22 (N3) (N4?). Less frequent than its immediate predecessor *G. anguliofficinalis* (see Blow 1969) and rarely common (Text-fig. 4.29, 4.30).

Text-fig. 4. Planktonic foraminifera from the Lower Globigerina member of sections on the islands of Malta and Gozo, Maltese Archipelago. 1 – “*Globigerina tripartita*” Koch. Umbilical view; Dingli section, sample 3205; 2–5 – “*Globigerina euapertura*” Jenkins. 2 – spiral view; Forna Point section (M19), sample 3224. 3 – side view, Forna Point section (M19), sample 3219. 4 – umbilical view; Forna Point section (M19), sample 3216. 5 – umbilical view; Sliema Point section, sample 3209; 6–8 – “*Globigerina sellii*” Borsetti. 6 – umbilical view; Sliema Point section, sample 3210. 7 – side view, Marsaskala section, sample 3196. 8 – spiral view, Forna Point section (M32), sample 3200. 9–10, 15–16 – *Subbotina gortanii* (Borsetti). 9 – umbilical-side view; Sliema Point section, sample 3212. 10 – spiral view; Forna Point section (M32), sample 3200. 15 – umbilical view; Dingli section, sample 3203. 16 – umbilical view; Forna Point section (M32), sample 3200; 11–12 – *Cassigerinella chipolensis* (Cushman and Ponton). Sliema Point section. 11 – umbilical view; sample 3212. 12 – spiral view, sample 3211; 13–14 – *Dentoglobigerina globularis* (Bermúdez). Sliema Point section, sample 3209. 13 – umbilical view. 14 – spiral view; 17–18 – *Globigerinoides quadrilobatus primordius* Blow and Banner. Sliema Point section. 17 – umbilical view, sample 3212. 18 – spiral view, sample 3209; 19–20 – *Globoturborotalita woodi* (Jenkins). Sliema Point section, sample 3209. 19 – umbilical view. 20 – spiral view; 21–22 – *Globigerina praebulloides praebulloides* Blow. Requa Point section, sample 21 – umbilical view. 22 – spiral view; 23–24 – *Globigerina praebulloides occlusa* Blow. 23 – umbilical view, Sliema Point section, sample 3209. 24 – spiral view; Marsaskala section, sample 3194; 25–26 – *Globoturborotalita brazieri* (Jenkins). Sliema Point section, sample 3209. 25 – umbilical view. 26 – spiral view; 27–28 – *Globigerina ciperoensis* Bolli. 27 – umbilical view; Sliema Point section, sample 3209. 28 – spiral view; Dingli section, sample 3207; 29–30 – *Globoturborotalita angulisuturalis* (Bolli). 29 – spiral view; Requa Point section, sample 2410; 30 – umbilical view; Dingli section, sample 3207; 31–32 – *Globoturborotalita anguliofficinalis* (Blow). Sliema Point section, sample 3209. 31 – umbilical view. 32 – spiral view; 33 – *Paragloborotalita siakensis* (LeRoy). Umbilical view; Sliema Point section, sample 3212; 34–36 – *Paragloborotalita opima nana* (Bolli). Sliema Point section. 34 – umbilical view, sample 3210. 35 – umbilical view, sample 3210. 36 – spiral view, sample 3212. 37–38 – *Globigerinella obesa* (Bolli). 37 – umbilical view; Xlendi section, sample MT13. 38 – spiral view; Debrani section, sample 3179; 39–40 – *Paragloborotalita pseudokugleri* (Blow). Sliema Point section, sample 3209. 39 – umbilical view. 40 – spiral view

*Globoturborotalita brazieri* (Jenkins, 1966): P22 (N3)–N8. Present in several samples, but only common or abundant in a few sections (Text-fig. 4.25, 4.26).

*Globoturborotalita woodi* (Jenkins, 1960): P18–N23. Present in several samples, but only common in a few (Text-fig. 4.19, 4.20).

*Paragloborotalia opima nana* (Bolli, 1957): P13–P22 (N3) (up to N12?). Its LO is dated at 21.7 Ma by Miller *et al.* (1985) and at 22.6 Ma by Zhang *et al.* (1993). However, both these ages fall within the N4 chronozone (fide Berggren *et al.* 1995 and Gradstein *et al.* 2004a, b). Ubiquitous, often common or abundant (Text-fig. 4.34–4.36).

*Paragloborotalia pseudokugleri* (Blow, 1969): FO in the lower half of P22 (N3) (at 25.9 Ma according to Berggren *et al.* 1995; at 26.3 Ma in Pearson and Chaisson 1997); LO in N4 (at 21.6 Ma according to Berggren *et al.* 1995; at 21.31 Ma in Gradstein *et al.* 2004a). Ubiquitous, often common or abundant (Text-fig. 4.39, 4.40).

*Paragloborotalia siakensis* (LeRoy, 1939): P19–N14. Well represented in many samples (Text-fig. 4.33).

*Subbotina gortanii* (Borsetti, 1959): P15–N4. Present in many samples, but common in only a few sections (Text-fig. 4.9, 4.10, 4.15, 4.16).

*Tenuitellinata angustumbilicata* (Bolli, 1957). Recorded from P16 to the Pleistocene. According to Li *et al.* (1992), who carried out an analysis of tenuitellid stock ranges from oceanic sites, the LO of this species occurred at the top of N5. Ubiquitous and generally abundant.

*Tenuitellinata uvula* (Ehrenberg, 1861): P21 (N2) – Recent. Very rare and present in only one section.

The planktonic foraminifera allow the studied succession of the Lower Globigerina Limestone member to be assigned to the P22 (N3) Zone. The main arguments are as follows: (i) *Paragloborotalia opima nana*, which occurs in the Lower Globigerina Limestone member, is a direct successor of *P. opima opima* (Bolli), whose LO marks the P21 (N2)/P22 (N3) zonal boundary; and (ii) the FO of *Paragloborotalia kugleri* (Bolli), which marks the P22 (N3)/N4 zonal boundary, as emended by Berggren and Miller (1988; see also Berggren *et al.* 1995), is above the top of the Lower Globigerina Limestone member. Only *P. pseudokugleri*, the immediate ancestor of *P. kugleri* (Blow 1969; Spezzaferri *et*

*al.* 1991, among others), is present in the Lower Globigerina Limestone member.

In addition to these arguments, the planktonic foraminifera from the Lower Globigerina Limestone member include elements that substantiate the biostratigraphic assignment of the member to the P22 (N3) Zone. Moreover, they suggest that the member represents rather the lower part of the zone, with its upper part undocumented. The following arguments are of importance: (i) the presence of taxa with their FOs almost at the base (*Globigerinoides quadrilobatus primordius*) or at least in the lowermost part (*Paragloborotalia pseudokugleri*) of the P22 (N3) Zone (*Globoturborotalita obesa*, *G. woodi*, *G. brazieri*, *Tenuitellinata uvula*); (ii) the presence of taxa characterized by their LO at the base of the P22 (N3) Zone (*Globoturborotalita anguliofficinalis*, *Subbotina gortanii gortanii*). The latter species, which also occur in the upper part of the sections, make it possible to restrict the stratigraphic range of the Lower Globigerina Limestone member to the lower part of the zone; (iii) the upper portion of the P22 (N3) Zone is additionally excluded by the occurrence of a low number of *Globigerinoides quadrilobatus primordius* specimens (relatively common only in one level of the Xlendi section and in some samples of the Sliema and Marsaskala sections); in fact, this taxon is recorded as being abundant only in the upper part of this zone.

### Calcareous nannofossils

Calcareous nannofossils were found in all of the samples studied. The assemblages are rather poor taxonomically and have highly variable states of preservation. The taxa recognised are listed in Text-figs 2 and 3 and in Tables 5, 6 and 7; the critical forms are illustrated in Text-fig. 5.

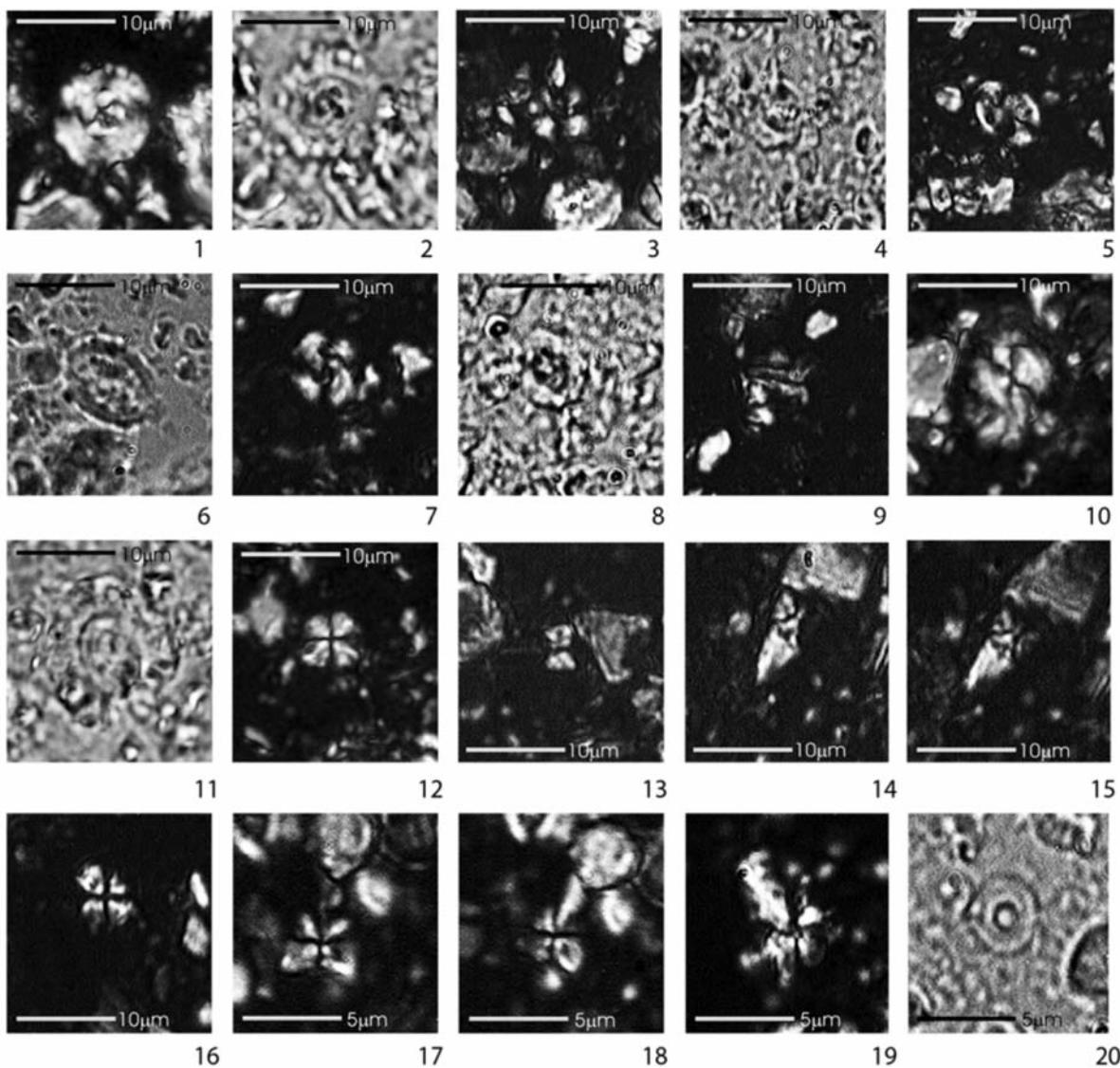
*Coccolithus pelagicus*, *Cyclicargolithus floridanus*, *Dictyococcites scrippsae*, *Pontosphaera multipora*, *Sphenolithus dissimilis*, *Umbilicosphaera sibogae* and *Zygrhablithus bijugatus* are generally common to abundant in almost all of the sections; *Dictyococcites bisectus* and *Helicosphaera euphratis* are rare; *Cyclicargolithus abisectus* and *Helicosphaera obliqua* are present in many samples but with a limited number of specimens; all other taxa are rare, sporadic, or very rare (see Tables 5, 6 and 7 in the Appendix).

Although no significant biostratigraphic events were found within the Lower Globigerina Limestone member, the calcareous nannofossils allow the studied succession to be assigned to the NP25 Zone. The critical arguments are as follows: (i) the co-occurrence of *Cyclicargolithus abisectus*, *Dictyococcites bisectus*, *D. scrippsae*, *Helicosphaera recta* and *Zygrhablithus bi-*

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*jugatus* is commonly indicative of Martini's (1971) NP24–NP25 interval (see Martini 1971; Bukry 1973, 1975, 1978; Martini and Muller 1986; Muller 1976; Bizon and Muller 1979; Perch-Nielsen 1985; Fornaciari *et al.* 1990, among others); (ii) in Martini's (1971) standard zonation, the FO of *Sphenolithus ciperoensis* marks the base of the NP24 Zone, while the LO of the same taxon marks the top of the NP25 Zone. Both the FO and LO of *S. ciperoensis* are also used in the low latitude zonation of Okada and Bukry (1980), to identify the lower

and upper boundaries (respectively) of the CP19 Zone. Furthermore, this taxon is well documented in the Mediterranean (e.g., Catalano and Di Stefano 1996; Fornaciari and Rio 1996; among others); (iii) the LO of *Sphenolithus distentus* defines the boundary between the NP24 (CP19a Subzone) and the NP25 (CP19b Subzone) zones; (iv) *S. dissimilis* evolves from *S. moriformis* in the late NP24 Zone and continues up to the middle–late NN2 Zone (Perch-Nielsen 1985); (v) the FO of *Pontosphaera enormis* and *Triquetrorhabdulus carinatus* are



Text-fig. 5. Calcareous nannofossils from the Lower Globigerina Limestone member of sections on the islands of Malta and Gozo, Maltese Archipelago. 1–2 – *Cyclcargolithus abiseptus* (Muller), crossed and parallel nicols, sample MH6, Qammieh section; 3–4 – *Cyclcargolithus floridanus* (Bukry), crossed and parallel nicols, sample BL5, II Blata section. 5–6 – *Pontosphaera multipora* (Kamptner), crossed and parallel nicols, sample BL3, II Blata section; 7–8 – *Dictyococcites scrippsae* (Bukry and Percival), crossed and parallel nicols, sample MH8, Qammieh section; 9 – *Zygrhablithus bijugatus* (Deflandre), crossed nicols, sample BL5, II Blata section; 10–11 – *Cocco lithus miopelagicus* (Wallich), crossed and parallel nicols, sample BL5, II Blata section; 12 – *Sphenolithus moriformis* (Bonnemann and Stradner), crossed nicols, sample MH8, Qammieh section; 13–15 – *Sphenolithus ciperoensis* (Bramlette and Wilcoxon), 0°, 20°, 45°, crossed nicols, sample MH6, Qammieh section; 16 – *Sphenolithus dissimilis* (Bukry and Percival), crossed nicols, sample BL5, II Blata section; 17–19 – *Sphenolithus delphix* (Bukry), 0°–20°–45°, crossed nicols, sample MH6, Qammieh section; 20 – *Umbilicosphaera* sp., parallel nicols, sample BL3, II Blata section

close to the NP24/NP25 boundary (Martini 1981; Perch-Nielsen 1985). According to these authors, the FO of *P. enormis* can also be used to define the base of the NP25 Zone in high latitudes or in areas with poor connections to the open ocean.

Summarising, the constant and abundant presence of *Dictyococcites bisectus*, *D. scrippsae*, *Sphenolithus dissimilis* and *Zygrhablithus bijugatus*, as well as the rare to very limited occurrence of *S. ciperoensis*, *Helicospaera recta*, *Pontosphaera enormis* and *Triquetrorhabdulus carinatus*, and the absence of *S. distentus*, allow all the sections studied to be assigned to NP25 Zone, excluding, most probably, its uppermost and lowermost parts.

This biostratigraphic result based on calcareous nannofossils agrees well with those obtained from the planktonic foraminifera (see Text-figs 2 and 3).

To clarify the biostratigraphic attribution, it was decided to highlight the calcareous plankton content (for calcareous nannofossils it is considered very rare to rare if it is <3%; common, between 3% and 10%; and abundant if it is >10%) of two sections (one located in Malta and one in Gozo) where the Lower Globigerina Limestone member is relatively thick and its base and top can be observed (Text-figs 2 and 3).

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The dating of the Maltese sections to the lower part of the foraminiferal P22 (N3) Zone and to the nannofossil NP25 Zone (probably with the exception of its upper and lowermost parts) definitely establishes the Chattian age of the Lower Globigerina Limestone member (Text-figs 2 and 3). This historically shared correspondence between the P22 (N3) (almost all) and NP25 zones and the Chattian age has recently been referred by the International Union of Geological Sciences (IUGS), through the acceptance of the Global Stratotype Section and Point (GSSP) for the base of the Neogene Period (that is the Oligocene/Miocene, as well as the Chattian/Aquitanian boundary) proposed by Steininger *et al.* (1997). The GSSP of the Paleogene/Neogene Period-System boundary was placed 35 m from the top of the Lemme-Carrosio Section in the Pedimont Basin (northern Italy), close to the base of Subchron C6Cn.2n, dated at 23.80 Ma (according to the Time Scale of Berggren *et al.* 1995). Billups *et al.* (2004) subsequently proposed a revised age of 23.03 Ma for this boundary and it was also reported in the Astronomically Tuned Neogene Time Scale (ATNTS) (Gradstein *et al.* 2004a, b). From a biostratigraphic viewpoint (see also Aubry and Villa 1996; Iaccarino *et al.* 1996; Steininger *et al.* 1996,

1997) the Oligocene/Miocene boundary is located in the uppermost part of the P22 (N3) Zone (2 m below the FO of *Paragloborotalia kugleri* s.s.) and in the NN1 Zone (MNN1b Subzone of Fornaciari and Rio 1996), close to the FO of *Sphenolithus capricornutus* and below the LO of *Sphenolithus delphix*.

## CLIMATE-STRATIGRAPHIC APPROACH

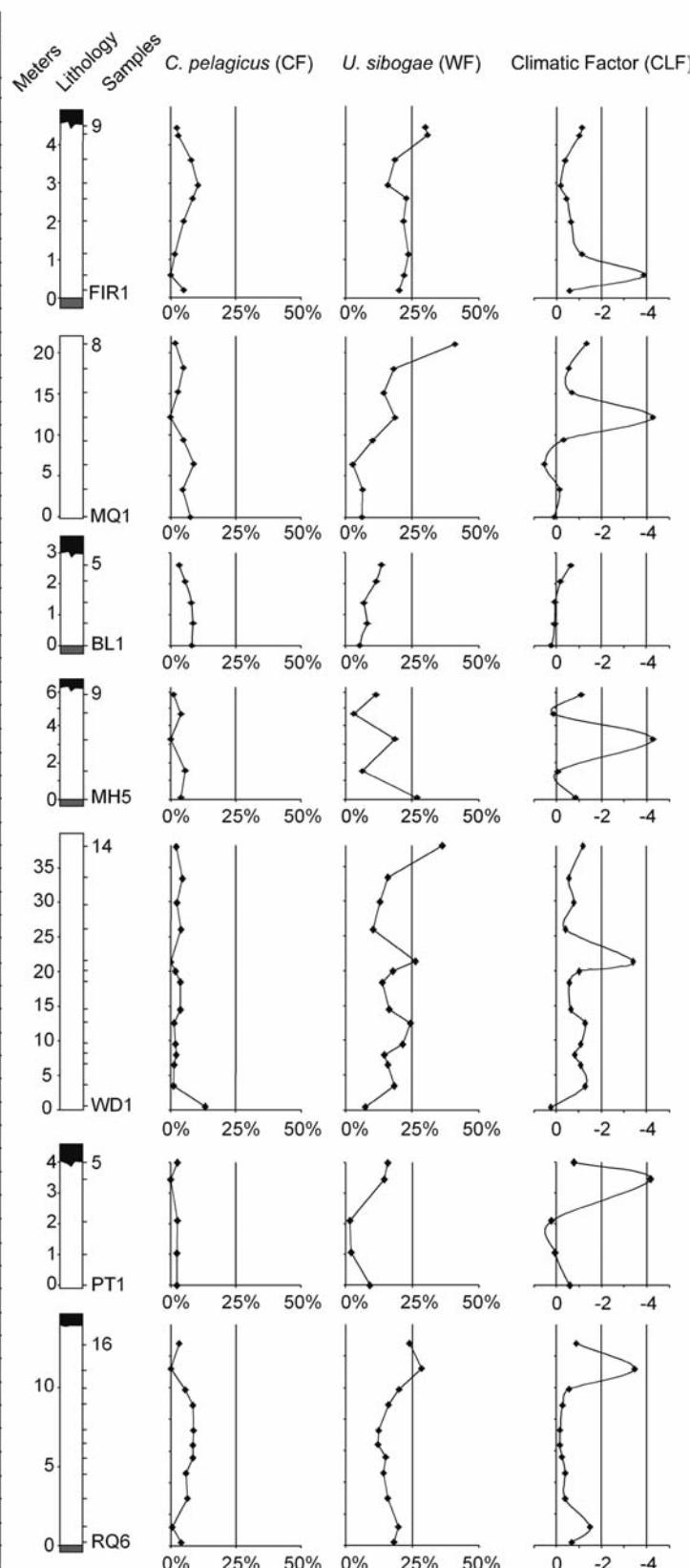
The quantitative analysis of the nannofossil assemblages (see Table 7 in the Appendix), carried out on selected sections (Island of Gozo: Reqqa Point and Wardija Point; Malta Island: Qammieh, Fomm Ir-Rih, Wied id Dis, Huttaf Gandolf and Il Blata), was the basis for a climate-stratigraphic approach.

Of the nannofossil taxa recognised in the sections studied, two extant forms were selected: the species *Coccolithus pelagicus* (<11 µm) and the genus *Umbilicosphaera*. Both show very strict surface water temperature preferences (Hasle 1960; Black 1965; McIntyre and Bé, 1967; McIntyre *et al.* 1970; Bartolini *et al.* 1970; Okada and Honjo 1973; Roth and Berger 1974; Geitzenauer *et al.* 1976; Haq and Lohmann 1976; Braarud 1979; Okada and McIntyre 1979; Haq 1980; Roth and Coulbourn 1982; Zhang and Siesser 1986; Kleijne *et al.* 1989; Honjo 1990; Sambleten and Schroeder 1992; Brand 1994; Winter *et al.* 1994; Baumann 1995; Wells and Okada 1996; Andrleit 1997; Flores *et al.* 1997, 1999; Findlay and Flores 2000; Geisen *et al.* 2002, 2004; Saez *et al.*, 2003; Sato *et al.* 2004; Hagino and Okada 2006; Marino *et al.* 2008; Bonnet *et al.* 2010). *Coccolithus pelagicus* (between 6 and 11 µm in dimensions, which Geisen *et al.* 2002, 2004 defined as *C. pelagicus* subsp. *pelagicus* and Saez *et al.* 2003 raised to the rank of species) characterizes the living assemblages of the Transitional and Subarctic floral zones; its optimum temperature is between 8° and 10°C (McIntyre and Bé 1967; McIntyre *et al.* 1970) in the Atlantic Ocean, and between 9° and 12°C (Roth and Berger 1974) in the Pacific Ocean. The genus *Umbilicosphaera* characterizes oceanic areas of Tropical and Subtropical zones where the sea surface temperature ranges between 24° and 28°C.

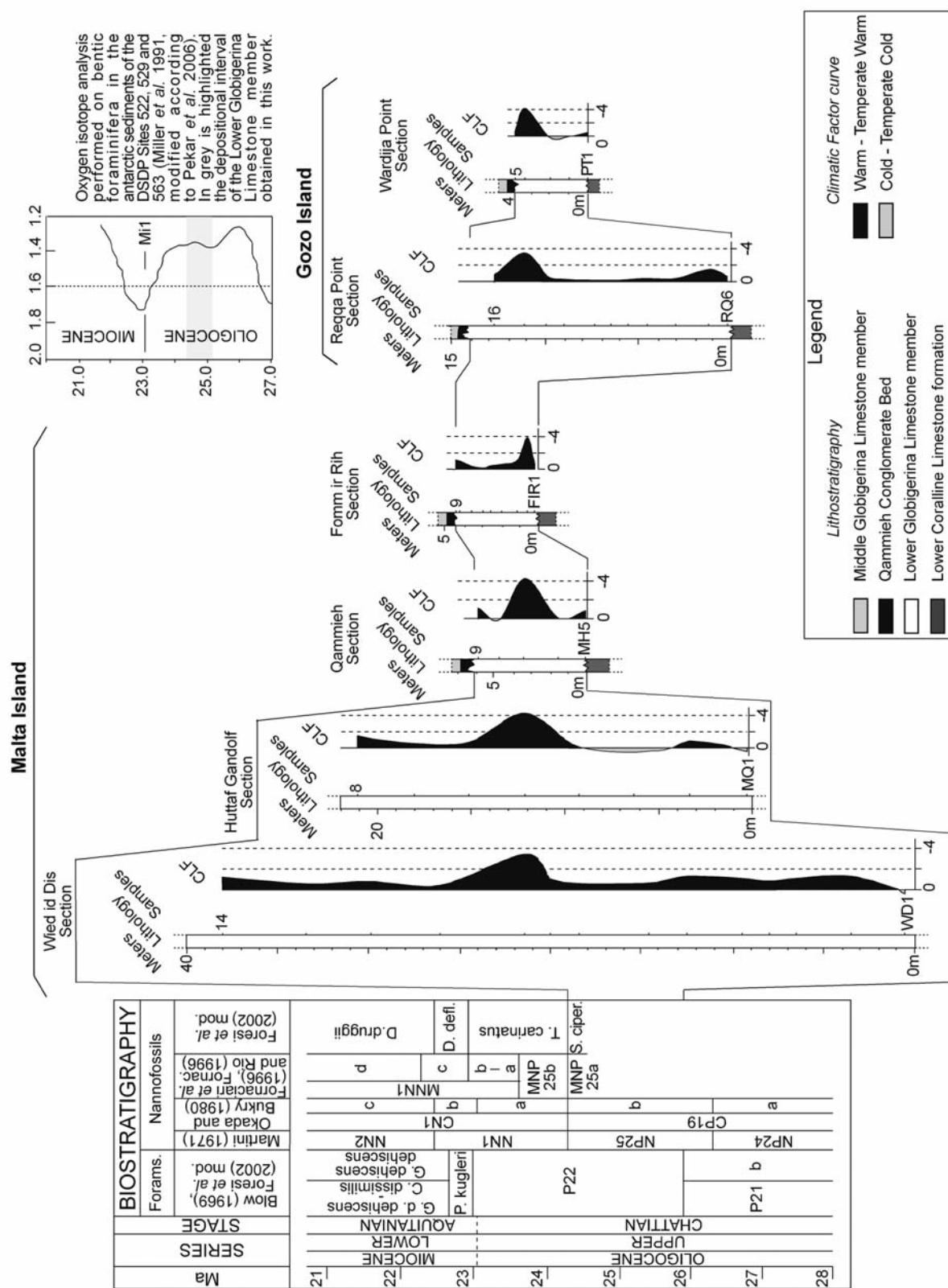
Based on temperature preferences of its modern counterparts *Coccolithus pelagicus* (<11 µm) and the genus *Umbilicosphaera* were selected as key species for the recognition of cold and warm surface waters respectively. Consequently, their relative abundance in relation to the total assemblages was used to construct Cold Factor (CF) and Warm Factor (WF) curves. An opposing trend of the two factors is generally evident (Text-fig. 6), thus proving the validity of the choice of these two

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Samples	<i>C. pelagicus</i> (CF)	<i>U. sibogae</i> (WF)	CLIMATIC FACTOR
Fomm ir-Rih			
FIR1	5.40%	20.03%	-0.6
FIR2	0.00%	21.89%	-3.86
FIR3	1.69%	23.54%	-1.14
FIR4	4.98%	21.76%	-0.64
FIR5	8.48%	22.80%	-0.43
FIR6	10.40%	15.86%	-0.19
FIR7	7.81%	18.48%	-0.37
FIR8	2.99%	30.76%	-1.01
FIR9	2.26%	29.86%	-1.12
Huttaf Gandolf			
MQ1	7.60%	5.97%	0.10
MQ2	4.68%	6.25%	-0.13
MQ3	8.86%	2.53%	0.54
MQ4	5.00%	10.00%	-0.30
MQ5	0.00%	18.51%	-4.27
MQ6	2.85%	14.28%	-0.70
MQ7	5.04%	18.06%	-0.55
MQ8	1.95%	41.02%	-1.32
II Blata			
BL1	8.33%	5.20%	0.20
BL2	8.20%	6.90%	0.07
BL3	7.87%	6.79%	0.06
BL4	6.50%	12.65%	-0.29
BL5	3.26%	13.40%	-0.61
Qammieh			
BL1	3.92%	26.62%	-0.83
BL2	5.55%	6.34%	-0.06
BL3	0.00%	18.49%	-4.27
BL4	4.08%	3.06%	0.12
BL5	0.94%	11.37%	-1.08
Wied id Dis			
WD1	13.27%	7.40%	0.25
WD2	0.94%	18.23%	-1.29
WD3	1.33%	16.00%	-1.08
WD4	2.24%	14.60%	-0.81
WD5	1.77%	21.35%	-1.08
WD6	1.24%	24.29%	-1.29
WD7	3.70%	16.40%	-0.65
WD8	3.61%	13.67%	-0.58
WD9	1.72%	17.73%	-1.01
WD10	0.00%	26.08%	-3.42
WD11	4.07%	10.29%	-0.40
WD12	2.25%	13.00%	-0.76
WD13	4.51%	15.78%	-0.54
WD14	2.12%	32.62%	-1.19
Wardija Point			
WD1	2.50%	9.10%	-0.56
WD2	2.40%	2.00%	0.08
WD3	2.70%	1.60%	0.23
WD4	0.00%	14.60%	-4.16
WD5	2.70%	15.70%	-0.76
Reqqa Point			
RQ6	3.84%	18.26%	-0.68
RQ7	0.62%	19.62%	-1.50
RQ8	6.35%	15.71%	-0.39
RQ9	5.70%	14.25%	-0.40
RQ10	8.29%	14.97%	-0.26
RQ11	8.29%	12.22%	-0.17
RQ12	8.59%	12.50%	-0.16
RQ13	8.35%	16.07%	-0.28
RQ14	5.56%	20.04%	-0.56
RQ15	0.00%	28.44%	-3.45
RQ16	3.24%	24.05%	-0.87



Text-fig. 6 Climate-stratigraphic framework of Lower Globigerina Limestone member



Text-fig. 7 Reconstruction of the Cold Factor (CF), Warm Factor (WF) and Climatic Factor (CLF) curves relative to the chosen sections. In figure, the Lower Coralline Limestone formation is indicated in light grey colour, the Lower Globigerina Limestone member in white and the Qammieh Conglomerate Bed in black colour

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taxa as climate indicators. The ratio between the percentages of CF and WF (represented in logarithmic form in order to highlight the negative values) was considered indicative of the Climatic Factor (CLF): high and low CLF values represent evidence of “cooling” and “warming” conditions respectively (Text-fig. 6).

Although the sampling performed is not adequate to define a climatic curve, it seems important to remember that all the CLF values are consistently negative and show no significant increase in any direction; consequently, the Lower Globigerina Limestone member can only have been deposited in the warm climatic phase of the oxygen isotope curve of Miller *et al.* (1987, 1989, 1991a, b, 2005) which follows the Oi2c event (according to Pekar *et al.* 2006 this  $\delta^{18}\text{O}$  event can be dated at about 25.1 Ma) and precedes the beginning of the cooling trend that culminates in the Mi1 event (Raffi *et al.* 2006, based on the calibration of the oxygen isotope curves of Zachos *et al.* 2001 and Billups *et al.* 2004 with the Astronomical Tuned Neogene Time Scale of Lourens *et al.* 2004, date this beginning at about 23.4 Ma).

By combining the obtained biostratigraphic data (lower part of the P22 (N3) Zone, the CLF curve and the reference isotope curve of Miller *et al.* (1991b modified according to Pekar *et al.* 2006), it is possible to define more accurately the time interval for the deposition of the member (Text-fig. 7). This deposition should start after 25.1 Ma (see above) and end well before 24.3 Ma (age estimated by Raffi *et al.* 2006 for the upper boundary of the NP25 Zone).

The obtained CLF curves were further applied to an attempt at fine-stratigraphic inter-correlation of particular sections (Text-fig. 7). In Text-fig. 7, it is shown that, although the sections are biostratigraphically referred to the same zone, their chronostratigraphic range within this zone varies. This is also supported by precise comparison of quantitative data from planktonic foraminifera and calcareous nannofossils (see Tables 3, 4, 5, 6 and 7 in the Appendix).

## CONCLUSIONS

The planktonic foraminiferal and calcareous nannofossil assemblages of the Lower Globigerina Limestone member of the Maltese Archipelago have confirmed that the member belongs to the lower part of the planktonic foraminiferal P22 (N3) Zone of Blow (1969) and to the nannofossil NP25 Zone of Martini (1971) (with the exception of its uppermost and lowermost parts). Consequently, the member is of late Chattian age.

A climate-stratigraphic approach based on quantitative analyses of calcareous nannofossil assemblages was applied to two sections on Gozo Island (Reqqa Point and Wardija Point) and five sections on Malta Island (Qammieh, Fomm Ir-Rih, Il Blata, Wied id Dis and Huttaf Gandolf). The species *Coccolithus pelagicus* (with diameter <11  $\mu\text{m}$ ) and the genus *Umbilicosphaera* were selected as key forms for the recognition of cold and warm surface waters respectively. The ratio of the percentages of these taxa enabled the construction of the Climatic Factor (CLF) curve.

All samples in the Lower Globigerina Limestone member provided CLF values consistent with a warm climatic phase which should be related to the Chattian heating documented by the oxygen isotope curve of Miller *et al.* (1987, 1989, 1991a, 1991b, 2005). In particular, our CLF curve matches well the portion of the Miller *et al.* curve above the Oi2c event ( $\delta^{18}\text{O}$  event highlighted by Pekar and Miller 1996; Sugarman *et al.* 1997; Shackleton *et al.* 1999; Pekar *et al.* 2002, 2006 and Miller *et al.* 2005) and below the beginning of the cooling trend that culminates in the Mi1 event. Based on the studies of Pekar *et al.* (2006) and Raffi *et al.* (2006), it is possible to date the two climatic events at about 25.1 Ma and 23.4 Ma respectively. Consequently, this is the time assumed for the deposition of the Lower Globigerina Limestone member.

Further clarification of the timing of the end of deposition of the Lower Globigerina Limestone member is implicitly provided by the biostratigraphic attribution based on the nannofossil assemblages; in fact, the recognition of the NP25 Zone (with the exception of its uppermost and lowermost parts) enables this sedimentary event to be dated well before 24.3 Ma (age estimated by Raffi *et al.* 2006 for the upper boundary of the NP25 Zone). Indirect confirmation in this sense comes from the study of Föllmi *et al.* (2008), which provides an age ranging from 23.5 to 22.0 Ma for the overlying “Lower Main Phosphate Bed” (the “Qammieh Conglomerate Bed” of Rose *et al.* 1992). It is also worth recalling the age of  $24.54 \pm 0.8$  Ma provided by Jacobs *et al.* (1996) for the Qammieh Conglomerate Bed.

The obtained CLF curve also appeared a very useful tool in a refined chronostratigraphic inter-correlation of particular sections. Based on the characteristics of the CLF curve (supported partly by quantitative micropalaeontological data), it can be inferred that particular sections of the Lower Globigerina Limestone member, although referable to a single biostratigraphic zone, represents various actual ranges. This suggests distinct sedimentary environments during the deposition of the Lower Globigerina Limestone member throughout the entire archipelago.

## REFERENCES

- Andruleit, H. 1997. Coccolithophore fluxes in the Norwegian-Greenland Sea: seasonality and assemblage alterations. *Marine Micropaleontology*, **31**, 45–64.
- Aubry, M.P. and Villa, G. 1996. Calcareous nannofossil stratigraphy of the Lemme-Carrosio Paleogene/Neogene Global Stratotype Section and Point. *Giornale di Geologia*, **58**, 51–69.
- Bartolini, C. and Gehin, C.E. 1970. Evidence of sedimentation by gravity-assisted bottom currents in the Mediterranean Sea. *Marine Geology*, **9**, M1–M5.
- Baumann, K.H. 1995. Morphometry of Quaternary *Coccolithus pelagicus* coccoliths from Northern North Atlantic and its paleoceanographical significance. In: Flores, J.A. and Sierra, F.J. (Eds), Fifth INA Conference Proceedings University of Salamanca, pp. 11–21. Salamanca.
- Bennett, S.M. 1979. A transgressive carbonate sequence spanning the Paleogene Neogene boundary on the Maltese Islands. *Annales Géologiques des Pays Helléniques*, **1**, 71–80.
- Bennett, S.M. 1980. Palaeoenvironmental studies in Maltese mid-Tertiary carbonates. Unpublished Ph.D. Dissertation, University of London.
- Berggren, W.A., Kent, D.V., Swisher III, C.C. and Aubry, M.-P. 1995. A Revised Cenozoic geochronology and chronostratigraphy. In: Berggren W.A., Kent D.V., Aubry M.-P. and Hardendol J. (Ed.), Geochronology, time scales and global stratigraphic correlation. *SEPM Special Publication*, **54**, 129–212.
- Billups, K., Palike, H., Channell, J.E.T., Zachos, J.C. and Shackleton, N.J. 2004. Astronomical calibration of the Late Oligocene through Early Miocene geomagnetic polarity time scale. *Earth Planetary Science Letters*, **224**, 33–44.
- Biolzi, M. 1985. The Oligocene/Miocene boundary in selected Atlantic, Mediterranean and Paratethyan sections based on biostratigraphic and stable isotope evidence. *Memorie di Scienze Geologiche*, **37**, 303–378.
- Biolzi, M., Bizon, G., Radovisc, A., Rögl, F. and Zachariasse, W.J. 1981. In search of the Paleogene/Neogene boundary stratotype. Part. 1. Potential boundary stratotype sections in Italy and Greece and a comparison with results from the deep-sea. Planktonic foraminifera. *Giornale di Geologia*, **44**, 98–104.
- Bizon, G., Bizon, J.J., Aubert, J. and Oertli, H.-J. 1972. Atlas des principaux foraminifères planctoniques du bassin méditerranéen Oligocène à Quaternaire, pp. 1–136. Edition Technip; Paris.
- Bizon, G. and Muller, C. 1979. Remarks on the Oligocene\ Miocene boundary based on the results obtained from the Pacific and Indian Oceans. *Annales Géologiques des Pays Helléniques*, **1**, 101–111.
- Black, M. 1965. Coccoliths. *Endeavour*, **24**, 131–137.
- Blow, W.H. 1969. Late Middle Eocene to Recent Planktonic foraminiferal biostratigraphy. In: Brönnimann, P. and Renz, H.H. (Eds), Proceedings of the First International Conference on Planktonic Microfossils, Geneva., pp. 199–421. E. Brill; Leiden.
- Boersma, A. 1984. Cretaceous – Tertiary planktonic foraminifers from the southeastern Atlantic Walvis Ridge Area, Deep Sea Drilling Project Leg 74. In: Moore, T.C. Jr., Rabinowitz, P.D. et al. (Eds), *Initial Reports of the Deep Sea Drilling Project*, **74**, 501–524.
- Boersma, A. and Premoli Silva, I. 1991. Distribution of Paleogene planktonic foraminifera – Analogies with the Recent? *Palaeogeography, Palaeoclimatology, Palaeoecology*, **83**, 29–48.
- Bonnet, S., De Vernal, A., Hillaire-Marcel, C., Radi, T. and Husum, K. 2010. Variability of sea-surface temperature and sea-ice cover in the Fram Strait over the last two millennia. *Marine Micropaleontology*, **74**, 59–74.
- Braarud, T. 1979. The temperature range of the non-motile stage of *Coccolithus pelagicus* in the North Atlantic region. *European Journal of Phycology*, **14**, 349–352.
- Brand, L.E. 1994. Physiological ecology of marine coccolithophores. In: Winter, A. and Siesser, A. (Eds), *Coccolithophores*, pp. 39–49. Cambridge University Press.
- Bukry, D. 1973. Low Latitude Coccolith Biostratigraphic Zonation. In: Edgar, N.T., Saunders, J.B. et al. (Eds), *Initial Reports of the Deep Sea Drilling Project*, **15**, 685–703.
- Bukry, D. 1975. New Miocene to Holocene stages in the ocean basins based on calcareous nannoplankton zones. In: Saito, T. and Burckle, L.H. (Eds), Late Neogene Epoch boundaries. *Micropaleontology Press, Special Publication*, 162–166.
- Bukry, D. 1978. Biostratigraphy of Cenozoic marine sediments by calcareous nannofossils. *Marine Micropaleontology*, **24**, 44–60.
- Carbone, S., Grasso, M., Lentini, F. and Pedley, H.M. 1987. The distribution and palaeoenvironmental of early Miocene phosphorites of southeast Sicily and their relationship with the Maltese phosphorites. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **58**, 35–53.
- Catalano, S. and Di Stefano, A. 1996. Nuovi dati geologici e stratigrafici sul Flysch di Capo d'Orlando nei Peloritani Orientali (Sicilia Nord-Orientale). *Memorie della Società Geologica Italiana*, **51**, 149–164.
- Challis, G.R. 1979. Miocene echinoid biofacies of the Maltese Islands. *Annales Géologiques des Pays Helléniques, Tome Hors Serie*, **1**, 253–261.
- Cooke, J.H. 1896a. Notes on the Globigerina Limestone of the Maltese Islands. *Geological Magazine*, **4** (3), 502–511.
- Cooke, J.H. 1896b. Contribution to the stratigraphy and palaeontology of the Globigerina Limestone of the Mal-

## CHRONOSTRATIGRAPHY OF THE MALTESE LOWER GLOBIGERINA LIMESTONE MEMBER

- tese Islands. *Quarterly Journal of the Geological Society of London*, **52**, 461–462.
- Dart, C.J., Bosence, D.W.J. and McClay, K.R. 1993. Stratigraphy and structure of the Maltese graben system. *Journal of the Geological Society, London*, **150**, 1153–1166.
- Felix, R. 1973. Oligo-Miocene stratigraphy of Malta and Gozo. *Mededelingen Landbouwhogeschool Wageningen*, **73**, 1–104.
- Findlay, C.S. and Flores, J.A. 2000. Subtropical Front fluctuations south of Australia (45°09'S, 146°17'E) for the last 130 ka years based on calcareous nannoplankton. *Marine Micropaleontology*, **40**, 403–416.
- Flores, J.A., Gersonde, R. and Sierro, J.F. 1999. Pleistocene fluctuations in the Agulhas Current retroflection based on the calcareous plankton record. *Marine Micropaleontology*, **37**, 1–22.
- Flores, J.A., Sierro, F.J., Frances, G., Vazquez, A. and Zamarreno, I. 1997. The last 100.000 years in the western Mediterranean surface water and frontal dynamics as revealed by coccolithophores. *Marine Micropaleontology*, **29**, 351–366.
- Föllmi, K.B., Gertsch, B., Renevey, J.-P., De Kaenel, E. and Stilles, P. 2008. Stratigraphy and sedimentology of phosphate-rich sediments in Malta and south-eastern Sicily (latest Oligocene to early Late Miocene). *Sedimentology*, **55**, 1029–1051.
- Foresi, L.M., Donia, F., Mazzei, R. and Salvatorini, G. 2007. Revised age of the Maltese Lower Globigerina Limestone member (Globigerina Limestone formation): preliminary results. *Rivista Italiana di Paleontologia e Stratigrafia*, **46**, 175–181.
- Foresi, L.M., Iaccarino, S.M., Mazzei, R., Salvatorini, G. and Bambini, A.M. 2001. Il plancton calcareo (Foraminiferi e nannofossili) del Miocene delle Isole Tremiti. *Palaeontographia Italica*, **88**, 1–64.
- Foresi, L.M., Mazzei, R. and Salvatorini, G. 2002. Schema di biostratigrafia integrata a plancton calcareo per il Neogene-Quaternario. In: Bossio, A. et al., Note illustrative della carta geologica della zona di S. Maria di Lèuca. *Atti della Società Toscana di Scienze Naturali, Memorie, Serie A*, **107**, 145–158.
- Fornaciari, E., Raffi, I., Rio, D., Villa, G., Backman, J. and Olafsson, G. 1990. Quantitative distribution patterns of Oligocene and Miocene calcareous nannofossils from western equatorial Indian Ocean. In: Duncan, R.A., Backman, J., Peterson, L.C. et al. (Eds), *Proceedings of the Ocean Drilling Program, Scientific Results*, **115**, 237–254.
- Fornaciari, E. and Rio, D. 1996. Latest Oligocene to early middle Miocene quantitative calcareous nannofossil biostratigraphy in the Mediterranean region. *Micropaleontology*, **42** (1), 1–36.
- Geisen, M., Billard, C., Broerse, A.T.C., Cros, L., Probert, I. and Young, J.R. 2002. Life-cycle associations involving pairs of holococcolith species: intraspecific variation or cryptic speciation? *European Journal of Phycology*, **37** (4), 531–550.
- Geisen, M., Young, J.R., Probert, I., Saez, A.G., Baumann K.-H., Sprengel, C., Bollmann, J., Cros, L., De Vargas, C. and Medlin, L.K. 2004. Species level variation in coccolithophores. In: Thierstein, H.R. and Young, J. (Eds), *Coccolithophores – From Molecular Processes to Global Impact*, 327–366.
- Geitzenaver, K.R., Poche, M.B. and McIntyre, A. 1976. Modern Pacific coccolith assemblages: derivation and application to late Pleistocene paleotemperature analysis. In: Cline, R.M. and Hays, J.D. (Eds), *Investigation of late Quaternary paleoceanography and paleoclimatology. Geological Society of America*, 423–448.
- Giannelli, L. and Salvatorini, G. 1972. I Foraminiferi planktonici dei sedimenti terziari dell'Arcipelago maltese. I. Biostratigrafia del "Globigerina Limestone". *Atti della Società Toscana di Scienze Naturali, Memorie, Serie A*, **79**, 49–74.
- Giannelli, L. and Salvatorini, G. 1975. I Foraminiferi planktonici dei sedimenti terziari dell'Arcipelago maltese. II. Biostratigrafia di: "Blue Clay", "Greensand" e "Upper Coralline Limestone". *Atti della Società Toscana di Scienze Naturali, Memorie, Serie A*, **82**, 1–24.
- Gradstein, F.M., Ogg, J.G. and Smith, A.G. 2004a. *A Geological Time Scale 2004*, pp. 1–589. Cambridge University Press; U.K.
- Gradstein, F.M., Ogg, J.G., Smith, A.G., Bleeker, W. and Lourens, L.J. 2004b. A new Geological Time Scale with special reference to Precambrian and Neogene. *Episodes*, **2**, 83–100.
- Gregory, J.W. 1890–91. The maltese fossil Echinoidea and their evidence on the correlation of the Maltese rocks. *Transactions of the Royal Society of Edinburgh*, **36**, 225–231.
- Hagino, K. and Okada, H. 2006. Intra- and infra-specific morphological variation in selected coccolithophore species in the equatorial and subequatorial Pacific Ocean. *Marine Micropaleontology*, **58**, 184–206.
- Haq, B.U. 1980. Biogeographic History of Miocene calcareous nannoplankton and paleoceanography of the Atlantic Ocean. *Micropaleontology*, **26**, 414–443.
- Haq, B.U. and Lohmann, G.P. 1976. Early Cenozoic calcareous nannoplankton biogeography of the Atlantic Ocean. *Marine Micropaleontology*, **1**, 119–194.
- Hasle, G.R. 1960. A quantitative study of phytoplankton from equatorial Pacific. *Deep Sea Research*, **6**, 38–59.
- Honjo, S. 1990. Particle fluxes and modern sedimentation in the polar oceans. In: Smith Jr., W.O. (Ed.), *Polar oceanography. Part B: Chemistry, biology, and geology*, 687–739.
- House, M.R., Dunham, K.C. and Wigglesworth, J.C. 1961. Geology and structure of the Maltese Islands. In: Bowen,

- J.H., Dewdney, J.C. and Fisher, W.B. (Eds), Malta, a background for development. Durham University Press, 25–47.
- Hyde, T.G. 1955. Geology of the maltese islands. *Proceedings of the Royal Society of Edinburgh*, **B65**, 299.
- Iaccarino, S.M., Borsetti, A.M. and Rögl, F. 1996. Planktonic foraminifera of the Neogene Lemme – Carrosio GSSP Section (Piedmont, Northern Italy). In: Steininger, F.F., Iaccarino, S. and Cati, F. (Eds), In search of the Paleogene/Neogene boundary. Part 3: The Global Stratotype Section and Point. The GSSP for the base of the Neogene. *Giornale di Geologia*, **58**, 35–49.
- Iaccarino, S.M., Premoli Silva, I., Biolzi, M., Foresi, L.M., Lirer, F. and Petrizzo, M.R. 2005. Pratical manual of Oligocene to Middle Miocene Planktonic Foraminifera. In: Biolzi, M., Iaccarino, S.M. and Rettori, R. (Eds), International School on Planktonic Foraminifera, 4° Course, Perugia, 1–141.
- Jacobs, E., Weissert, H., Shield, G. and Stille, P. 1996. The Monterey event in the Mediterranean: A record from shelf sediments of Malta. *Paleoceanography*, **11**, 717–728.
- Kienel, U., Rehfeld, U. and Bellas, S.M. 1995. The Miocene Blue Clay formation of the Maltese Islands: Sequence-stratigraphic and paleoceanographic implications based on calcareous nannofossil stratigraphy. *Berliner Geowissenschaftliche Abhandlungen*, **16**, 533–557.
- Kleijne, A., Kroon, D. and Zevenboom, W. 1989. Phytoplankton and foraminiferal frequencies in northern Indian Ocean and Red Sea surface waters. *Netherlands Journal of Sea Research*, **24** (4), 531–539.
- Lourens, L., Hilgen, F., Shackleton, N.J., Laskar, J. and Wilson, D. 2004. The Neogene Period. In: Gradstein, F., Ogg, J. and Smith, A. (Eds), A Geological Timescale 2004. Cambridge University Press, 409–440.
- Marino, M., Maiorano, P. and Lirer, F. 2008. Changes in calcareous nannofossil assemblages during the Mid-Pleistocene Revolution. *Marine Micropaleontology*, **69**, 70–90.
- Martini, E. 1971. Standard Tertiary and Quaternary calcareous nannoplankton zonation. In: Farinacci, A. (Ed.), Proceedings of the second International Conference (Roma 1970) of Planktonic Microfossils 2, pp. 739–785. Roma.
- Martini, E. and Muller, C. 1986. Current Tertiary and Quaternary calcareous nannoplankton stratigraphy and correlations. *Newsletters on Stratigraphy*, **16**, 99–112.
- Mazzei, R. 1980. Studio biostratigrafico di alcune sezioni mioceniche dell'Arcipelago Maltese sulla base del Nannoplanton calcareo. *Paleontologia Stratigrafica e Evoluzione*, **1**, 149–152.
- Mazzei, R. 1986. The Miocene sequence of the Maltese Islands: biostratigraphic and chronostratigraphic references based on nannofossils. *Atti della Società Toscana di Scienze Naturali, Memorie*, **92**, 165–197.
- McIntyre, A. and Bè, A.W.H. 1967. Modern coccolithophorids of the Atlantic Ocean-I. Placoliths and cyrtoliths. *Deep Sea Research and Oceanographic Abstracts*, **14**, 561–564.
- McIntyre, A., Bè, A.W.H. and Roche, M.B. 1970. Modern Pacific Coccolithophorida: a paleontological thermometer. *New York Academy of Sciences Series II*, **32**, 720–731.
- Menesini, E. 1979a. Echinidi fossili dell'Arcipelago Maltese. *I. Atti della Società Toscana di Scienze Naturali, Memorie, Serie A*, **86**, 51–64.
- Menesini, E. 1979b. Maltese Fossil Echinoids. *Annales Géologiques des Pays Helléniques*, **2**, 799–806.
- Miller, K.G., Aubry, M.-P., Khan, M.J., Melillo, A.J., Kent, D.V. and Berggren, W.A. 1985. Oligocene-Miocene biostratigraphy, magnetostratigraphy and isotopic stratigraphy of the western North Atlantic. *Geology*, **13**, 257–261.
- Miller, K.G., Feigenson, M.D., Wright, J.D. and Clement, B.M. 1991a. Miocene isotope reference section Deep Sea Drilling Project Site 608: an evaluation of isotope and biostratigraphic resolution. *Paleoceanography*, **6**, 33–52.
- Miller, K.G. and Kent, D.V. 1987. Testing Cenozoic eustatic changes: the critical role of stratigraphic resolution. *Cushman Foundation for Foraminiferal Research Special Publication*, **24**, 51–56.
- Miller, K.G., Kominz, M.A., Browning, J.V., Wright, J.D., Mountain, G.S., Katz, M.E., Sugarman, P.J., Cramer, B.S., Christie-Blick, N. and Pekar, S.F. 2005. The Phanerozoic Record of Global Sea-Level Change. *Science*, **310**, 1293–1298.
- Miller, K.G., Wright, J.D. and Fairbanks, R.G. 1991. Unlocking the Ice House: Oligocene-Miocene Oxygen Isotopes, Eustasy, and Margin Erosion. *Journal of Geophysical Research*, **96**, 6829–6848.
- Miller, K.G., Wright, J.D. and Brower, A.N. 1989. Oligocene to Miocene stable isotope stratigraphy and planktonic foraminifer biostratigraphy of the Sierra Leone Rise (DSDP Site 366 and ODP Site 667). In: Ruddiman, W., Sarnthein, M. et al. (Eds), *Proceedings of the Ocean Drilling Program, Scientific Results*, **108**, 279–296.
- Muller, C. 1976. Tertiary and Quaternary calcareous nannoplankton from the Norwegian-Greenland Sea, DSDP Leg 38. In: Talwani, M., Volitsev, G. et al. (Eds), *Initial Reports of the Deep Sea Drilling Project*, **38**, 823–841.
- Murray, J. 1890. The Maltese Islands with special reference to their geological structure. *Scottish Geographical Magazine*, **6**, 449–488.
- Okada, H. and Bukry, D. 1980. Supplementary modification and introduction of code numbers to the low latitude coccolith biostratigraphy zonation (Bukry 1973, 1975). *Marine Micropaleontology*, **51**, 321–325.
- Okada, H. and Honjo, S. 1973. The distribution of oceanic coccolithophorids in the Pacific. *Deep Sea Research and Oceanographic Abstracts*, **20**, 355–364.
- Okada, H. and McIntyre, A. 1979. Seasonal distribution of

## CHRONOSTRATIGRAPHY OF THE MALTESE LOWER GLOBIGERINA LIMESTONE MEMBER

- modern coccolithophores in the western North Atlantic Ocean. *Marine Biology*, **54**, 319–328.
- Pearson, P.N. and Chaisson, W.P. 1997. Late Paleocene to Middle Miocene planktonic foraminifer biostratigraphy of the Ceara Rise. In: Shackleton, N.J., Curry, W.B., Richter, C. and Bralower, T.J. (Eds), *Proceedings of the Ocean Drilling Program, Scientific Results*, **154**, 33–68.
- Pedley, H.M. 1975. Oligocene-Miocene stratigraphy of the Maltese islands. Unpublished Ph.D. Dissertation, University of Hull; Hull.
- Pedley, H.M. 1976. A paleoecological study of the Upper coralline Limestone *Terebratula-Aphelesia* Bed (Miocene, Malta) based on bryozoan growth-forms and brachiopod distribution. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **20**, 209–234.
- Pedley, H.M. 1978a. A new lithostratigraphical and paleoenvironmental interpretation for the coralline limestone formations (Miocene) of the Maltese Islands. *Institute of Geological Sciences, Overseas Geology and Mineral Resources*, **54**, 273–291.
- Pedley, H.M. 1978b. A new record of fish bearing strata from the Maltese Islands and its palaeoenvironmental significance. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **24**, 73–83.
- Pedley, H.M. 1989. Controls of Cenozoic sedimentation in the Maltese Islands: review and interpretation. *Memorie della Società Geologica Italiana*, **38**, 81–94.
- Pedley, H.M. 1993. Geological Maps of the Maltese Islands. Scale: 1:25,000, 2 sheets. In: Oil Exploration Directorate, Office of the Prime Minister, Valletta, Malta, Geological map of the Maltese Islands (sheet 1, Malta; sheet 2 Gozo and Comino), British Geological Survey, Keyworth.
- Pedley, H.M. and Bennett, S.M. 1985. Phosphorites, hardgrounds and syndepositional subsidence: A paleoenvironmental model from Miocene of the Maltese Islands. *Sedimentary Geology*, **45**, 1–34.
- Pedley, H.M., Cugno, G. and Grasso, M. 1992. Gravity slide and resedimentation processes in a Miocene carbonate ramp, Hyblean Plateau, southeastern Sicily. *Sedimentary Geology*, **79**, 189–202.
- Pedley, H.M., House, M.R. and Waugh, B. 1976. The geology of Malta and Gozo. *Proceedings of the Geologists' Association*, **87**, 325–341.
- Pekar, S.F., Christie-Blick, N., Kominz, M.A. and Miller, K.G. 2002. Calibration between eustatic estimates from backstripping and oxygen isotopic records for the Oligocene. *Geology*, **30**, 903–906.
- Pekar, S.F., De Conto, R.M. and Harwood, D.M. 2006. Resolving a late Oligocene conundrum: Deep-Sea warming and Antarctic glaciation. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **231**, 29–40.
- Pekar, S.F. and Miller, K.G. 1996. New Jersey Oligocene “Icehouse” sequences (ODP leg 150X) correlated with the global  $\delta^{18}\text{O}$  and Exxon eustatic records. *Geology*, **24**, 567–570.
- Perch-Nielsen, K. 1985. Cenozoic calcareous nannofossils. In: Bolli, H.M., Saunders, J.B. and Perch-Nielsen, K. (Eds), *Plankton Stratigraphy*. Cambridge University Press, 427–554.
- Pratt, S.K. 1990. Hardgrounds genesis in pelagic carbonates from the Miocene of Malta and Cretaceous of southern England. Unpublished Ph.D. Dissertation, University of London, London.
- Raffi, I., Backman, J., Fornaciari, E., Palike, H., Rio, D., Lourens, L. and Hilgen, F. 2006. A review of calcareous nannofossils astrobiochronology encompassing the past 25 million years. *Quaternary Science Reviews*, **25**, 3113–3137.
- Rehfeld, U. and Janssen, A.W. 1995. Development of phosphatized hardgrounds in the Miocene Globigerina Limestones of the Maltese archipelago, including a description of *Gamopleura melitensis* sp. nov. (Gasteropoda, Euthecosomata). *Facies*, **33**, 91–106.
- Rizzo, C. 1932. Geology of the Maltese Islands. Government Printing Office, Malta.
- Roman, F. and Roger, J. 1939. Observations sur la faune des pectinides de Malta. *Bulletin de la Société Géologique de France*, **9**, 59–79.
- Rose, E.P.F. 1974. Stratigraphical and facies distribution of irregular echinoids in Miocene limestone of Gozo, Malta and Cyrenaica, Libia. 5th Congres du Neogene Mediteranneen, Mémoire du Bureau de Recherches Géologiques et Minières Tome 1, pp. 349–355. Lyon
- Rose, E.P.F., Pratt, S.K. and Bennett, S.M. 1992. Evidence for Sea-Level Changes in the Globigerina Limestone formation (Miocene) of the Maltese Islands. *Paleontologia i Evolució*, **24-25**, 265–276.
- Roth, P.H. and Coulbourn, W.T. 1982. Floral and solution patterns of coccoliths in surface sediments of the North Pacific. *Marine Micropaleontology*, **7**, 1–52.
- Roth, P.H. and Berger, W.H. 1974. Distribution and dissolution of coccoliths in the south and central Pacific. In: Sliter, W.V., Be, A.W.H. and Berger, W.H. (Eds), *Carbonate dissolution. Cushman Foundation for Foraminiferal Research Special Publication*, **13**, 87–113.
- Saez, A.G., Probert, I., Geisen, M., Quinn, P., Young, J.R. and Medlin, L.K. 2003. Pseudo-cryptic speciation in coccolithophores. *Proceedings of the National Academy of Sciences of United States of America*, **100**, 7163–7168.
- Sambleten, C. and Schroeder, A. 1992. Living coccolithophore communities in the Norwegian-Greenland Sea and their record in sediments. *Marine Micropaleontology*, **19**, 333–354.
- Sato, T., Yoguchi, S., Takayama, T. and Kameo, K. 2004. Drastic change in the geographical distribution of the cold-water nannofossil *Coccolithus pelagicus* (Wallich)

- Schiller at 2.74 Ma in the late Pliocene, with special reference to glaciation in the Arctic Ocean. *Marine Micropaleontology*, **52**, 181–193.
- Shackleton, N.J., Crowhurst, S.J., Weedon, G.P. and Laskar, J. 1999. Astronomical calibration of Oligocene – Miocene time. *Philosophical Transactions of the Royal Society of London, Series Mathematical, Physical and Engineering Sciences A*, **357**, 1907–1929.
- Spezzaferri S, Premoli Silva I (1991) Oligocene planktonic foraminiferal biostratigraphy and paleoclimatic interpretation from Hole 538A, DSDP Leg 77, Gulf of Mexico. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **83**, 217–263
- Spezzaferri, S. 1994. Planktonic foraminiferal biostratigraphy and taxonomy of the Oligocene and Lower Miocene in the oceanic record. An overview. *Palaeontographia Italica*, **81**, 1–187.
- Stainforth, R.M., Lamb, J.L., Luterbacher, H., Beard, J.H. and Jeffords, R.M. 1975. Cenozoic planktonic foraminiferal zonation and characteristics of index forms. *University of Kansas Paleontological Contributions*, **62**, 1–162.
- Steininger F.F., Aubry M.P., Berggren, W.A., Biolzi, M., Borsetti, A.M., Cartlidge, E.J.E., Cati, F., Corfield, R., Gelati, R., Iaccarino, S., Napoleone, C., Ottner, F., Rögl, F., Roetzel, R., Spezzaferri, S., Tateo, F., Villa, G. and Zevenboom, D., 1997. The Global Stratotype Section and Point (GSSP) for the base of the Neogene. *Episodes*, **20**, 23–28.
- Steininger, F.F., Iaccarino, S. and Cati, F. (Eds) 1996. In search of the Paleogene/Neogene boundary. Part 3 – The Global Stratotype Section and Point. The GSSP for the base of the Neogene. *Giornale di Geologia*, **58**, 1–192.
- Sugarman, P.J., McCartan, L., Miller, K.G., Feigenson, M.D., Pekar, S., Kistler, R.W. and Robinson, A.G. 1997. Strontium-isotopic correlation of Oligocene to Miocene sequences, New Jersey and Florida. In: Miller K.G. and Snyder S.W. (Eds), *Proceedings of the Ocean Drilling Program Scientific Results*, **150**, 147–159.
- Theodoridis, S. 1984. Calcareous nannofossil biozonation of the Miocene and revision of the helicoliths and *discoasters*. *Utrecht Micropaleontological Bulletins*, **32**, 1–271.
- Wells, P. and Okada, H. 1996. Holocene and Pleistocene glacial palaeoceanography off southeastern Australia, based on foraminifers and nannofossils in Vema cored hole V18-222. *Australian Journal of Earth Science*, **43**, 509–523.
- Winter, A., Jordan, R.W. and Roth, P.H. 1994. Biogeography of living coccolithophores in ocean waters. In: Winter A. and Siesser W.G. (Ed.), *Coccolithophores*. Cambridge University Press, pp. 161–177.
- Wright, T. 1855. On a new genus of fossil Cidaridae with a synopsis of the species included therein. *Annals and Magazine of Natural History*, **16**, 94–100. and *Proceedings of the Cotteswold Naturalists' Field Club*, **2**, 121.
- Zachos, J., Pagani, M., Sloan, L., Thomas, E. and Billups, K. 2001. Trends, Rhythms and Aberrations in Global Climate 65 Ma to Present. *Science*, **292**, 686–693.
- Zhang, J., Miller, K.G. and Berggren, W.A. 1993. Neogene planktonic foraminiferal biostratigraphy of the northeastern Gulf of Mexico. *Micropaleontology*, **39**, 299–326.
- Zhang, J. and Siesser, W.G. 1986. Calcareous nannoplankton in continental-shelf sediments, East China Sea. *Micropaleontology*, **32**, 271–281.

*Manuscript submitted: July 21<sup>th</sup> 2010*

*Revised version accepted: October 15<sup>th</sup> 2012*

## APPENDIX

GOZO ISLAND		THICKNESS OF LGNm	SAMPLES	REMARKS
SECTIONS	LOCATION			
<b>Reqqa Point</b> (Section 12 in Giannelli and Salvatorini 1972 and in Foresi <i>et al.</i> 2007)	Western side of Reqqa Point, along the northern coast of Gozo Island, 2 km WNW of Marsalforn	14 m	25 •	The succession composed of (from bottom to top): <b>(a)</b> upper part of the LCLf (including the "ScutellaBed") up to the TLCLHg; <b>(b)</b> the LGNm (with its phosphatized base) up to the TLGLHg; <b>(c)</b> the QCB, 30-50 cm thick. In its lower part (7 m thick), the LGNm is brown-yellow in colour, medium-size grained, very rich in bioturbation and with abundant <i>Scutella</i> in the first 1.50 m. Its upper half is light yellow in colour, friable, marly and with abundant <i>Flabellipecten</i> (concentrated in a medium level) and echinoids. The hardground at the top (less than 1 m thick) is hardly phosphatized
<b>Forna Point</b> (M19 and M32 sections of Foresi <i>et al.</i> 2007)	Along the northern coast of Gozo Island	14 m	36 ..	The succession composed of (from bottom to top): <b>(a)</b> the upper part of the LCLf up to the TLCLHg; <b>(b)</b> the LGNm (with its phosphatized base) up to the TLGLHg; <b>(c)</b> the QCB, 30-50 cm thick. The LGNm crops out with thickness and lithological features very similar to those of the Reqqa Point section
<b>Santa Lucija</b>	4 km NW of Santa Lucija, by the side of the road climbing from Wied Ilma up to in the Ghajn Abdul hill	14 m	8	The succession composed of (from bottom to top): <b>(a)</b> the upper part of the LGNm; the member is well diagenized, finely grained, yellowish in colour and characterized by bioturbations and bivalves. <b>(b)</b> A not well evident in outcrop TLGLHg; <b>(c)</b> the QCB, 20-40 cm thick.
<b>Dabroni</b> (Section 15 in Giannelli and Salvatorini 1972 and in Foresi <i>et al.</i> 2007)	Along the Marsalforn Valley, on the South-eastern flank of Dabroni hill	16 m	6	The succession composed of the LGNm. The unit starts above the phosphatized base for about 16 m; the "Scutella Bed" crops out in the lowermost part. At present, the LGNm sediments are widely covered with detritus and vegetation. Giannelli and Salvatorini (1972) report that the succession of the member is alike the foregoing sections
<b>Wardija Point</b>	North-western coast of Gozo Island, WNW of Xlendi Bay and just South of Wardija Point	4 m	5	The succession composed of the upper part of the LGNm up to the TLGLHg; the QCB crops out for 30-40 cm
<b>Xlendi</b> (Section 17 in Foresi <i>et al.</i> 2007)	South-western coast of Gozo Island, just South of Xlendi Bay	22 m	10 ...	The succession composed of (from bottom to top): <b>(a)</b> upper part (7-8 m above sea level) of the LCLf with abundant, large-size <i>Lepidocyclina</i> , common pectinids, and rare <i>Scutella</i> ; <b>(b)</b> the LGNm up to the TLGLHg; <b>(c)</b> the QCB, 30-50 cm thick. The LGNm shows a lowermost part (50 cm thick) composed of a yellow-ocre, coarse-grained limestone, rich in <i>Scutella</i> and bryozoans. Also present are others bivalves, echinoids, gasteropods and bioturbations. The 14 m of overlying hard sediments are medium-coarse grained, brown in colour, with abundant <i>Flabellipecten</i> particularly at the top. Above, for 7.5 m, the member is more friable and marly, yellowish, with rarer bioturbations and rich in echinoids and pectinids. At the top, the TLGLHg is 40-50 cm thick, with numerous <i>Thalassinoides</i> burrows, (filled with phosphatic elements in light matrix)
<b>Dahlet Qorrot</b>	Along the North-eastern coast of Gozo, just South of the homonymous locality	5m	4	The succession composed of the upper part of the LGNm up to the TLGLHg; the LGNm is well cemented, with medium-fine grained deposits, yellowish in colour, rich in <i>Flabellipecten</i> and echinoids in the upper part. The TLGLHg is 70 cm thick. The QCB, up to 60 cm thick, characterized by phosphatic elements up to 20-30 cm in diameter
<b>Ras-il-Hobz</b>	Just West of Ras-il-Hobz, on the southern coast of Gozo and East of Mgarr Ix-Xini	9 m	5	The succession composed of (from bottom to top): <b>(a)</b> the LGNm (with the exception of its lowermost deposits) with many bioturbations and, in the lower part, common <i>Scutella</i> (6 m thick). In the upper part, about 2 meters thick, <i>Flabellipecten</i> and echinoids occur up to the TLGLHg (70-80 cm thick); <b>(b)</b> the QCB is about 20-30 cm thick
.	14 samples for foraminifera. + 11 samples for calcareous nannofossils			
..	17 samples for foraminifera + 19 samples for calcareous nannofossils			
...	10 samples for foraminifera + 7 samples for calcareous nannofossils			

Table 1. Location, thickness of Lower Globigerina Limestone member, number of collected samples and main lithostratigraphic features of the studied sections on Gozo

MALTA ISLAND					
SECTIONS	LOCATION	THICKNESS OF LGLm	SAMPLES	REMARKS	
<b>Qammieh</b>	Along the north-western coast of Malta Island, at Rdum il-Qammieh	9.50 m	5	The succession composed of (from bottom to top): <b>(a)</b> the upper part (3.50 m thick) of LCLf (very rich in large <i>Lepidocyclina</i> ); 3 hardgrounds (10 cm, 15-30 cm and 20-30 cm thick respectively, with the highest one representing the TLCLHg, with common <i>Scutella</i> , <i>Flabellites</i> , other bivalves and bryozoans; <b>(b)</b> the LGLm (with phosphatized base) up to TLGLHg; <b>(c)</b> the QCB 40-50 cm thick. LGLm (6 m thick) with lower part characterized by many bio-turbations and abundant bryozoans, pectinids and common <i>Scutella</i> (in lowermost 3 m of the unit), and an upper part with abundant pectinids and echinoids. Upwards (at 2/3 from the base) the member becomes light yellow, poorly bioturbated and more friable. Higher up there is a hardground, with phosphatic elements and pectinids. Other discontinuity surfaces are present at the top of TLGLHg	
<b>Fomm-ir-Rih</b> (Section M14 in Foresi et al. 2007)	Most internal part of Fomm-ir-Rih Bay, almost at the mouth of Wied Gerzuma, just near the "Victoria lines Fault"	4.50 m	16	The succession composed of (from bottom to top): <b>(a)</b> the uppermost part of LCLf (very rich in foraminifera, bryozoans and echinoids) up to the TLCLHg (with planar phosphatized upper surface); <b>(b)</b> the LGLm up to TLGLHg; <b>(c)</b> QCB 20-35 cm thick. LGLm characterized by bioturbations and abundant bryozoans, pectinids and echinoids	
<b>Il-Blata</b> (Section 6 in Giannelli and Salvatorini 1972 and in Foresi et al. 2007)	Along the western coast of Malta Island, about 1 Km SW of Fomm-ir-Rih Section	3.40 m	8 ..	The succession composed of (from bottom to top): <b>(a)</b> the uppermost part of LCLf up to the TLCLHg; <b>(b)</b> the LGLm (with phosphatized base) up to the TLGLHg; <b>(c)</b> the QCB 20-100 cm thick. The LGLm characterized by abundant bryozoans and numerous bioturbations, echinoids and pectinids; the latter concentrated in the lower and upper part. <i>Scutella</i> is abundant in the lowermost part	
<b>Mthaleb</b> (Section 4 in Giannelli and Salvatorini 1972 and in Foresi et al. 2007)	Mthaleb area, just North of Il-Qaws (NW of Dingli, Malta Island)	1.30 m	1 ...	The succession composed of (from bottom to top): <b>(a)</b> the LGLm up to TLGLHg; <b>(b)</b> the QCB 30-40 cm thick. LGLm characterized by numerous bioturbations (upper part), bryozoans, pectinids, and echinoids. Rare <i>Scutella</i> occur in the basal coarse-grained yellowish biotitic limestone (30 cm thick)	
<b>Dingli</b> (Section 3 in Giannelli and Salvatorini 1972 and in Foresi et al. 2007)	South-western part of Malta Island, at the cliffs of Dingli, just 1Km South of the homonymous village	1.40 m	5	The succession composed of (from bottom to top): <b>(a)</b> the LGLm up to TLGLHg; <b>(b)</b> the QCB which is about 15-25 cm thick. The LGLm is represented by whitish marly limestones (60-100 cm thick); the macrofossils are rare	
<b>Wied-id-Dis</b>	Eastern part of Malta Island, along the way up to Ghargħur, about 4Km NW of Sliema	40 m	14	The succession composed of LGLm (yellowish limestones and light yellow marly limestones) characterized by rare pectinids	
<b>Sliema Point-Qui Si Sana</b> (Sliema Section in Foresi et al. 2007)	Composite section: lower part c. 200 m NW of the Sliema Point Tower; upper part at Qui-Si-Sana, c. 500 m SSE of the Sliema Point Tower	7 m	8	The succession composed of (from bottom to top): <b>(a)</b> the LGLm up to the TLGLHg (the base of the unit crops out along the coast of Balluta Bay, there, at the top of LCLf, there is a hardground, with seaweeds and <i>Scutella</i> ; <b>(b)</b> the QCB 15-30 cm thick. The LGLm characterized by common pectinids ( <i>Flabellites</i> ), echinoids, and numerous bioturbations	
<b>Fort St Elmo</b> (St.Elmo Section in Foresi et al. 2007)	North side of the Grand Harbour (Valletta), along the coast next to Fort St. Elmo	15 m	6	The succession composed of the lower part of the LGLm, which is especially rich in echinoids and bioturbations	
<b>Valletta</b>	Along the Triq L-Assedju L-Kbir, that borders the Marsamxett Harbour, and SSW of St. Michael's Bastion	33 m	6 :	The succession composed of the middle-upper part of LGLm (yellowish, fine calcarenous with scattered echinoids) up to TLGLHg; the QCB (5-15 cm thick) crops out along the Triq Marsamxett, next to St. Paul's Cathedral)	

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<b>MALTA ISLAND</b>		<b>THICKNESS OF LGLm</b>	<b>SAMPLES</b>	<b>REMARKS</b>
<b>SECTIONS</b>	<b>LOCATION</b>			
<b>Kalkara</b>	Along the rock walls around the Grand Harbour below Kalkara (lower part) and to the side of the road facing the Kalkara Creek (upper part)	8 m	3	The succession composed of a portion of the LGLm, developed as fine-grained, marly and bioturbated yellowish sediment, with abundant echinoids
<b>Huttaf Gandolf</b>	Next to Huttaf Gandolf locality (South of Luqa Airport), in a quarry on the side of the road Luqa-Mqabba	22 m	18 ::	The succession composed of the LGLm (yellowish fine-grained limestones). Macrofossils are rare
<b>Marsaskala</b> (Marsaskala Section in Foresi <i>et al.</i> 2007)	North coast of Marsaskala Bay (South eastern part of Malta Island)	13 m	8	The succession composed of (from bottom to top): (a) uppermost part of the LCLf (with <i>Lepidocyclina</i> and <i>Scutella</i> ) up to the TLCLHg; (b) the LGLm up to the TLGLHg; (c) the QCB, 20-35 cm thick (along the coastal road and in Zonqor Point). The LGLm characterized by common bioturbations, bryozoans, pectinids and echinoids in the lower part; rare pectinids in the upper part
.	7 samples for foraminifera + 9 samples for calcareous nannofossils			
..	3 samples for foraminifera + 5 samples for calcareous nannofossils			
...	1 sample for foraminifera			
:	6 samples for foraminifera and 5 samples for calcareous nannofossils			
::	18 samples for foraminifera and 8 samples for calcareous nannofossils			

Table 2. Location, thickness, member, number of collected samples and main lithostratigraphic features of the studied sections on Malta

Malta Island	<i>Cassigerinella chipolensis</i>	<i>Dentoglobigerina globularis</i>	<i>Globigerina ciporenensis</i>	<i>Globigerina praebulloides</i> s.l.	<i>"Globigerina" euapertura</i>	<i>"Globigerina" sellii</i>	<i>"Globigerina" tripartita</i>	<i>Globigerinoides quad: primordius</i>	<i>Globoturborotalita anguliofficinalis</i>	<i>Globoturborotalita angulisuturalis</i>	<i>Globoturborotalita brasieri</i>	<i>Globoturborotalita woodi</i>	<i>Paragloborotalia opima nana</i>	<i>Paragloborotalia pseudokugleri</i>	<i>Subtina gortani</i>	<i>Tenuitellinata angustumobilis</i>	<i>Tenuitellinata uvula</i>	
<b>Qammieh Section</b>																		
MH5	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
MH6	R	/	R	R	/	/	/	/	R	R	/	/	R	R	/	C	/	
MH7	R	/	/	C	/	/	/	R	/	R	/	/	R	R	/	C	/	
MH8	R	/	/	C	/	/	/	/	C	C	/	/	R	R	/	A	/	
MH9	/	R	/	C	/	/	/	R	C	/	/	/	R	R	/	C	/	
<b>Fomm Ir-Rih Section</b>																		
FIR5	/	/	/	R	/	/	/	/	/	/	/	/	/	/	/	R	/	
A	/	/	/	/	/	/	/	R	/	/	/	/	R	/	/	R	/	
FIR6	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
FIR7	R	/	/	C	/	/	/	/	/	/	/	/	/	R	/	R	/	
FIR8	/	/	/	R	/	/	/	/	/	/	/	/	/	/	/	/	/	
B	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
FIR9	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
<b>II Blata Section</b>																		
709	R	R	R	R	R	/	/	R	R	R	/	/	R	R	/	R	/	
MT10	/	/	R	/	/	/	/	/	/	/	/	/	R	/	/	R	/	
0	/	/	/	R	/	/	/	/	/	/	/	/	/	/	/	R	/	
<b>Sliema Point-Qui Si Sana Section</b>																		
3208	R	C	R	A	R	/	/	R	R	C	R	C	C	A	A	/	R	A
3209	R	A	C	A	C	C	/	R	C	A	C	R	C	A	C	R	A	/
3210	R	A	R	A	C	C	R	/	C	R	R	A	C	C	A	/	R	A
3211	R	C	R	A	R	C	/	R	R	R	/	R	C	C	/	/	A	/
3212	R	A	C	A	C	/	R	R	R	R	/	R	A	R	R	C	/	
3213	C	C	/	/	C	/	/	R	C	R	/	/	C	R	/	R	A	/
QS1	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
3214	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
<b>Fort St. Elmo Section</b>																		
3165	R	C	/	R	C	/	/	/	R	/	/	/	C	C	C	/	R	/
<b>Kalkara Section</b>																		
VT1	R	C	/	A	C	R	R	R	R	C	/	/	A	A	C	C	A	/
VT2	R	C	R	R	R	R	R	/	R	R	/	/	C	C	R	C	A	/
VT3	R	R	/	R	R	R	/	/	R	R	/	/	R	R	/	R	R	/
<b>Marsaskala Section</b>																		
MK4	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
MK5	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
3194	C	C	R	A	R	/	/	R	R	R	/	R	C	C	/	R	A	/
3195	C	C	C	A	R	/	/	C	C	R	/	R	C	R	/	/	A	/
3196	C	C	C	A	R	R	R	R	C	R	/	/	A	C	/	R	A	/
3197	C	A	R	A	R	/	/	R	R	R	/	/	A	C	/	R	A	/
<b>Wied id-Dis Section</b>																		
WD1	A	R	/	A	/	R	/	/	R	R	R	/	C	R	C	R	A	/
WD2	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
WD3	C	R	/	C	C	R	/	/	R	R	/	/	R	/	C	/	C	/
WD4	C	A	R	A	C	R	/	C	R	A	R	/	C	C	C	R	A	R
WD5	R	R	/	R	R	/	/	/	R	/	/	/	C	C	R	R	R	/
WD6	R	/	/	R	/	/	/	/	/	/	/	/	R	R	R	/	C	/
WD7	R	/	/	R	/	/	/	/	/	/	/	/	R	R	R	/	R	/
WD8	R	/	/	R	R	/	/	/	R	/	/	/	R	R	R	/	R	/
WD9	C	/	/	R	/	/	/	R	/	/	/	/	R	R	R	/	C	/
WD10	/	/	/	/	/	/	/	R	/	/	/	/	/	/	/	/	/	/
WD11	R	/	/	C	/	/	/	R	/	/	/	/	R	R	R	/	C	/
WD12	R	R	/	R	/	/	/	/	/	/	/	/	R	R	R	/	R	/
WD13	/	C	/	/	/	/	/	/	/	/	/	/	/	/	/	R	/	/
WD14	R	/	/	C	R	R	/	/	/	/	/	/	R	R	/	R	/	/

## CHRONOSTRATIGRAPHY OF THE MALTESE LOWER GLOBIGERINA LIMESTONE MEMBER

Malta Island	<i>Cassigerinella chipolensis</i>	<i>Dentoglobigerina globularis</i>	<i>Globigerina ciporenensis</i>	<i>Globigerina praebullioides</i> s.l.	<i>"Globigerina" euapertura</i>	<i>"Globigerina" sellii</i>	<i>"Globigerina" tripartita</i>	<i>Globigerinoides quad: primordius</i>	<i>Globoturborotalita angulifascinatis</i>	<i>Globoturborotalita angulusuturalis</i>	<i>Globoturborotalita brazieri</i>	<i>Globoturborotalita woodi</i>	<i>Paragloborotalia opima nana</i>	<i>Paragloborotalia pseudokugleri</i>	<i>Subtina gortani</i>	<i>Tenuitellinata angustumobilis</i>	<i>Tenuitellinata uvula</i>
<b>Dingli Section</b>																	
3203	R	A	R	A	C	R	R	R	/	/	C	R	R	C	C	/	
3204	/	R	R	R	R	/	/	R	/	/	R	/	/	R	R	/	
3205	R	C	R	A	R	R	R	R	C	/	/	A	/	C	R	/	
3206	R	A	R	A	C	/	R	R	C	C	/	R	A	C	R	/	
3207	/	A	R	A	A	/	/	R	R	C	/	C	A	C	C	/	
<b>Huttaf Gandolf Section</b>																	
MQ1	R	R	/	C	R	/	/	/	R	/	/	C	/	C	/	C	/
3180	R	A	R	A	C	R	R	R	C	R	/	C	R	R	R	A	R
MQ2	/	R	/	C	R	/	/	/	R	/	/	R	R	R	/	C	/
3181	R	C	R	A	R	/	R	R	/	R	/	C	C	R	R	A	/
3182	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
MQ3	/	/	/	C	/	/	/	/	R	/	/	R	R	/	/	R	/
3183	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
MQ4	R	/	/	R	/	/	/	R	/	/	/	R	/	/	/	R	/
3184	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
3185	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
MQ5	/	/	/	R	/	/	/	/	R	/	/	R	R	/	/	R	/
3186	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
MQ6	R	R	/	R	/	/	/	/	/	/	/	R	/	/	/	R	/
3187	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
3188	R	R	/	C	/	/	/	/	R	/	/	R	R	/	/	R	/
MQ7	R	R	/	R	/	/	/	/	R	/	/	R	/	/	/	R	/
3189	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
MQ8	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/
<b>Mtahleb Section</b>																	
643	/	C	R	A	R	/	/	R	R	R	/	/	R	R	/	C	A

Table 3. Semiquantitative foraminiferal content of the analyzed sections of Lower Globigerina Limestone member on Malta

<b>Gozo Island</b>		<i>Cassigerinella chipolensis</i>	<i>Dentoglobigerina globularis</i>	<i>Globigerina ciperoensis</i>	<i>Globigerina praebulloides</i> s.l.	<i>"Globigerina" euapertura</i>	<i>"Globigerina" sellii</i>	<i>"Globigerina" tripartita</i>	<i>Globigerinella obesa</i>	<i>Globigerinoides quad: primordius</i>	<i>Globoturborotalita anguliofficinalis</i>	<i>Globoturborotalita angulusaturalis</i>	<i>Globoturborotalita brazieri</i>	<i>Globoturborotalita woodi</i>	<i>Paragloborotalita opima nana</i>	<i>Paragloborotalita pseudokugleri</i>	<i>Paragloborotalita siakensis</i>	<i>Subbotina gortani</i>	<i>Tenuitellinata angustumibilis</i>	<i>Tenuitellinata uvula</i>
<b>Reqqa Point Section</b>																				
12C	/	/	/	R	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
12D	R	/	/	R	/	/	/	/	/	R	/	/	/	R	/	/	/	R	/	
12E	R	/	/	R	C	/	/	/	/	R	/	/	/	R	/	/	/	R	/	
788	R	R	R	C	/	/	/	R	R	R	/	/	R	R	R	/	R	A	/	
2408	/	R	/	R	/	/	/	/	/	R	/	/	/	R	/	/	R	/	/	
2409	C	R	R	A	/	/	/	R	/	C	R	/	R	R	C	C	/	A	/	
789	C	A	R	A	R	R	/	R	R	C	R	/	R	C	C	C	R	A	/	
2411	R	R	/	C	/	/	/	/	/	R	/	/	/	R	C	R	/	R	/	
2412	/	R	R	R	C	/	/	/	R	R	R	/	R	A	A	/	A	/	/	
790	R	A	R	A	R	R	/	/	R	R	R	/	R	A	A	/	R	A	/	
2413	R	R	/	C	R	/	/	R	R	R	/	R	R	R	C	R	R	/	/	
2414	/	C	R	C	/	R	/	/	/	/	/	/	/	C	C	R	C	/	/	
<b>Forna Point Section</b>																				
3216	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
3217	R	R	/	A	/	/	/	/	R	R	/	/	R	/	C	/	C	/	/	
3218	/	R	R	C	/	/	/	/	/	R	/	/	/	R	/	R	/	R	/	
3219	R	C	R	C	R	R	/	/	R	R	R	/	R	C	C	R	C	/	/	
3220	R	R	/	C	R	/	/	/	/	C	R	/	/	R	R	R	R	/	/	
3221	/	R	R	R	C	R	/	/	/	C	R	/	/	R	R	R	R	C	/	
3222	C	C	R	C	R	/	/	/	/	C	R	/	/	R	R	R	R	/	/	
3223	R	C	R	C	/	/	/	/	/	C	R	/	/	R	R	C	/	R	/	
3224	C	C	R	A	R	/	/	R	R	C	R	/	/	C	R	C	C	R	/	
3199	C	C	R	C	R	/	/	R	R	C	R	/	/	R	C	C	/	C	/	
3225	R	R	/	R	R	/	/	/	R	R	R	/	/	R	R	R	/	R	/	
3226	R	R	/	R	R	/	/	R	R	R	/	/	/	R	/	R	/	R	/	
3200	R	A	R	A	C	C	/	R	/	R	/	/	/	C	C	A	R	C	/	
3227	/	R	/	C	R	R	/	/	R	/	/	/	R	R	/	C	/	A	/	
3201	R	C	R	A	R	/	/	R	/	R	/	/	R	C	R	R	C	/	/	
3228	/	A	R	A	R	/	/	/	R	R	R	/	/	C	C	/	R	A	/	
3229	/	C	R	C	R	R	/	/	R	R	R	/	/	C	C	/	R	A	/	
<b>Dahlet Qorrot Section</b>																				
DQ1	R	/	/	R	/	/	/	R	R	R	/	/	R	R	R	/	R	/	/	
DQ2	R	/	/	R	/	R	/	R	/	C	R	/	/	R	R	R	/	R	/	
DQ3	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
DQ4	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
<b>Wardija Point Section</b>																				
PT1	C	C	R	A	C	/	R	R	R	R	R	/	R	C	/	A	R	A	/	
PT2	R	R	R	R	R	R	/	R	C	/	/	R	C	R	C	R	A	A	/	
PT3	R	R	R	A	R	/	/	R	R	R	R	/	R	C	R	C	R	A	/	
PT4	R	A	/	A	C	/	R	/	R	R	R	/	R	C	R	C	/	R	/	
PT5	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
<b>Xlendi Section</b>																				
MT10	R	/	/	C	/	/	/	/	/	/	/	/	R	R	R	/	C	/	/	
MT11	R	R	R	A	R	/	/	/	C	/	/	/	R	C	R	/	C	/	/	
MT12	R	R	/	C	R	/	/	/	/	/	/	/	R	R	R	/	A	/	/	
MT12bis R	R	C	R	A	/	C	R	/	/	R	R	/	R	A	R	C	/	C	/	
MT12ter /	R	R	R	C	/	/	R	/	/	R	R	/	R	R	C	R	/	C	/	
3191	R	A	R	A	C	R	/	R	R	R	R	/	R	R	C	C	R	A	/	
MT13	C	C	C	R	R	R	/	R	R	C	R	/	R	A	R	R	C	C	/	
MT14	/	R	R	R	A	R	R	/	R	R	R	/	R	A	R	R	R	A	/	
3192	/	A	R	A	C	R	/	R	C	R	R	/	R	A	C	R	R	A	/	
3193	R	/	R	C	R	/	/	R	/	R	R	/	R	R	C	C	R	A	/	

## CHRONOSTRATIGRAPHY OF THE MALTESE LOWER GLOBIGERINA LIMESTONE MEMBER

Gozo Island	<i>Cassigerinella chipolensis</i>	<i>Dentoglobigerina globularis</i>	<i>Globigerina ciporenensis</i>	<i>Globigerina praebulloides</i> s.l.	<i>"Globigerina" euapertura</i>	<i>"Globigerina" sellii</i>	<i>"Globigerina" tripartita</i>	<i>Globigerinella obesa</i>	<i>Globigerinoides quad: primordius</i>	<i>Globoturborotalita angulif officinalis</i>	<i>Globoturborotalita angulisuturalis</i>	<i>Globoturborotalita brazieri</i>	<i>Globoturborotalita woodi</i>	<i>Paragloborotalia opima nana</i>	<i>Paragloborotalia pseudokugleri</i>	<i>Paragloborotalia siakensis</i>	<i>Subbotina gortani</i>	<i>Tenuitellinata angustumobilis</i>	<i>Tenuitellinata uvula</i>
<b>Dabrani Section</b>																			
3174	R	/	/	R	/	/	/	/	/	/	/	/	/	R	R	/	R	/	
3175	/	/	/	R	/	/	/	/	/	R	/	/	/	R	R	/	R	/	
3176	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
3177	/	R	R	C	R	/	/	R	/	C	R	/	/	C	C	R	/	A	
3178	R	R	/	C	R	R	/	/	R	R	/	R	C	C	R	/	A	/	
3179	/	C	R	A	C	R	/	R	R	/	R	/	C	C	C	R	A	/	
<b>Ras Il-Hobz Section</b>																			
HO1	R	R	/	C	/	/	/	/	R	R	/	/	R	R	R	/	R	R	
HO2	R	R	/	A	R	/	/	R	/	C	R	/	/	C	C	/	C	/	
HO3	R	R	/	R	/	/	/	/	R	/	/	/	R	R	/	/	R	/	
HO4	R	C	/	A	R	R	/	R	R	C	/	/	C	C	/	R	A	/	
HO5	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
<b>Santa Lucija Section</b>																			
SL1	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	
SL2	R	R	/	R	/	/	/	/	R	/	/	/	R	R	/	R	/		
SL3	R	R	/	R	/	/	/	/	R	/	/	/	R	/	/	R	/		
SL4	/	C	R	A	C	C	/	/	R	C	R	/	/	R	A	R	R	C	
SL5	A	C	R	A	R	/	/	R	R	/	/	/	R	R	R	R	R	/	
SL6	R	C	/	C	C	/	R	/	R	R	/	/	C	R	R	R	A	/	
SL7	R	C	R	A	C	/	/	R	R	R	/	/	C	R	C	R	A	/	
SL8	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	/	

Table 4. Semiquantitative foraminiferal content of the analyzed sections of Lower Globigerina Limestone member on Gozo

Malta Island	<i>Braarudosphaera bigelowii</i>	<i>Coccolithus miopelagicus</i>	<i>Coccolithus pelagicus</i>	<i>Cyclargolithus abisectus</i>	<i>Cyclargolithus floridanus</i>	<i>Dicyoecites bisectus</i>	<i>Dicyoecites scriptae</i>	<i>Discoaster deflatarei</i>	<i>Helicosphaera euphratis</i>	<i>Helicosphaera obliqua</i>	<i>Helicosphaera intermedia</i>	<i>Helicosphaera recta</i>	<i>Pontosphaera enormis</i>	<i>Pontosphaera multipora</i>	<i>Sphenolithus ciperensis</i>	<i>Sphenolithus compactus</i>	<i>Sphenolithus delphix</i>	<i>Sphenolithus dissimilis</i>	<i>Sphenolithus moriformis</i>	<i>Umbilicosphaera sibogae</i>	<i>Zygrolithus bijugatus</i>
<b>Qammieh Section</b>																					
MH5	/	/	C	R	A	R	C	/	/	/	/	/	/	C	/	/	/	/	A	A	
MH6	/	R	C	R	A	R	A	/	/	/	/	/	/	C	R	/	/	A	R	A	
MH7	/	/	/	/	A	R	A	/	/	/	R	/	C	/	/	/	R	R	A	A	
MH8	/	/	C	C	A	C	A	/	R	/	/	/	C	/	/	/	C	R	C	A	
MH9	/	/	R	C	A	C	A	/	R	/	/	R	R	R	/	/	R	R	A	C	
<b>Fomm ir-Rih Section</b>																					
FIR1	/	/	C	/	A	C	A	/	/	/	/	/	/	R	/	/	/	C	R	A	
FIR2	/	/	A	/	A	C	A	/	/	/	/	/	/	C	/	/	/	R	R	A	
FIR3	/	/	R	R	A	R	A	/	/	/	/	/	/	C	/	/	/	R	R	A	
FIR4	/	/	C	R	A	/	A	/	R	R	/	/	/	R	/	/	/	C	C	A	
FIR5	/	/	C	R	A	R	A	/	R	/	/	/	C	R	/	/	R	R	A	A	
FIR6	/	/	A	R	A	R	A	/	R	R	/	R	/	C	/	/	C	R	A	A	
FIR7	/	/	C	R	A	C	A	/	/	R	/	R	/	R	/	/	C	R	A	A	
FIR8	/	/	R	/	A	/	A	/	/	/	/	/	C	/	/	/	R	/	A	A	
FIR9	/	/	R	/	A	/	A	/	/	/	/	/	C	/	/	/	R	/	A	A	
<b>II Blata Section</b>																					
BL1	/	R	C	C	A	/	A	/	/	/	/	/	/	R	/	/	C	/	C	A	
BL2	/	R	C	C	A	R	A	/	/	/	/	/	/	R	/	/	C	R	C	C	
BL3	/	/	C	C	A	R	A	/	/	/	/	/	/	R	/	/	R	C	R	C	
BL4	/	R	C	C	A	R	A	/	/	/	/	/	/	R	/	/	C	/	A	A	
BL5	/	/	R	C	A	R	A	/	/	/	/	/	/	R	/	/	R	C	R	A	
<b>Sliema Point-Qui Si Sana Section</b>																					
3208	/	/	C	R	A	R	A	/	C	/	/	R	R	R	/	R	/	C	R	A	
3209	/	R	R	R	A	R	A	R	C	/	R	/	R	R	/	/	C	R	A	R	
3210	/	/	R	R	A	R	C	/	R	/	/	R	R	R	/	/	C	R	C	R	
3211	R	R	C	R	A	R	A	/	C	/	/	R	R	R	/	/	C	C	A	R	
3212	/	R	A	C	C	R	C	R	C	/	R	R	/	C	R	/	/	C	R	A	
3213	/	/	R	R	C	R	A	/	C	R	/	/	R	/	/	/	A	R	A	C	
QS1	/	/	R	/	R	R	C	/	R	/	/	/	/	/	/	/	R	/	C	R	
3214	/	/	C	R	A	R	C	/	R	/	/	R	/	/	/	/	C	R	A	C	
<b>Fort St. Elmo Section</b>																					
MT2	/	/	/	/	/	R	/	/	/	/	/	/	/	R	/	/	R	/	R	R	
MT3	/	/	R	/	R	/	/	/	/	R	/	/	R	/	/	/	R	/	R	R	
<b>Kalkara Section</b>																					
VT1	/	C	C	R	A	R	A	/	R	/	R	R	/	R	R	/	R	C	C	R	
VT2	/	R	R	/	A	/	A	/	/	/	/	/	/	R	/	/	R	R	C	R	
VT3	/	/	C	R	A	R	A	/	R	/	/	/	R	/	/	/	R	R	C	R	
<b>Wied id-Dis Section</b>																					
WD1	/	/	A	R	A	C	A	/	C	/	/	/	/	C	R	/	/	C	R	C	
WD2	/	/	R	R	A	C	A	/	R	/	/	/	/	C	/	/	R	R	A	C	
WD3	/	/	R	R	A	C	A	/	R	/	/	/	/	C	/	/	R	R	A	C	
WD4	/	/	R	R	A	C	A	/	R	/	/	/	/	R	/	/	R	R	A	R	
WD5	/	/	R	R	A	C	C	/	C	/	/	/	/	C	R	/	/	R	/	A	
WD6	/	/	R	R	A	C	A	/	R	/	/	/	/	R	R	/	/	C	/	A	
WD7	/	/	C	R	A	C	A	/	C	/	/	/	/	C	R	/	/	C	/	A	
WD8	/	/	C	/	A	C	A	/	C	/	/	/	/	R	R	/	/	C	R	A	
WD9	/	/	R	/	A	C	A	/	R	/	/	/	/	C	R	/	/	R	R	A	
WD10	/	/	/	/	A	R	A	/	R	/	/	/	/	R	/	/	C	R	A	A	
WD11	/	/	C	/	A	C	A	/	R	/	/	/	/	R	R	/	/	R	/	C	
WD12	/	/	R	/	A	C	A	/	R	/	/	/	/	R	/	/	R	/	A	C	
WD13	/	/	C	/	A	C	A	/	R	/	/	/	/	C	/	/	/	/	A	A	
WD14	/	/	R	/	A	/	C	/	/	/	/	/	/	R	/	/	/	/	A	A	

<b>Malta Island</b>	<i>Braarudosphaera bigelowii</i>	<i>Coccolithus miopelagicus</i>	<i>Coccolithus pelagicus</i>	<i>Cyclicargolithus abisectus</i>	<i>Cyclicargolithus floridanus</i>	<i>Dictyococcites bisectus</i>	<i>Dictyococcites scrippae</i>	<i>Discoaster deflandrei</i>	<i>Helicosphaera euphratis</i>	<i>Helicosphaera intermedia</i>	<i>Helicosphaera obliqua</i>	<i>Helicosphaera recta</i>	<i>Pontosphaera enomis</i>	<i>Pontosphaera multipora</i>	<i>Sphenolithus ciperensis</i>	<i>Sphenolithus compactus</i>	<i>Sphenolithus delphix</i>	<i>Sphenolithus dissimilis</i>	<i>Sphenolithus moriformis</i>	<i>Umbilicosphaera sibogae</i>	<i>Zygralithus bijugatus</i>
<b>Huttaf Gandolf Section</b>																					
MQ1	/	/	C	C	A	A	A	/	R	R	/	/	/	R	R	/	/	C	R	C	R
MQ2	/	/	C	/	A	C	A	/	R	/	/	/	/	/	/	/	/	C	/	C	C
MQ3	/	/	C	/	A	C	A	/	/	/	/	/	/	/	/	/	/	C	/	R	A
MQ4	/	/	C	R	A	R	A	/	R	R	/	/	/	/	/	/	/	C	/	A	A
MQ5	/	/	/	/	A	A	A	/	R	/	/	/	/	/	/	/	/	/	/	A	A
MQ6	/	/	R	R	A	C	A	/	/	R	/	/	R	C	R	/	/	C	/	A	A
MQ7	/	/	C	R	A	C	A	/	/	R	/	/	R	C	R	/	/	C	/	A	A
MQ8	/	/	R	/	C	/	R	/	/	/	/	/	R	A	R	/	R	R	/	A	A
<b>Marsaskala Section</b>																					
MK3	/	/	R	R	C	R	A	/	R	/	/	R	/	R	R	/	/	C	R	C	R
MK4	/	/	C	C	A	C	A	/	/	/	/	/	/	/	/	/	/	C	/	A	C
MK5	/	R	C	R	A	C	C	/	R	/	/	/	R	R	/	/	/	R	/	A	R
MK7	/	R	R	R	C	R	C	/	R	/	/	/	/	/	/	/	/	R	R	C	R
<b>Valletta Section</b>																					
VT4	/	/	/	/	R	/	R	/	/	/	/	/	/	C	/	/	/	C	R	A	A
VT5	/	/	C	/	R	/	R	/	/	/	/	/	/	C	/	/	/	C	/	A	A
VT6	/	/	R	/	C	/	R	/	/	/	/	/	/	R	/	/	/	C	/	A	A
VT7	/	/	R	/	C	/	C	/	/	/	/	/	/	C	/	/	/	C	R	A	A
VT8	/	/	R	/	R	/	R	/	/	/	/	/	/	C	/	/	/	C	R	A	A

Table 5. Semiquantitative calcareous nannofossil content of the analyzed sections of Lower Globigerina Limestone member on Malta

Gozo Island	<i>Coccolithus miopelagicus</i>	<i>Coccolithus pelagicus</i>	<i>Cyclargolithus abisectus</i>	<i>Cyclargolithus floridanus</i>	<i>Dictyococcites biscutus</i>	<i>Dictyococcites scrippae</i>	<i>Helicosphaera euphratis</i>	<i>Helicosphaera obliqua</i>	<i>Pontosphaera multipora</i>	<i>Sphenolithus ciperensis</i>	<i>Sphenolithus conicus</i>	<i>Sphenolithus dissimilis</i>	<i>Sphenolithus moriformis</i>	<i>Sphenolithus pseudoradians</i>	<i>Triquetrorhabdulus carinatus</i>	<i>Umbilicosphaera sibogae</i>	<i>Zygrolithus bijugatus</i>
<b>Reqqa Point Section</b>																	
RQ6	/	C	R	A	C	A	R	/	C	/	/	/	/	/	/	A	A
RQ7	/	R	R	A	C	A	A	/	C	/	/	R	/	/	/	A	A
RQ8	/	C	C	A	C	A	R	/	C	/	/	C	R	/	/	A	A
RQ9	/	C	C	A	C	A	R	/	C	/	/	C	R	/	/	A	C
RQ10	R	C	R	A	C	A	C	/	C	R	/	C	/	/	/	A	C
RQ11	R	C	R	A	C	A	R	/	R	R	/	C	R	/	/	A	C
RQ12	/	C	R	A	C	A	R	/	C	R	/	C	R	/	/	A	C
RQ13	/	C	R	A	C	A	C	/	C	/	/	C	R	/	/	A	C
RQ14	/	C	C	A	C	A	R	/	C	/	/	R	/	/	/	A	A
RQ15	/	R	R	A	C	A	/	/	C	/	/	C	/	/	/	A	A
RQ16	/	C	R	A	C	A	/	/	R	R	/	R	/	/	/	A	A
<b>Forna Point Section</b>																	
DB1	/	R	R	A	R	C	R	/	R	/	/	R	/	/	/	A	A
DB2	/	C	R	A	R	A	/	/	C	/	/	/	R	/	/	A	A
DB3	/	C	R	A	C	A	R	/	R	R	/	R	R	/	/	A	C
DB4	/	C	C	A	C	A	C	/	R	R	/	C	R	/	/	A	C
DB5	R	C	C	C	C	C	R	/	R	/	/	C	R	/	/	A	C
DB6	/	R	R	C	R	A	R	/	C	/	/	R	/	/	/	A	R
DB7	R	C	R	A	C	C	R	/	R	C	/	C	/	/	/	A	A
DB8	R	C	R	C	C	A	C	R	C	R	/	C	/	/	/	A	A
DB9	/	R	C	C	C	A	C	/	R	/	R	C	R	/	/	A	A
DB10	/	C	R	A	R	A	R	/	R	R	/	C	R	/	/	A	A
DB11	/	R	R	C	R	C	R	/	R	/	/	R	/	/	/	A	C
DB12	R	C	R	A	C	A	/	/	C	/	/	C	R	/	/	A	C
DB13	/	R	R	A	R	C	/	/	C	/	/	R	/	/	/	A	C
DB14	R	C	C	A	R	A	C	/	C	R	R	C	/	/	/	A	A
DB15	R	C	R	A	R	A	C	/	C	/	/	C	R	/	/	A	R
DB16	/	R	C	A	C	A	R	/	R	/	/	C	R	/	/	A	C
DB17	/	C	R	C	R	C	C	/	R	/	/	C	/	/	/	A	C
DB18	R	R	R	A	R	C	C	/	C	R	/	R	/	/	/	A	C
DB19	/	C	R	A	R	A	R	/	R	/	/	C	R	/	/	A	C
<b>Wardija Point Section</b>																	
PT1	/	R	/	A	C	A	/	/	/	/	/	C	/	/	/	C	A
PT2	/	R	/	A	R	A	/	C	/	/	C	R	/	/	/	R	A
PT3	/	R	C	A	R	A	R	R	R	R	/	C	R	/	/	R	C
PT4	/	/	/	A	/	A	/	/	/	/	C	R	/	/	/	A	A
PT5	/	R	C	A	R	A	/	/	R	R	R	C	/	/	/	A	C
<b>Xlendi Section</b>																	
MT10	/	R	R	A	R	C	R	/	R	R	/	/	R	/	/	A	C
MT11	/	R	R	C	R	C	R	/	R	/	/	C	R	/	/	C	R
MT12	/	R	R	C	/	C	/	/	R	/	/	R	/	/	/	C	R
MT12bis	/	R	R	C	R	C	R	/	R	/	/	R	/	/	/	A	R
MT12ter	/	R	R	C	R	C	R	/	R	/	/	C	/	/	/	C	C
MT13	/	R	R	C	R	C	R	/	R	/	/	C	/	/	/	A	R
MT14	/	C	R	C	R	C	R	/	R	/	/	C	C	/	/	A	R
<b>Dahlet Qorrot Section</b>																	
DQ1	/	C	/	A	R	A	R	/	C	/	/	C	R	/	/	A	A
DQ2	/	C	R	A	C	A	/	/	C	/	/	C	/	/	/	A	C
DQ3	/	C	/	A	/	A	R	/	C	/	/	R	/	/	/	A	A
DQ4	/	C	/	A	C	A	/	/	C	R	/	R	R	/	/	A	A

## CHRONOSTRATIGRAPHY OF THE MALTESE LOWER GLOBIGERINA LIMESTONE MEMBER

<b>Gozo Island</b>	<i>Coccolithus miopelagicus</i>	<i>Coccolithus pelagicus</i>	<i>Cyclicargolithus abisectus</i>	<i>Cyclicargolithus floridanus</i>	<i>Dictyococcites bissectus</i>	<i>Dictyococcites scrippae</i>	<i>Helicosphaera euphratis</i>	<i>Helicosphaera obliqua</i>	<i>Pontosphaera multipora</i>	<i>Sphenolithus ciperensis</i>	<i>Sphenolithus conicus</i>	<i>Sphenolithus dissimilis</i>	<i>Sphenolithus moriformis</i>	<i>Sphenolithus pseudoradians</i>	<i>Triquetrorhabdulus carinatus</i>	<i>Umbilicosphaera sibogae</i>	<i>Zygrolithus bijugatus</i>
<b>Ras II-Hobz Section</b>																	
HO1	R	C	C	A	R	A	/	/	R	/	/	/	R	/	/	C	R
HO2	/	R	C	A	R	A	/	/	R	/	/	R	C	/	/	C	C
HO3	R	R	C	A	/	A	R	/	R	/	/	R	R	/	/	C	C
HO4	R	R	C	A	R	A	R	/	R	/	/	R	/	/	/	A	C
HO5	R	C	R	A	/	A	/	/	R	/	/	C	R	/	/	A	C
<b>Santa Lucija Section</b>																	
SL1	/	C	C	A	/	A	/	/	R	/	/	C	C	/	/	C	A
SL2	R	A	R	A	R	A	R	/	/	/	/	C	/	/	/	C	C
SL3	/	C	/	A	/	A	R	/	R	/	/	R	R	/	/	C	A
SL4	R	C	/	A	R	A	R	/	/	/	/	R	R	/	/	A	C
SL5	/	R	C	A	R	A	/	/	R	/	/	R	R	/	/	A	C
SL6	R	C	C	A	C	A	R	/	/	/	/	C	C	/	/	C	C
SL7	/	C	C	A	R	A	/	/	R	/	/	R	R	/	/	A	C
SL8	/	C	R	A	R	A	/	/	R	/	/	C	C	/	/	C	A

Table 6. Semiquantitative calcareous nannofossil content of the analyzed sections of Lower Globigerina Limestone member on Gozo

	<i>Braarudosphaera bigelovii</i>	<i>Chiasmolithus altus</i>	<i>Coccolithus miopelagicus</i>	<i>Coccolithus pelagicus</i>	<i>Cyclicargolithus abisectus</i>	<i>Cyclicargolithus floridanus</i>	<i>Dicytoccites bisectus</i>	<i>Dicytoccites scriptae</i>	<i>Helicosphaera euphratis</i>	<i>Helicosphaera obliqua</i>	<i>Helicosphaera recta</i>	<i>Pontosphaera enormis</i>	<i>Pontosphaera multipora</i>	<i>Sphenolithus ciperoensis</i>	<i>Sphenolithus delphix</i>	<i>Sphenolithus dissimilis</i>	<i>Sphenolithus moriformis</i>	<i>Umbilicosphaera sibogae</i>	<i>Zygrabilithus bijugatus</i>
<b>Qammieh Section</b>																			
MH5	0.00	0.00	0.00	3.92	1.85	10.18	1.85	8.33	0.00	0.00	0.00	0.00	9.25	0.00	0.00	0.00	26.62	37.96	
MH6	0.00	0.00	1.58	5.55	1.58	33.33	2.38	23.8	0.00	0.00	0.00	0.00	4.76	1.19	0.00	12.30	1.58	6.34	5.55
MH7	0.00	0.00	0.00	0.00	0.00	28.36	2.83	22.69	0.00	0.00	0.70	0.00	5.67	0.00	0.00	1.41	2.83	18.49	17.02
MH8	0.00	0.00	0.00	4.08	3.57	35.71	8.16	20.40	0.51	0.00	0.00	0.00	4.08	0.00	0.00	3.06	2.04	3.06	15.30
MH9	0.47	0.47	0.47	0.94	3.79	34.19	3.31	27.48	1.89	0.00	0.00	1.42	1.89	0.94	0.00	2.84	0.94	11.37	8.53
<b>Fomm ir-Rib Section</b>																			
FIR1	0.00	0.00	0.00	5.04	0.00	23.26	4.95	16.78	0.00	0.00	0.00	0.00	2.33	0.00	0.00	4.45	2.43	20.03	20.69
FIR2	0.00	0.00	0.00	0.00	0.00	30.84	5.40	13.18	0.00	0.00	0.00	0.00	6.78	0.00	0.00	2.67	1.10	21.89	18.10
FIR3	0.00	0.00	0.00	1.69	1.21	24.84	0.48	21.60	0.00	0.00	0.00	0.00	3.88	0.00	0.00	2.66	1.85	23.54	18.20
FIR4	0.00	0.00	0.00	8.48	1.36	25.99	0.00	17.91	1.58	1.81	0.00	0.00	2.04	0.00	0.00	3.40	3.26	21.76	15.87
FIR5	0.00	0.00	0.00	8.48	1.71	31.29	0.76	13.74	1.92	0.00	0.00	0.00	3.81	0.38	0.00	2.29	1.55	22.80	11.25
FIR6	0.00	0.00	0.00	10.1	2.94	27.89	2.31	16.06	1.89	1.47	1.05	0.00	3.59	0.00	0.00	3.15	0.68	15.96	12.60
FIR7	0.00	0.00	0.00	7.81	2.60	22.17	4.50	13.27	0.00	2.13	2.92	0.00	3.75	0.00	0.00	4.97	2.65	18.48	14.69
FIR8	0.00	0.00	0.00	2.99	0.00	22.65	0.00	12.82	0.00	0.00	0.00	0.00	6.41	0.00	0.00	1.68	0.00	30.76	22.64
FIR9	0.00	0.00	0.00	2.26	0.00	26.24	0.00	11.76	0.00	0.00	0.00	0.00	7.69	0.00	0.00	1.80	0.00	29.86	20.36
<b>II Blata Section</b>																			
BL1	0.00	0.00	1.04	8.33	4.16	32.81	0.00	30.20	0.00	0.00	0.00	0.00	1.56	0.00	6.24	0.00	5.20	10.41	
BL2	0.00	0.00	0.85	8.20	3.50	31.25	1.40	30.00	0.00	0.00	0.00	0.00	1.35	0.54	7.06	0.96	6.79	8.15	
BL3	0.00	0.00	0.00	7.87	3.08	27.61	1.63	34.90	0.00	0.00	0.00	0.00	1.35	0.54	7.06	0.96	6.79	8.15	
BL4	0.00	0.00	0.40	6.50	3.30	24.75	1.55	27.95	0.00	0.00	0.00	0.00	1.10	0.00	7.25	0.00	12.65	14.55	
BL5	0.00	0.00	0.00	3.26	4.71	22.82	0.72	23.55	0.00	0.00	0.00	0.00	1.44	0.36	9.42	1.45	13.40	18.84	
<b>Huttaf Gandolf Section</b>																			
MQ1	0.00	0.00	0.00	7.60	3.80	32.60	10.86	26.08	0.54	1.08	0.00	0.00	0.54	1.63	0.00	6.52	1.63	5.97	1.08
MQ2	0.00	0.00	0.00	4.68	0.00	35.43	4.16	34.37	2.08	0.00	0.00	0.00	0.00	0.00	0.00	6.25	0.00	6.25	6.77
MQ3	0.00	0.00	0.00	8.86	0.00	26.58	5.06	30.14	0.00	0.00	0.00	0.00	1.26	0.00	7.59	0.00	2.53	17.98	
MQ4	0.00	0.00	0.00	5.00	0.83	30.00	2.50	21.66	0.83	0.83	0.00	0.00	1.66	0.00	5.00	0.00	18.51	34.25	
MQ5	0.00	0.00	0.00	0.00	0.00	22.22	10.18	12.03	2.77	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	18.51	34.25
MQ6	0.00	0.00	0.00	2.85	0.95	30.00	8.09	18.09	0.00	1.90	0.00	1.42	3.80	0.95	0.00	3.33	0.00	14.28	13.33
MQ7	0.00	0.00	0.00	5.04	2.36	26.82	4.20	21.21	0.00	0.84	0.00	2.52	3.04	1.68	0.00	4.04	0.00	18.06	10.10
MQ8	0.00	0.00	0.00	1.95	0.00	5.12	0.00	2.56	0.00	0.00	0.00	2.12	11.82	2.84	0.00	2.56	0.00	41.02	28.64
<b>Wied id-Dis Section</b>																			
WD1	0.00	0.00	0.00	13.27	2.16	21.60	9.41	27.20	6.63	0.00	0.00	0.00	3.54	0.77	0.00	4.16	1.54	7.40	2.31
WD2	0.00	0.00	0.00	0.94	1.03	25.78	7.54	31.13	2.83	0.00	0.00	0.00	3.77	0.00	0.00	1.88	1.14	18.23	5.66
WD3	0.00	0.00	0.00	1.33	1.33	32.00	9.00	26.33	1.66	0.00	0.00	0.00	3.33	0.00	0.00	2.00	0.66	16.00	6.33
WD4	0.00	0.00	0.00	2.24	0.93	31.46	6.17	32.20	1.49	0.74	0.00	0.00	1.31	0.00	0.00	2.62	0.56	14.60	2.62
WD5	0.00	0.00	0.00	1.77	1.06	23.84	6.40	27.75	3.20	0.00	0.00	0.00	5.33	1.06	0.00	2.13	0.00	21.35	6.04
WD6	0.00	0.00	0.00	1.24	0.31	25.23	3.11	28.03	1.55	0.00	0.00	0.00	2.80	0.31	0.00	3.12	0.00	24.29	9.34
WD7	0.00	0.00	0.00	3.70	0.52	33.86	6.34	22.75	3.17	0.00	0.00	0.00	3.17	0.79	0.00	4.76	0.79	16.40	3.70
WD8	0.00	0.00	0.00	3.61	0.00	32.54	4.71	26.10	2.98	0.00	0.00	0.00	1.57	0.78	0.00	5.18	1.25	13.67	7.54
WD9	0.00	0.00	0.00	1.72	0.00	26.60	4.18	27.17	1.92	0.00	0.00	0.00	3.20	0.49	0.00	2.21	0.49	17.73	14.28
WD10	0.00	0.00	0.00	0.00	0.00	28.57	1.86	19.25	0.00	0.00	0.00	0.00	2.48	0.00	0.00	3.72	1.86	26.08	16.14
WD11	0.00	0.00	0.00	4.07	0.00	29.51	6.99	33.59	2.52	0.00	0.00	0.00	1.94	0.77	0.00	2.71	0.00	10.29	7.57
WD12	0.00	0.00	0.00	2.25	0.00	39.00	4.00	30.00	1.00	0.00	0.00	0.00	1.50	0.00	0.00	1.76	0.00	13.00	7.50
WD13	0.00	0.00	0.00	4.51	0.00	27.06	6.76	28.57	2.63	0.00	0.00	0.00	3.38	0.00	0.00	0.00	0.00	15.78	11.27
WD14	0.00	0.00	0.00	2.12	0.00	19.14	0.00	19.14	0.00	0.00	0.00	0.00	0.70	0.00	0.00	0.00	0.00	36.17	32.62
<b>Reqqa Point Section</b>																			
RQ6	0.00	0.00	0.00	3.84	1.44	28.84	4.32	21.15	0.96	0.00	0.00	0.00	6.73	0.00	0.00	0.00	0.00	18.26	14.42
RQ7	0.00	0.00	0.00	0.62	1.86	28.34	3.73	28.03	0.00	0.00	0.00	0.00	3.11	0.00	0.00	2.80	0.00	19.62	11.83
RQ8	0.00	0.00	0.00	6.35	3.34	30.10	5.01	16.38	2.00	0.00	0.00	0.00	4.01	0.00	0.00	3.01	1.33	15.71	11.70
RQ9	0.00	0.00	0.00	5.70	3.08	28.74	4.03	21.37	0.95	0.00	0.00	0.00	7.36	0.00	0.00	4.98	0.71	14.25	8.78
RQ10	0.00	0.00	1.84	8.29	2.99	29.03	5.52	18.20	3.45	0.00	0.00	0.00	5.06	0.69	0.00	3.22	0.00	14.97	6.68
RQ11	0.00	0.00	2.18	8.29	2.18	28.82	6.76	19.86	1.74	0.00	0.00	0.00	2.83	0.65	0.00	5.67	1.09	12.22	7.64
RQ12	0.00	0.00	0.00	8.59	1.95	25.19	6.44	22.65	2.53	0.00	0.00	0.00	4.68	0.78	0.00	6.05	0.58	12.50	8.00
RQ13	0.00	0.00	0.00	8.35	1.87	29.01	6.05	21.50	3.54	0.00	0.00	0.00	3.54	0.00	0.00	3.54	0.20	16.07	6.26
RQ14																			

## CHRONOSTRATIGRAPHY OF THE MALTESE LOWER GLOBIGERINA LIMESTONE MEMBER

	<i>Braarudosphaera bigelowii</i>	<i>Chiasmolithus altus</i>	<i>Coccolithus miopelagicus</i>	<i>Coccolithus pelagicus</i>	<i>Cyclicargolithus abisectus</i>	<i>Cyclicargolithus floridanus</i>	<i>Dictyococcites bisectus</i>	<i>Dictyococcites scrippsae</i>	<i>Helicosphaera euphratia</i>	<i>Helicosphaera obliqua</i>	<i>Helicosphaera recta</i>	<i>Pontosphaera enormis</i>	<i>Pontosphaera multipora</i>	<i>Sphenolithus ciperoensis</i>	<i>Sphenolithus delphix</i>	<i>Sphenolithus dissimilis</i>	<i>Sphenolithus moriformis</i>	<i>Umbilicosphaera sibogae</i>	<i>Zygrolithus bijugatus</i>
<b>Wardija Point Section</b>																			
PT1	0.00	0.00	0.00	2.50	0.00	25.00	5.00	20.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	5.00	0.00	9.20	25.60
PT2	0.00	0.00	0.00	2.40	0.00	44.00	2.40	17.00	0.01	4.00	0.00	0.00	0.00	0.00	0.00	9.00	2.80	2.00	16.40
PT3	0.00	0.00	0.00	2.70	6.00	36.00	0.90	32.70	0.90	0.90	0.00	0.00	0.90	0.90	0.00	9.01	1.15	1.60	6.05
PT4	0.00	0.00	0.00	0.00	0.00	14.30	0.00	15.30	0.00	0.00	0.00	0.00	0.00	0.00	0.00	9.50	3.10	14.60	43.50
PT5	0.00	0.00	0.00	2.70	3.70	22.70	1.80	24.60	0.00	0.00	0.00	0.00	0.90	0.90	0.90	9.20	0.00	15.70	5.80

Table 7. Quantitative calcareous nannofossil content (percentages) of Lower Globigerina Limestone member in selected sections of the Maltese Archipelago